

# Observatoire de Corinthe

## Corinth Rift Laboratory (CRL)

P. Bernard, H. Lyon-Caen, A. Deschamps, S. Lambotte, P. Briole, A. Nercessian, M. Aissaoui, M. Grunberg, M. Ford, Ch. Beck, A. Hubert-Ferrari, F. Cornet, P. Elias, M. Vidal, N. Meyer, A. Canitano, M. Godano, O. Scotti, A. Boiselet, S. El Arem, C. Duverger, P. Dublanchet, A. Serpetsidaki, E. Sokos, P. Papadimitriou, V. Kapetanidis, N. Voulgaris, P. Albini, D. Katsonopoulou, J. Zahradnick, K. Makropoulos, A. Tselentis,  
et al ...

France : IPGP, ENS, GEOAZUR, EOST, CRPG, IRSN, Ecole Mines, Univ. Montpellier, ISTerre ....

Greece: NKUA, Univ. Patras, NOA

Belgium : Univ. Liège

Italy : INGV (Rome, Milano, Napoli)

Czech Rep., : Charles Univ. Prague

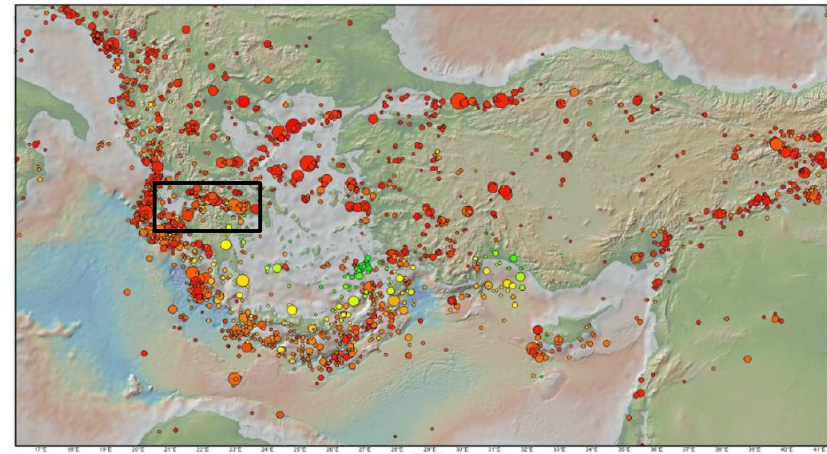
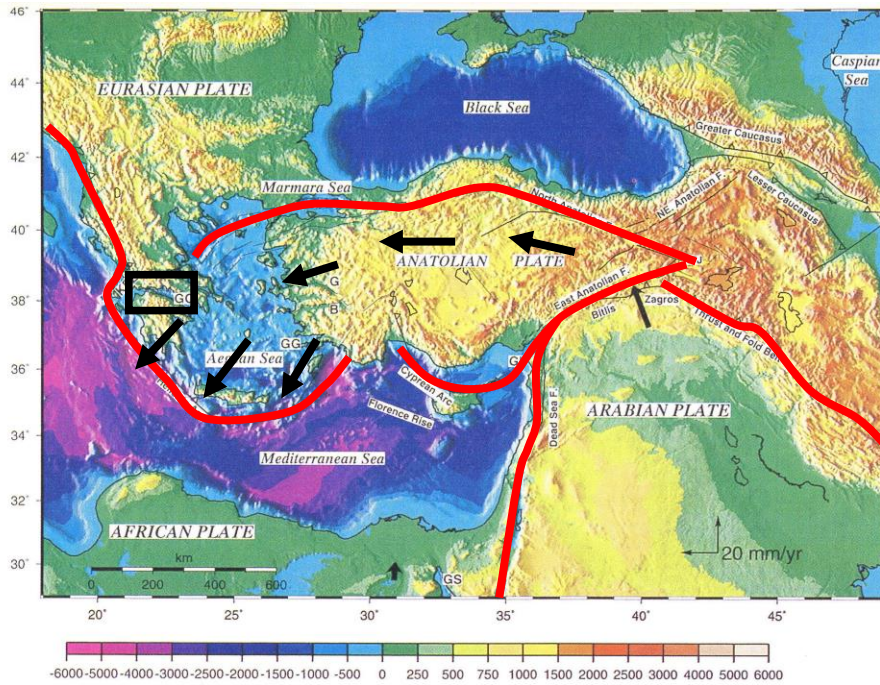
<http://www.crlab.eu>

**EC projects:**

**... CORSEIS, 3HAZ, REAKT, SERA**

**French ANR SISCOR**

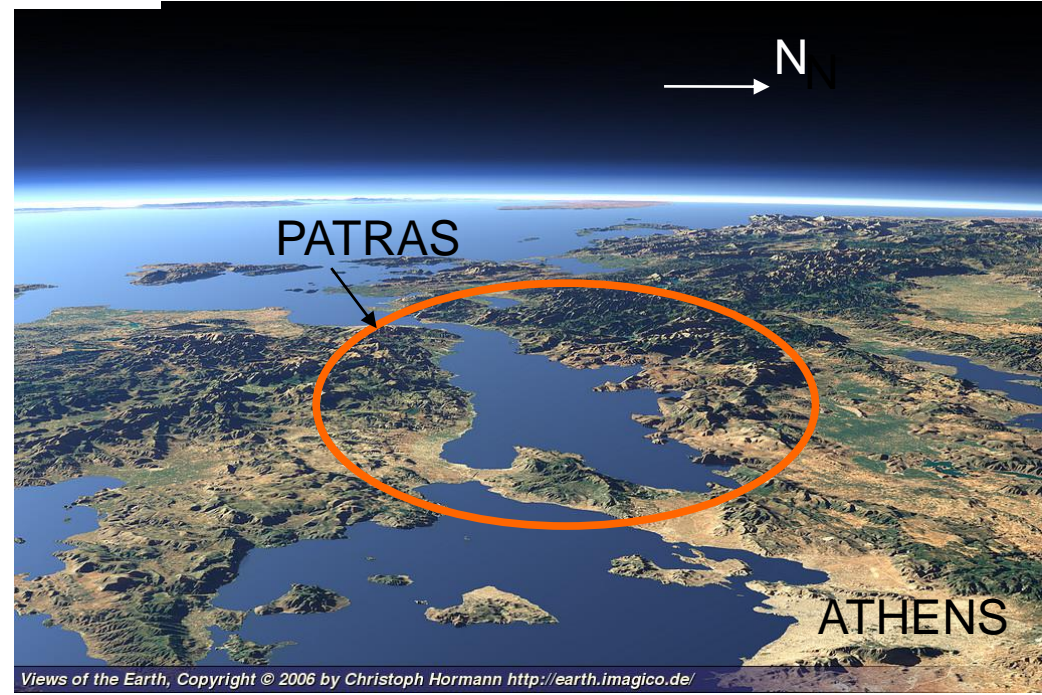


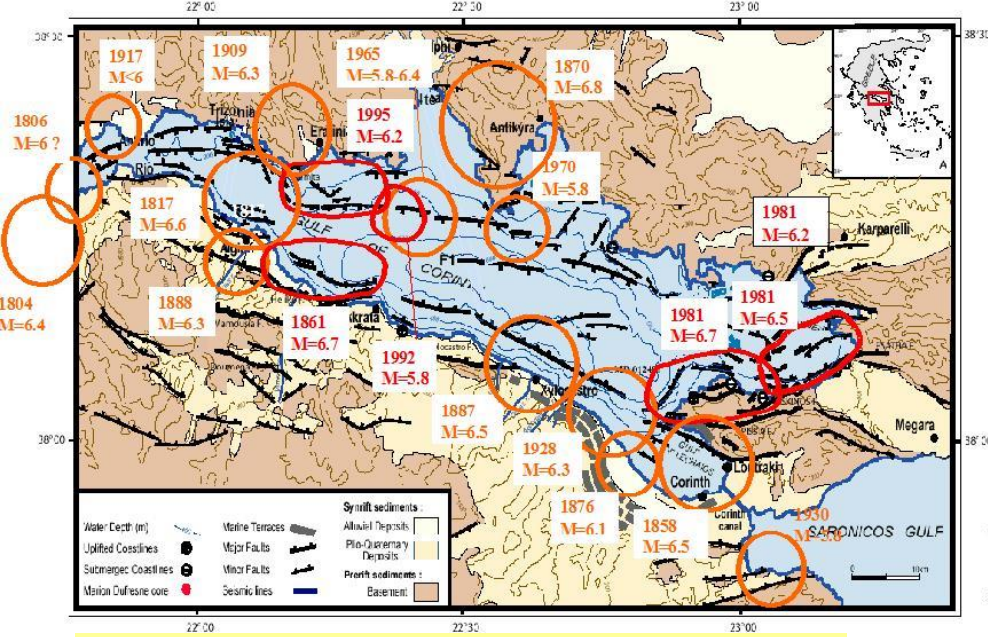


USGS-NEIC  
1970+ M>4.5

## The rift of Corinth:

- Extension of the Aegean plate due to the roll-back of the subducting African plate
- connection with northern branch of the North Anatolian Fault
- Associated to a large instrumental and historical earthquakes –  $M \sim 6.5-7$

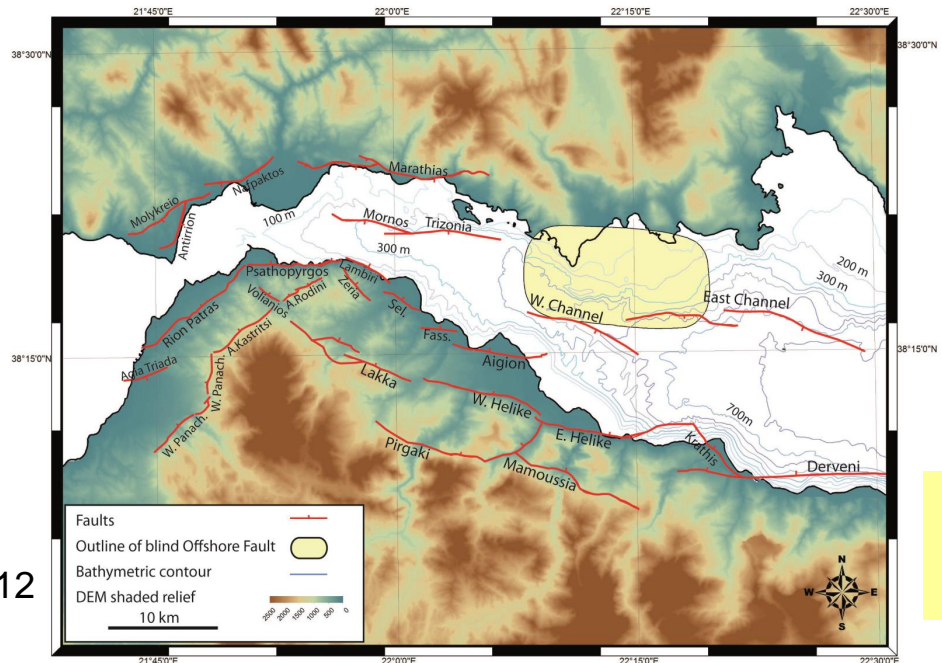




Major Earthquakes since 1800



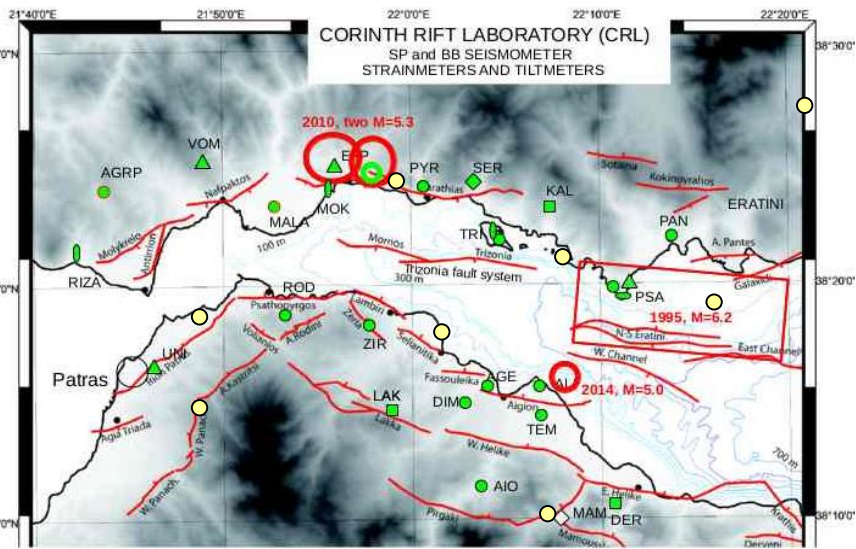
Helike fault scarp



Ford, Meyer, 2012  
- ANR SISCOR

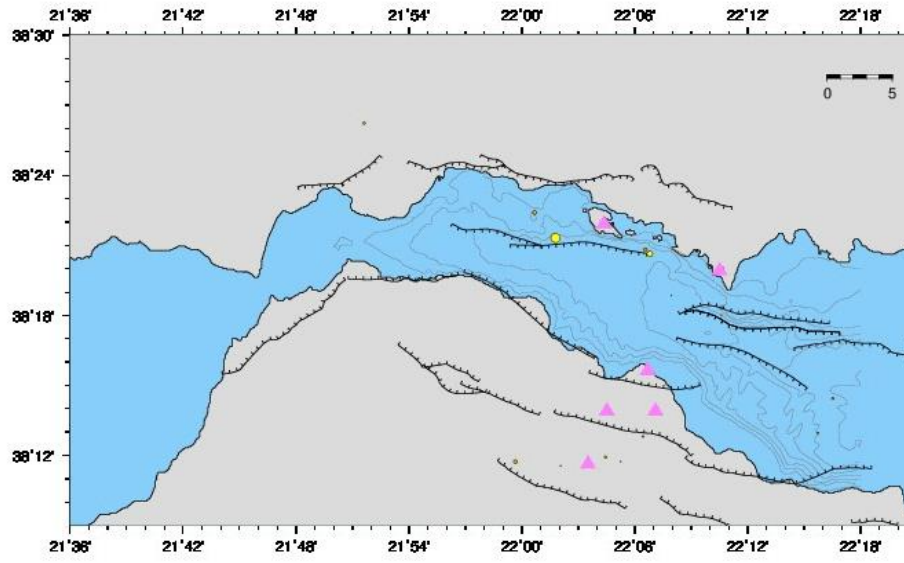


Main faults in the western rift of Corinth



- NKUA
- ◆ NOA
- ▲ PATRAS
- ◆ PRAGUE
- FRANCE CNRS
- Seismometers
- Seismic antenna
- strainmeter
- Tiltmeter
- ▲ MT
- cGPS

Automatically determined May 2000



Seismicity 2001-2013

- 1995 Aigion earthquake, M=6.2, low-dip, blind normal fault

**CRL starts in 2000, seismology  
+ multiparameter monitoring (GPS, strain, MT, ...)**

**Seismicity 2000-2014 : ~ 10 000 EQ/year – large swarms, days to months**

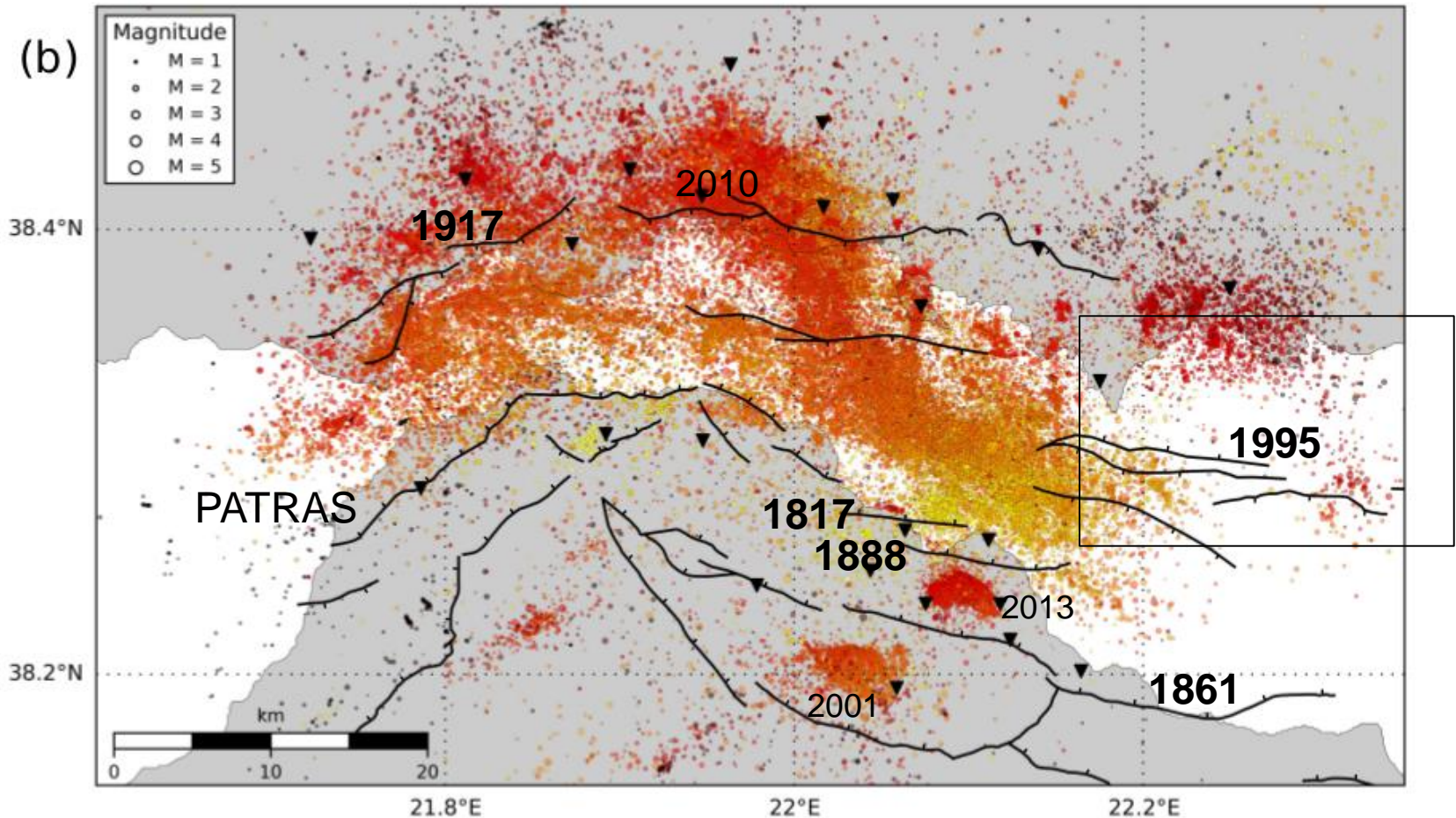
**M compl. ~ 1.5**

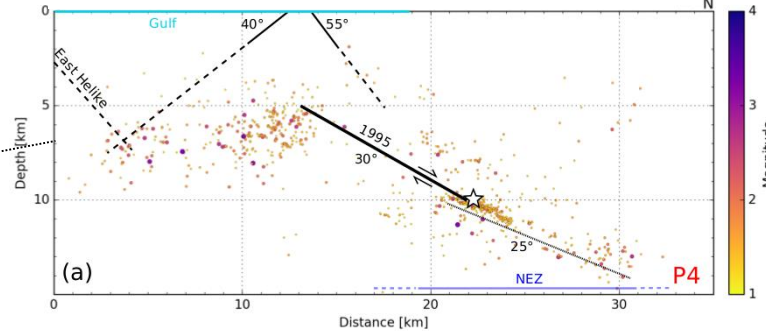
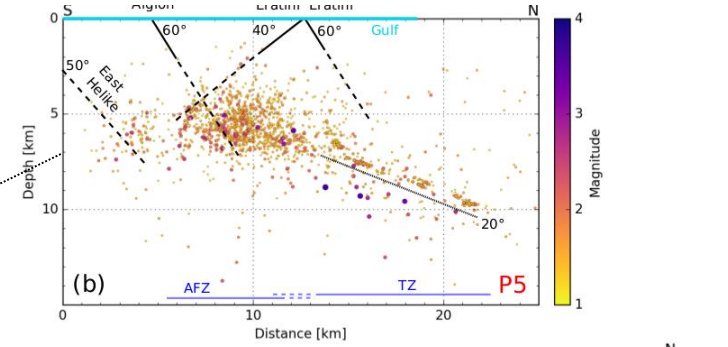
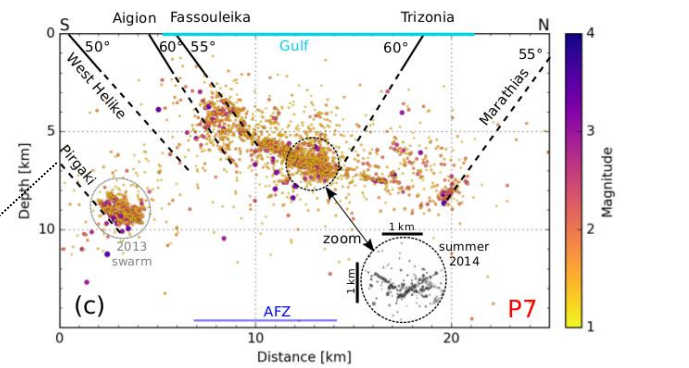
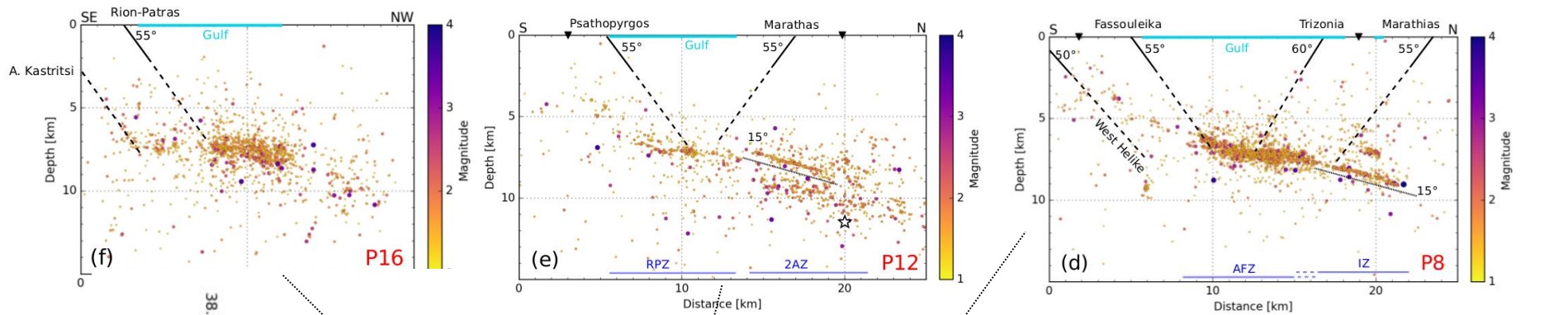
**two M=5.3, 2010**

**M=5.0 , 2014**

# Relocated seismicity 2000-2015

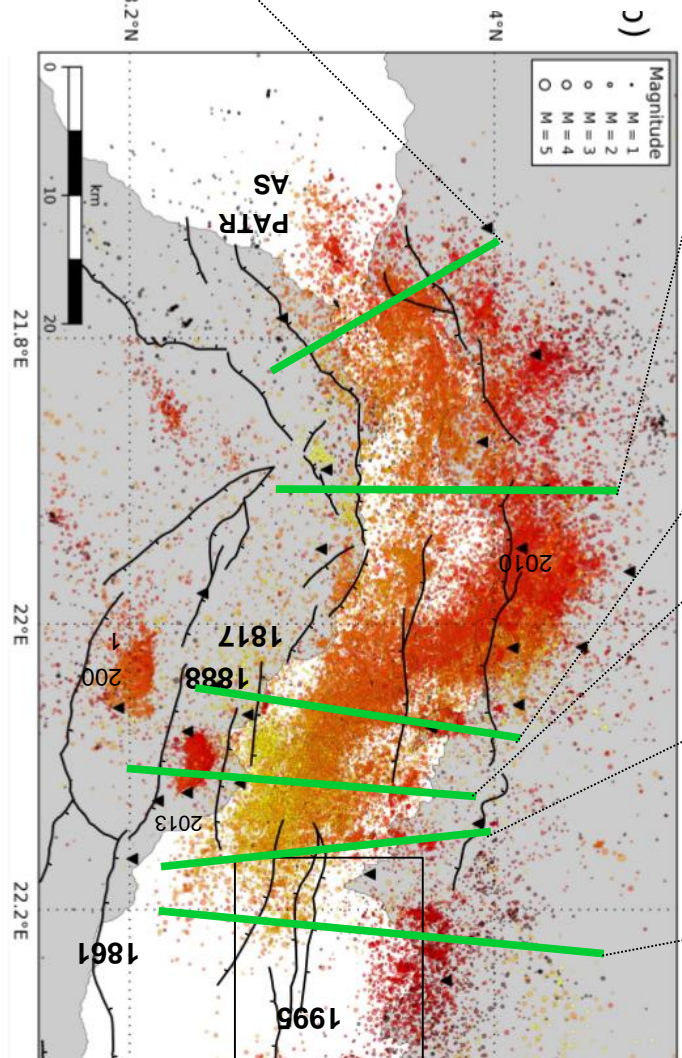
HypoDD –  
200,000 evts





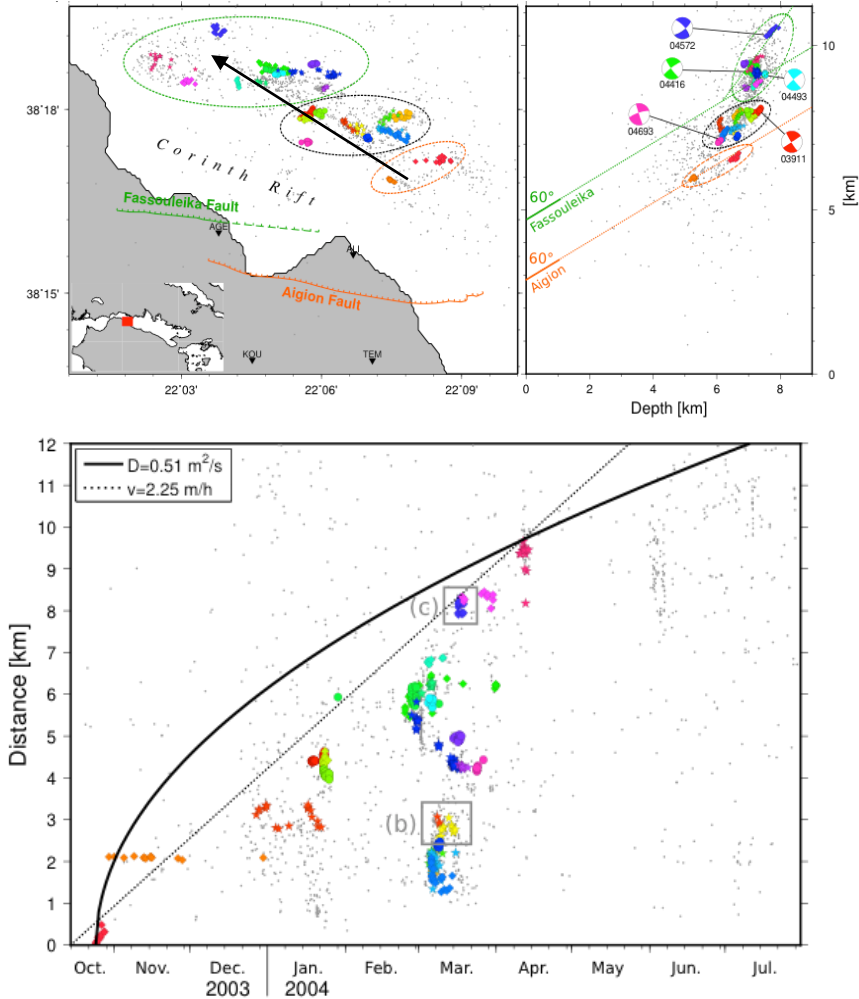
# Seismicity 2000-2015

Duverger et al.  
2018



# Space-Time characteristics of microseismicity: Swarms and migration

## Swarm 2004



Duverger et al. 2015

## Swarm 2014

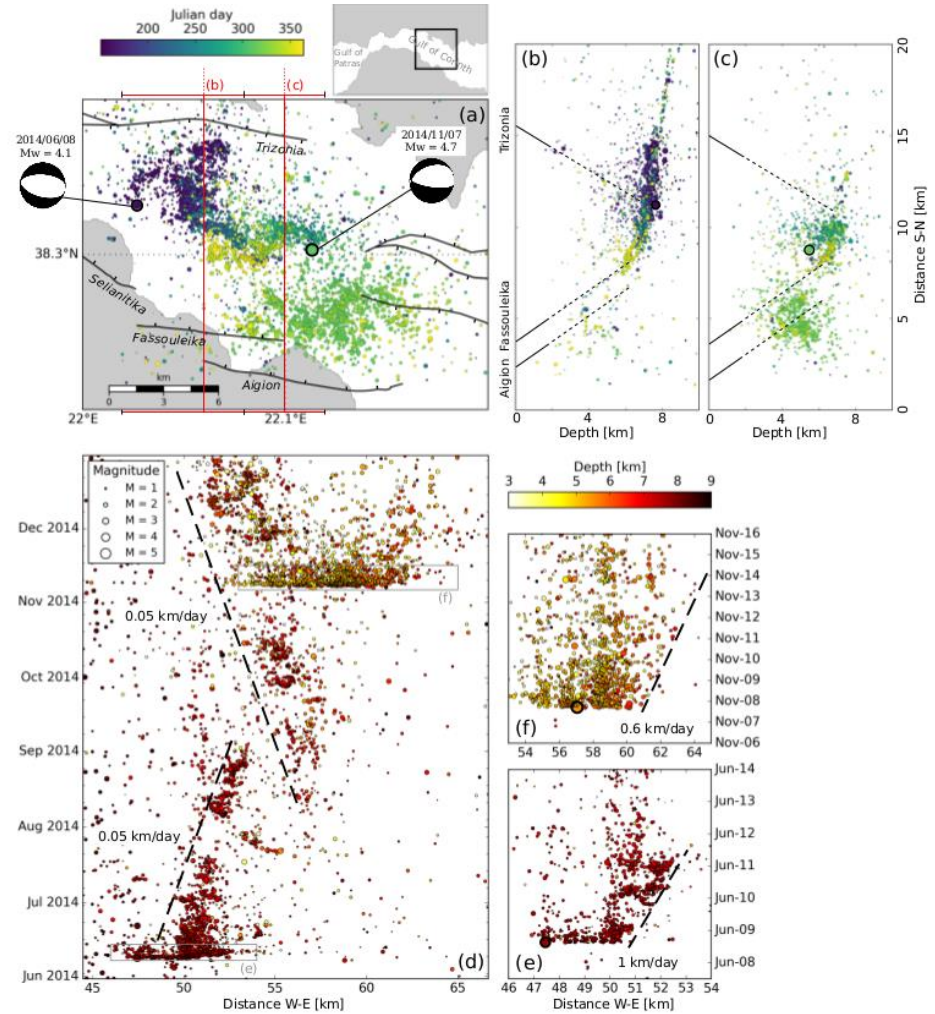
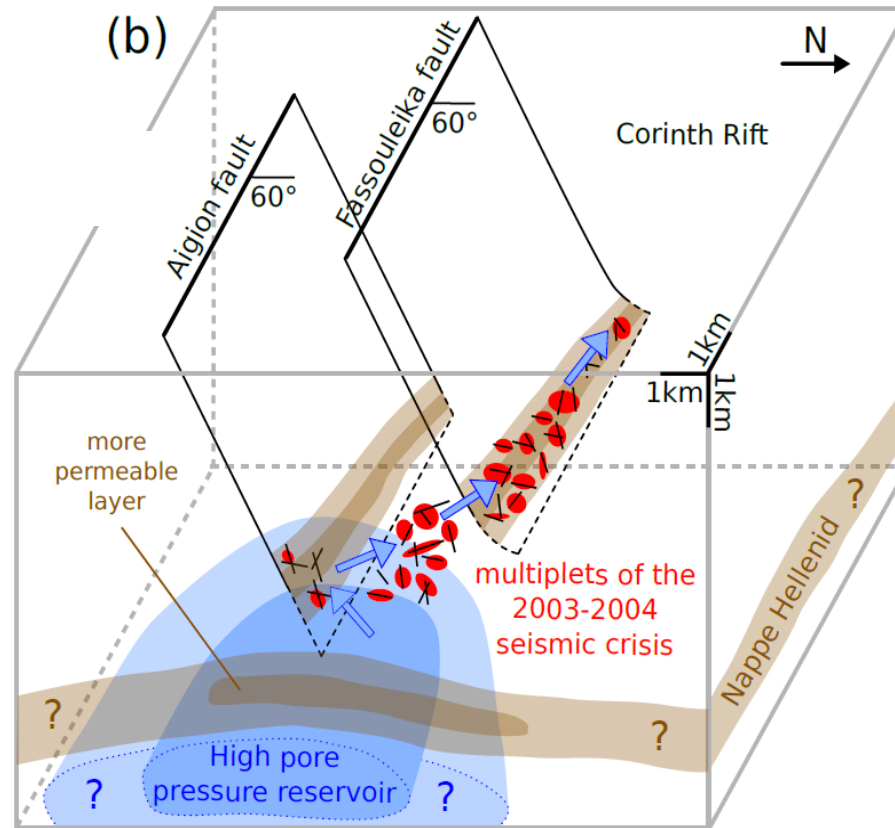


Figure 7. Seismic migration velocities during the western (ZW) swarm in 2014. (a) Map of the seismicity

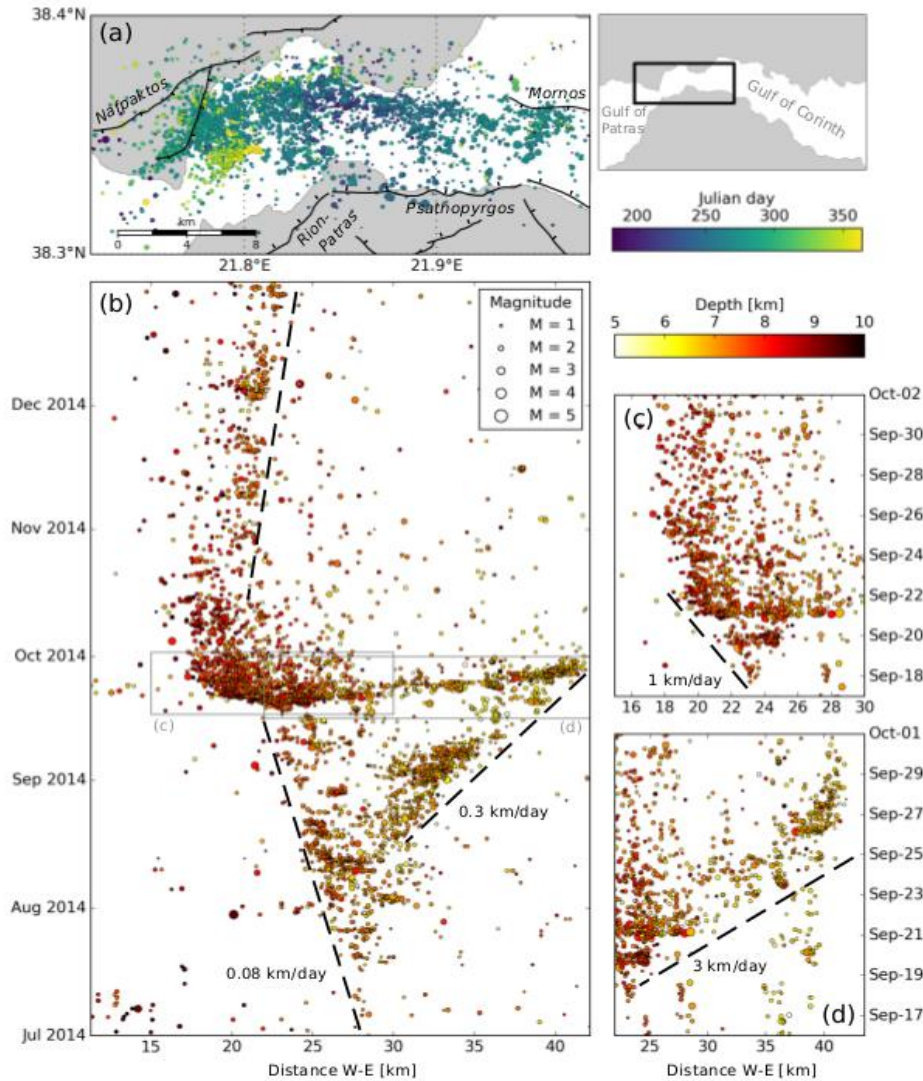
Duverger et al. 2018



Permeable corridor at the intersection of the active normal faults and a geological structure (Phyllade nappe? )



# Space-Time characteristics of microseismicity: Swarms and migration



Migration velocity < 0.5-1 km/day  
**Pore pressure** diffusion control

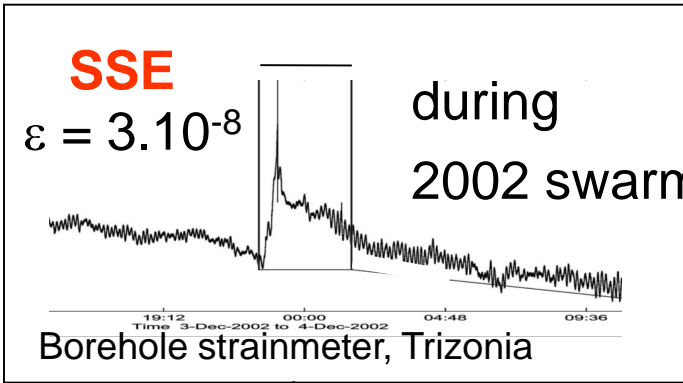
Migration velocity > 1 km/day  
**Creep** diffusion control

Swarm summer 2014

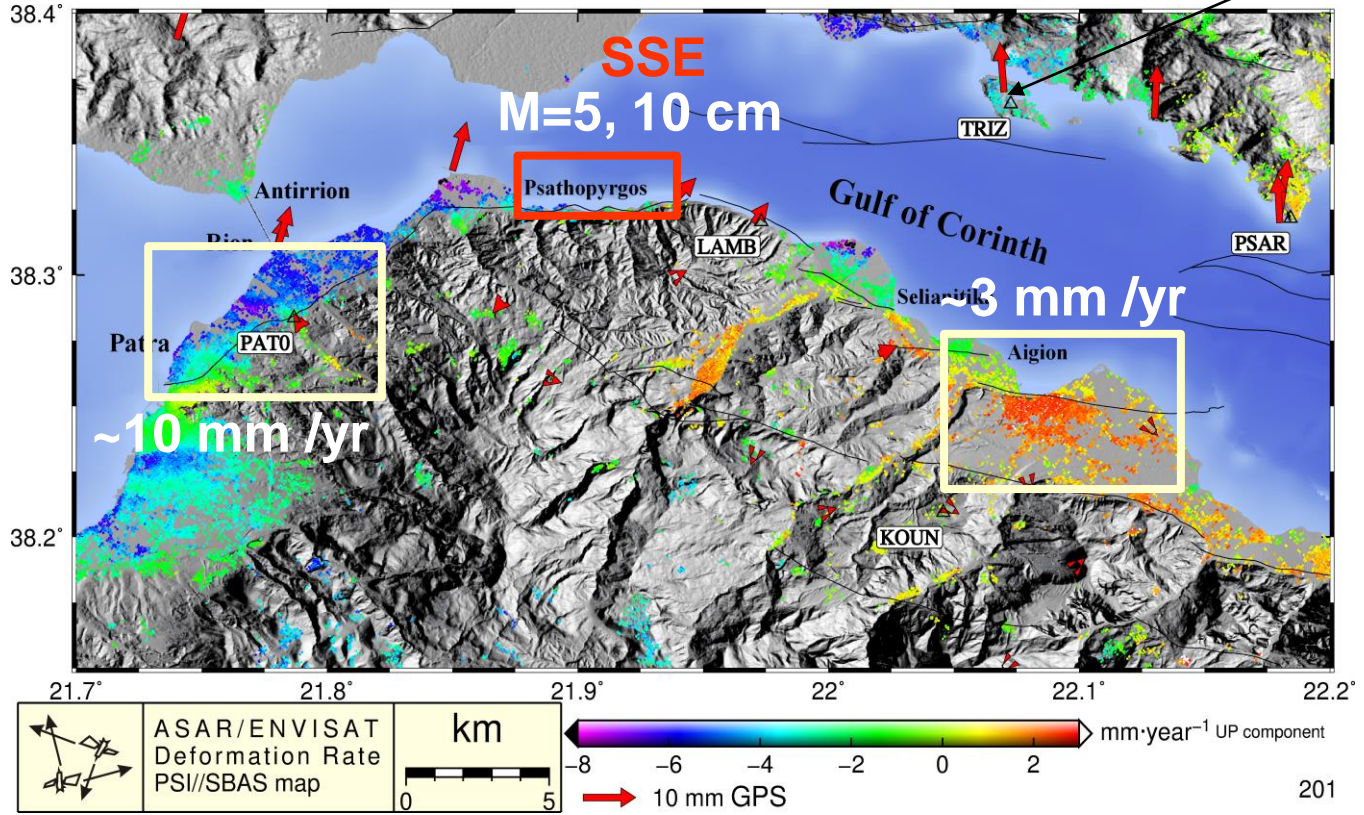
# Aseismic creep on major faults

from GPS, InSAR, and borehole strainmeters

Psathopyrgos (transient creep), Rio-Patras, and Aigion faults

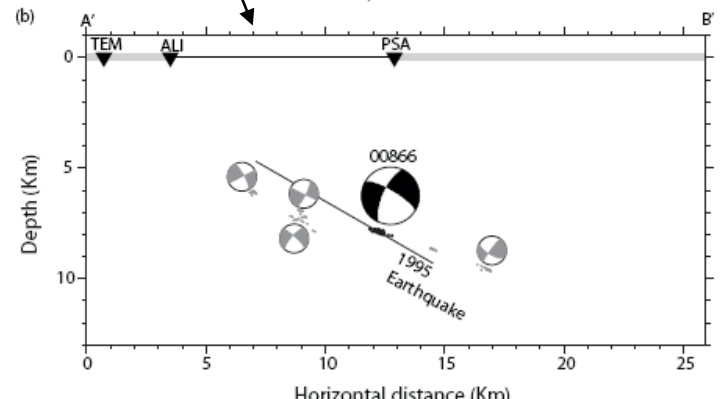
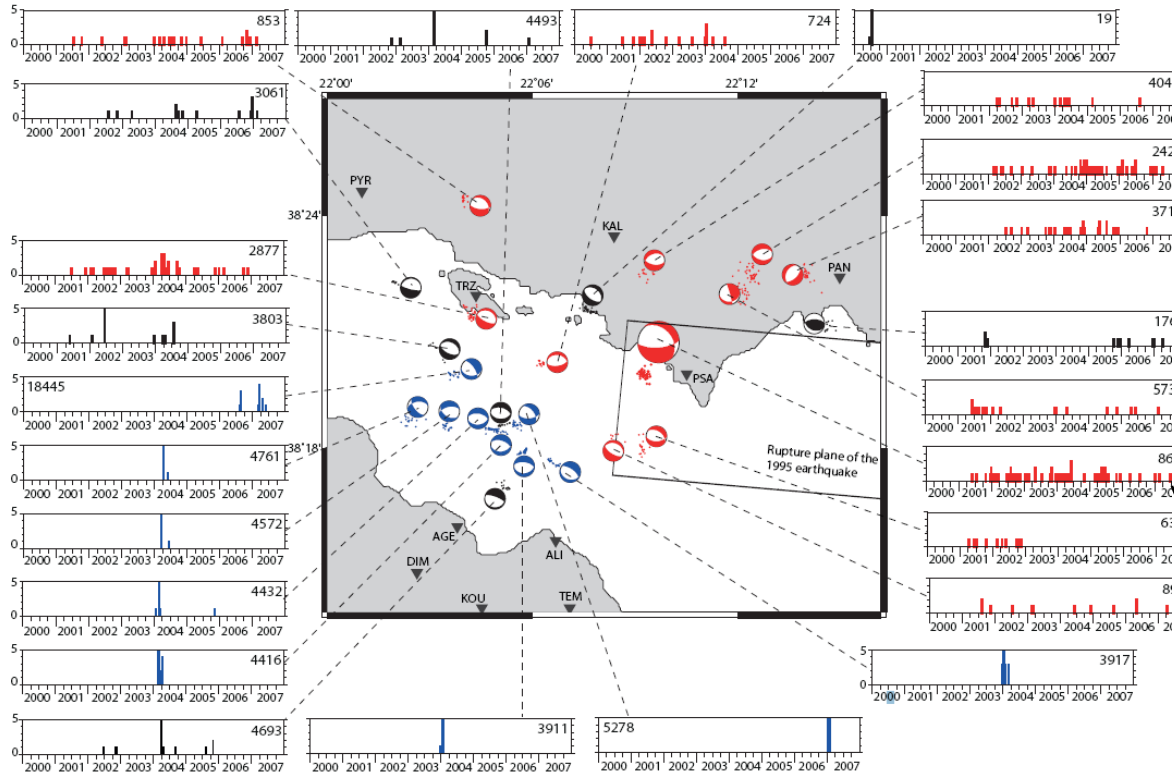


Bernard et al., 2008



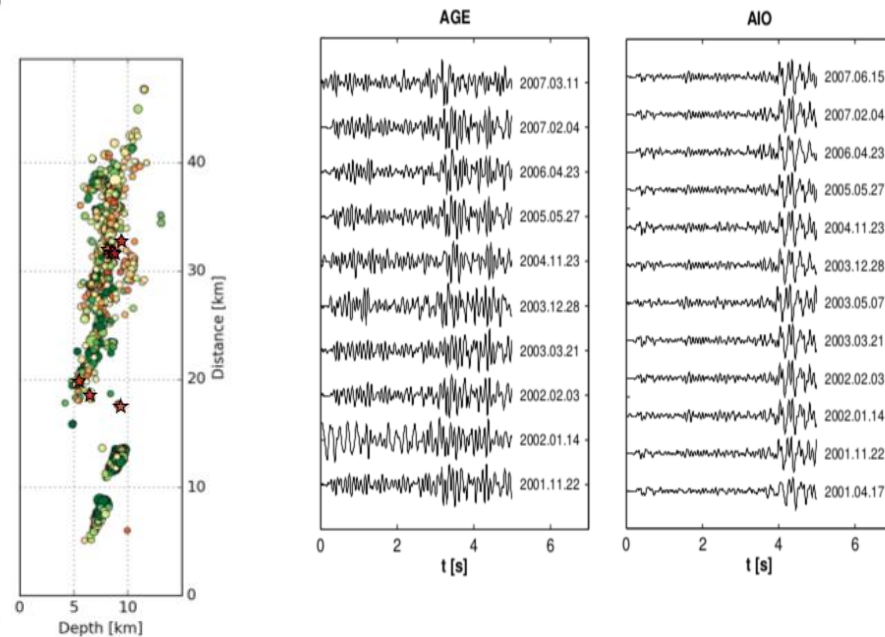
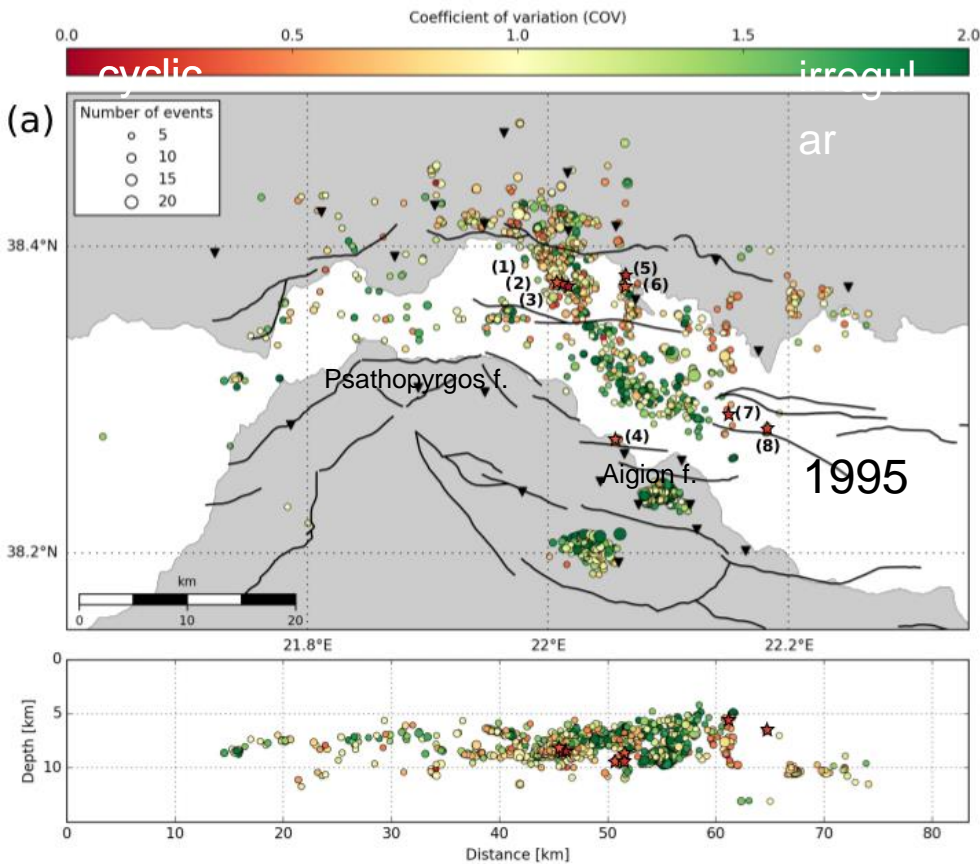
Elias and Briole, 2013

# Continuous creep on deep faults: Evidence from persistent multiplets?



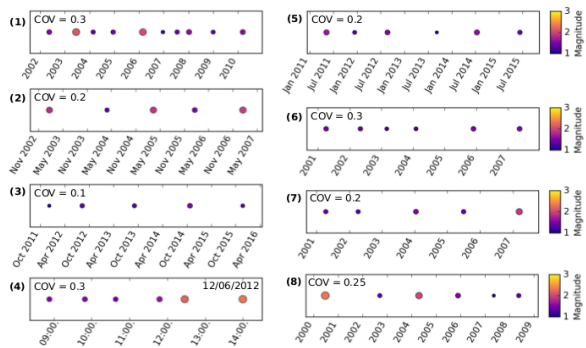
Godano et al. , 2015

# Interevent times of earthquakes within multiplets



**cyclic multiplets / repeaters**  
surrounding:

- 1995 fault zone
- deep Aigion fault zone
- deep Lambiri/Psathopyrgos fault zone

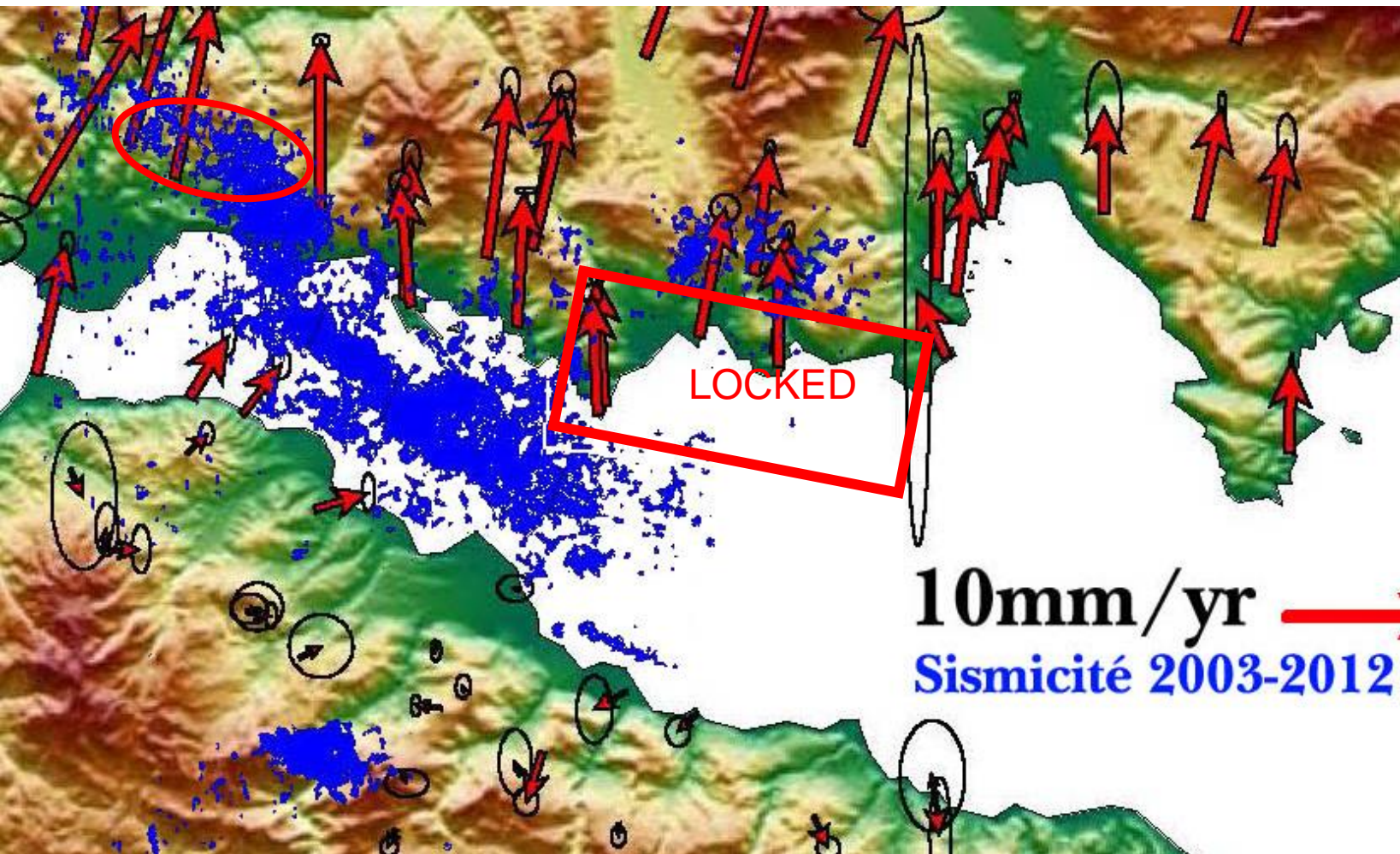


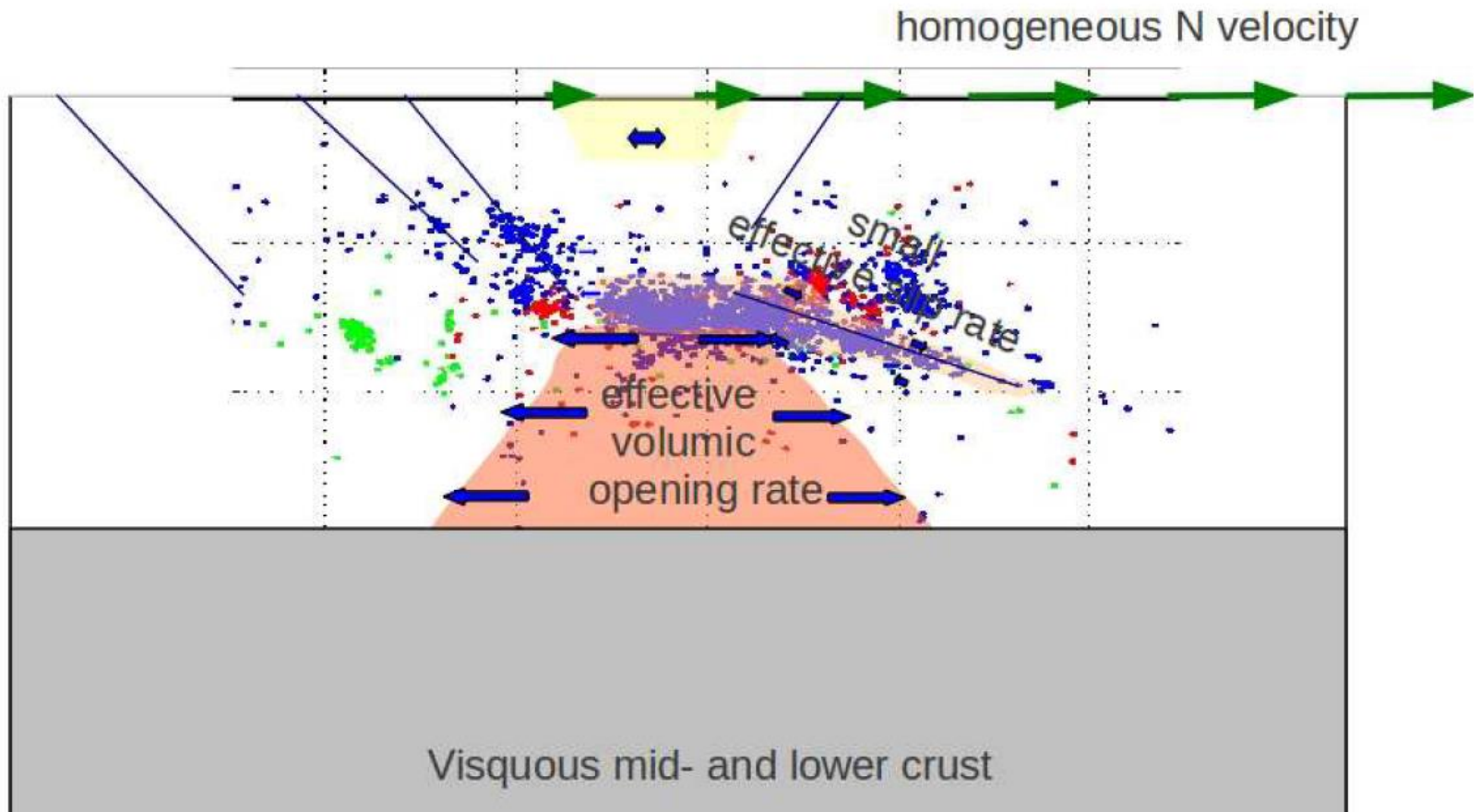
Duverger et al.  
2018

Locked fault surfaces on 1995 and 2010 rupture zones :

Does not seem to affect GPS rates –

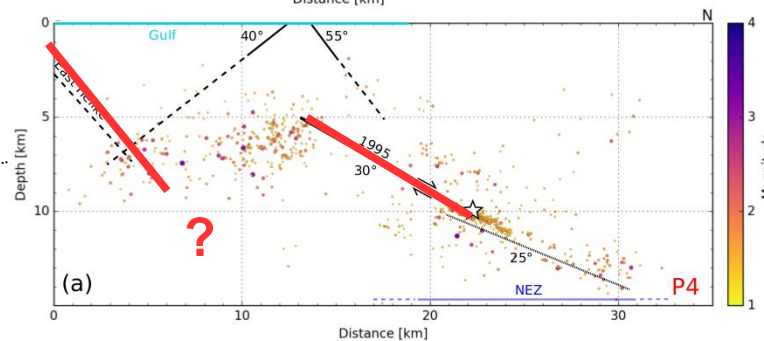
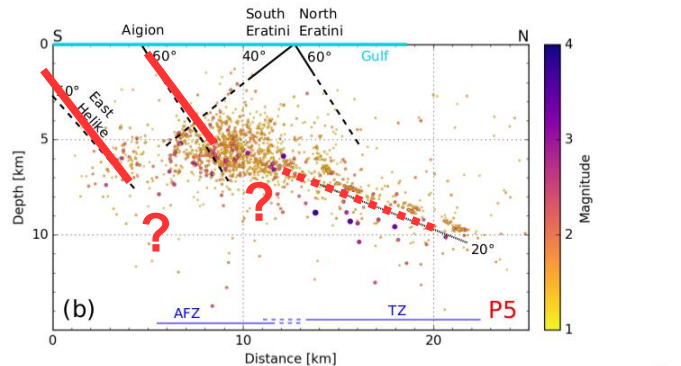
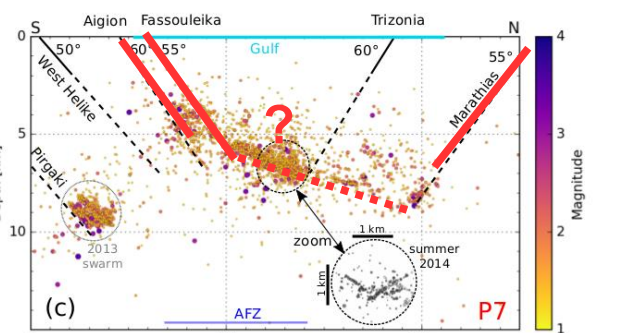
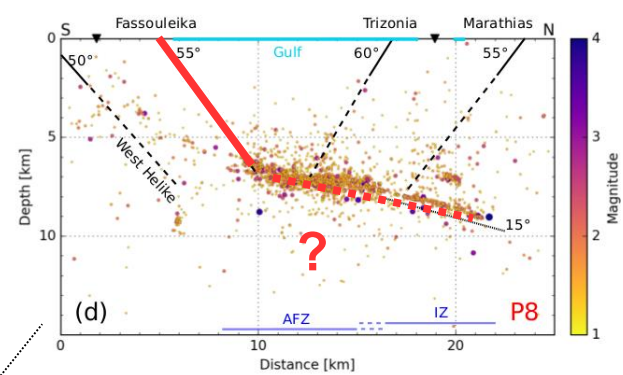
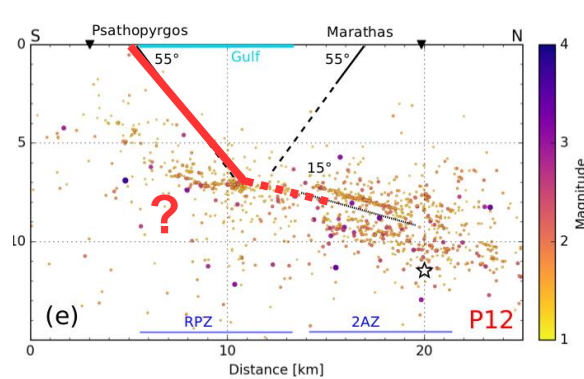
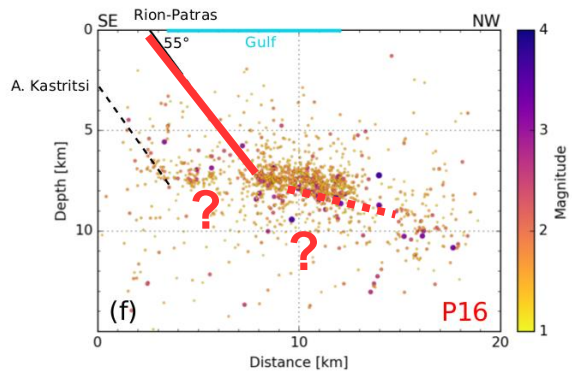
Suggests deeper or shallower interseismic sources





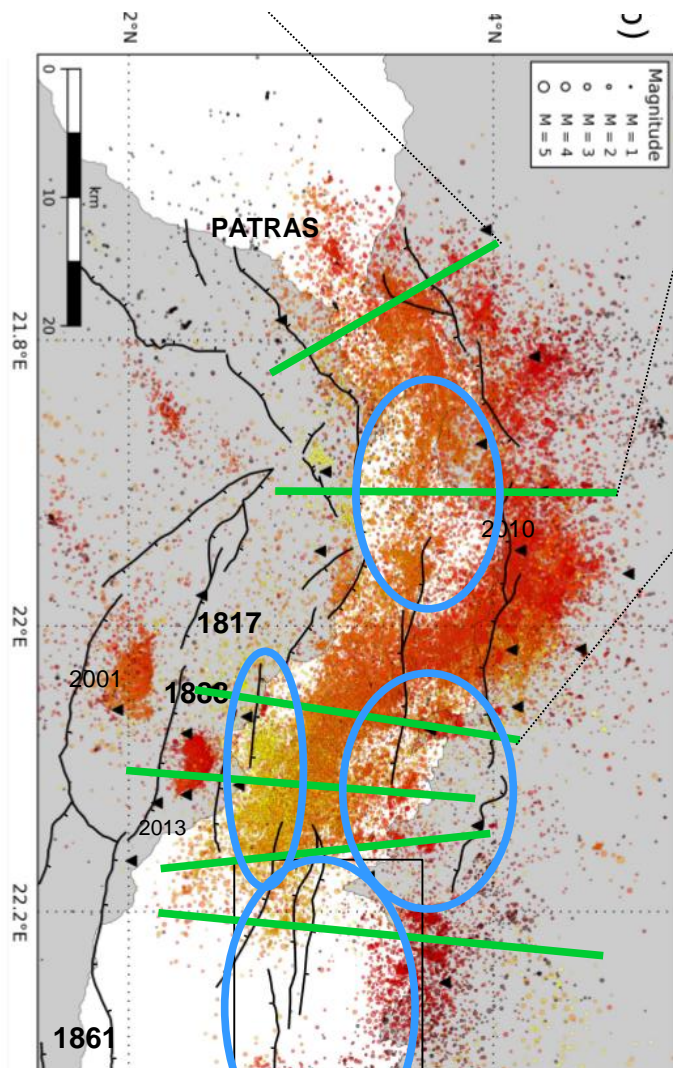
**The « detachment » is probably unmaturing, growing downwards,**

- not yet efficiently connected to the visquous mid-crust**
- not the main guide for the western rift opening...**

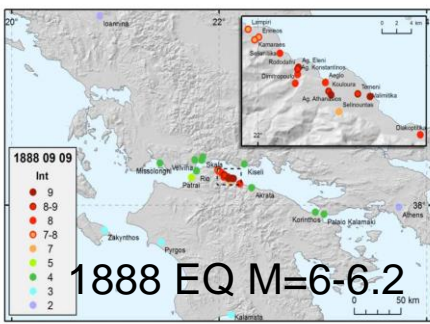


## Seismicity 2000-2015

Duverger et al.  
2018



# Earthquake rupture forecast for $M > 6$

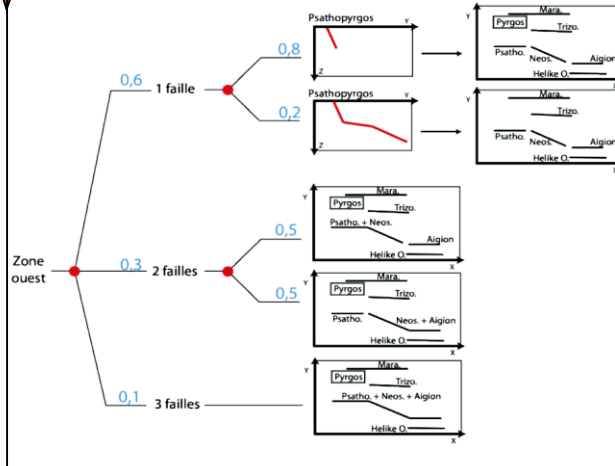


Albini, 2013;

- Fault slip rates
  - Strain from geodesy
  - Paleoseismicity
  - Historical seismicity
  - Instrumental seismicity
- + uncertainties**

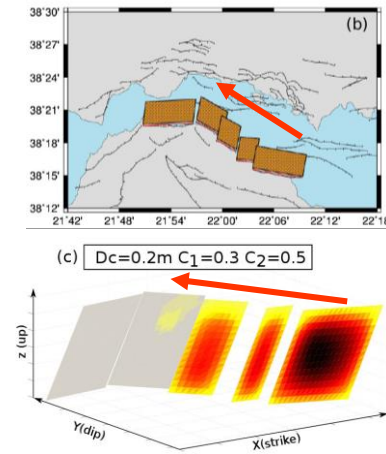
Logic tree:

- Seismicity models
- fault geometries



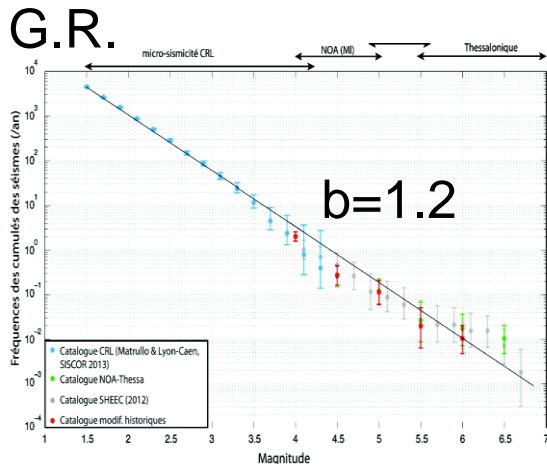
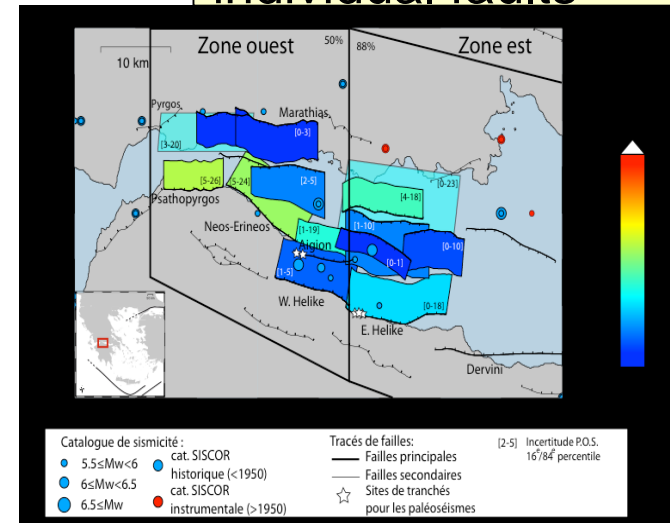
Boiselet et al. , 2014

Dynamic ruptures across fault segments



Durand et al. 2014

ERF forecast for individual faults



**Next 30 years:**

**P=1 % to 20 % of rupture on individual faults**

**P>50% on the fault system**



## études en cours sur CRL

### - **Imagerie de la Structure :**

ondes de surface par corrélation de bruit (UPatras)

temps d'arrivées de la microsismicité (Ecole des Mines, ENS,...)

Fonction récepteur : interfaces crustales (Montpellier, Ecole des Mines)

Anisotropie des ondes S ( NKUA)

### - **Imagerie des Sources :** ENS, IPGP, Géoazur, Isterre, ...

analyse des multiplets long et court terme

dynamique spatio-temporelle de la sismicité

### - **aléa sismique**

lois d'atténuation pour évaluation rapide de dégats : Aigion et Patras (NKUA)

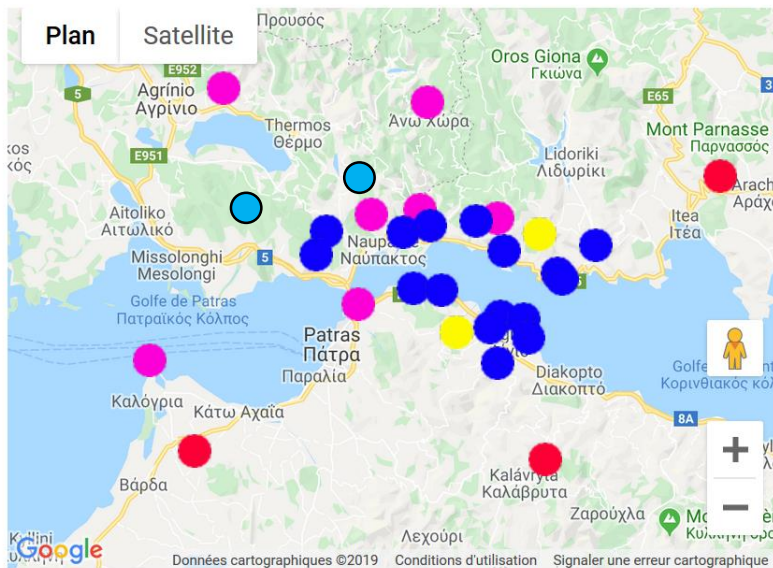
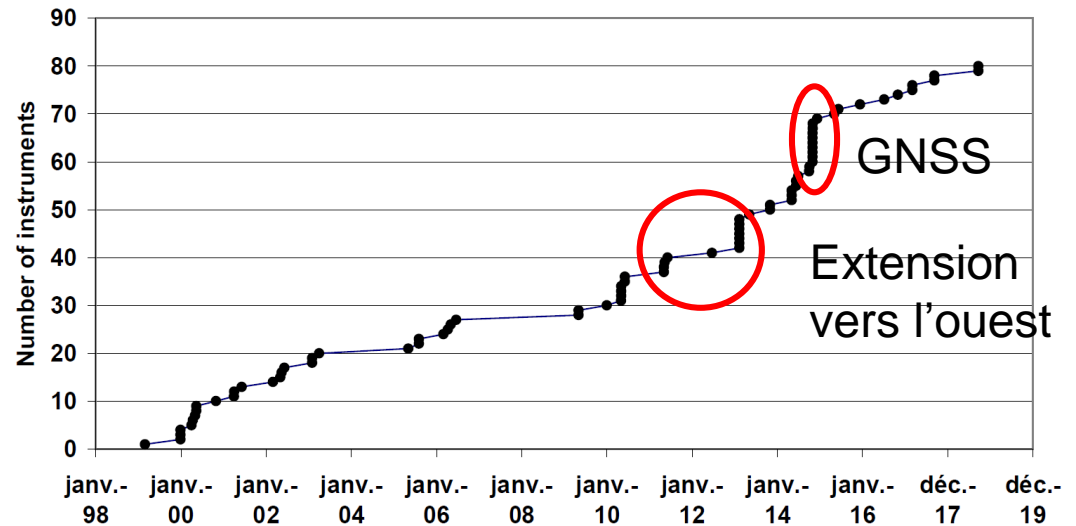
utilisation des flux temps réel pour alerte précoce au séisme (Upatras)

# CRL : Near Fault Observatory (NFO) , EPOS

crlab.eu

70 articles depuis 2005  
>20 thèses

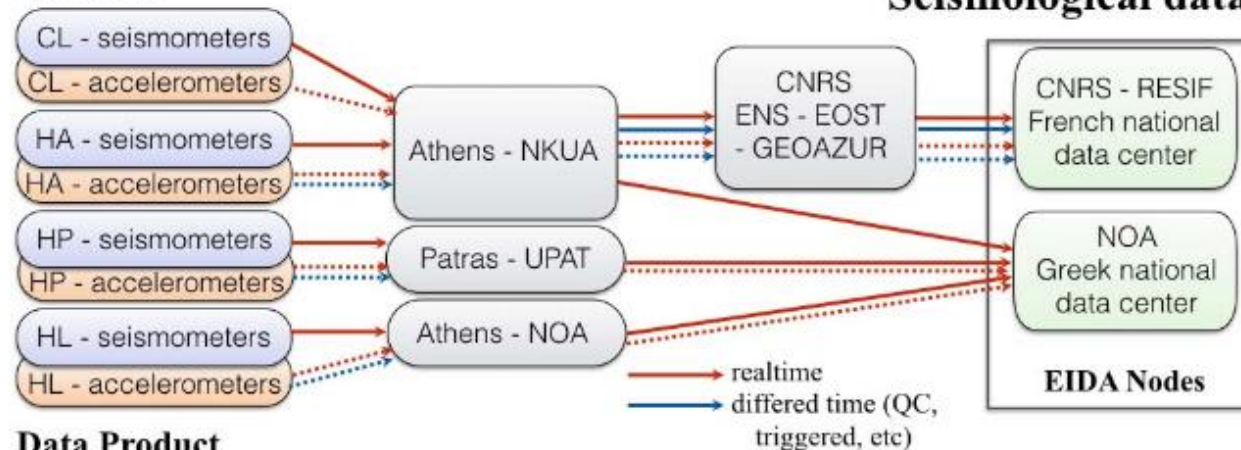
2019: 83 instruments on  
70 différent sites operated by:  
51 CNRS  
10 UPAT  
12 NOA  
7 NKUA  
3 CUP



Velocimétrie et déformation  
en bleu les instruments du CNRS

Toutes les données (velocimétrie et  
GNSS) sont collectées en temps réel

## Raw data



## Seismological data

## Data Product

### Focal mechanism

Athens - NOA

Athens - NKUA

Patras - UPAT

available through  
respective websites, and  
EMSC - webservice

### Earthquake catalog and parameters

CNRS - ENS - EOST

available through  
CRL website  
CNRS - webservice

Athens - NOA

Athens - NKUA

Patras - UPAT

available through  
respective websites  
EMSC - webservice

### Products in development

Multiplet catalogs

CNRS - EOST - IPGP  
- ENS

Seismic Velocity  
and Attenuation Models

CNRS - ENS

Athens - NKUA

# CRL :

## - Near Fault Observatory (NFO) dans EPOS depuis 2017

- données CRL sismo et GNSS disponibles en temps réel par RESIF
- comité de coordination CRL, bureau de direction des 5 NFOs
- financement grec indépendant (HELPOS)

## - “Site instrumenté” de l'INSU depuis 2017

mais pas de soutien récurrent du CNRS : 0 ke en 2017-2018, 5 ke en 2019

→ financement sur projet ANR et EC jusqu'en 2015

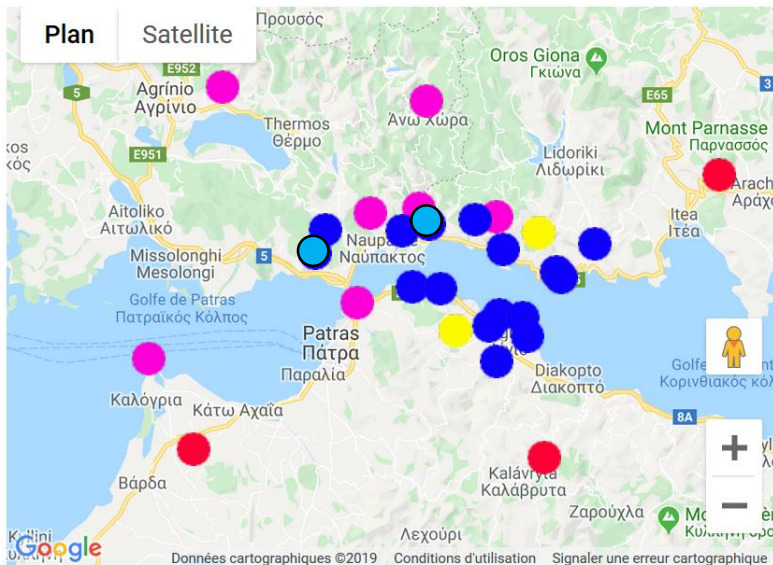
+ soutiens ponctuels des labos

→ **Infrastructure de terrain CRL en situation critique pour maintenance**

→ **besoin de 75 ke/an pour les 50 instruments du CNRS**

- maintenance terrain
- base de données
- management

**(112 ke/an pour CRL)**



# CRL : observatoire international depuis 2001

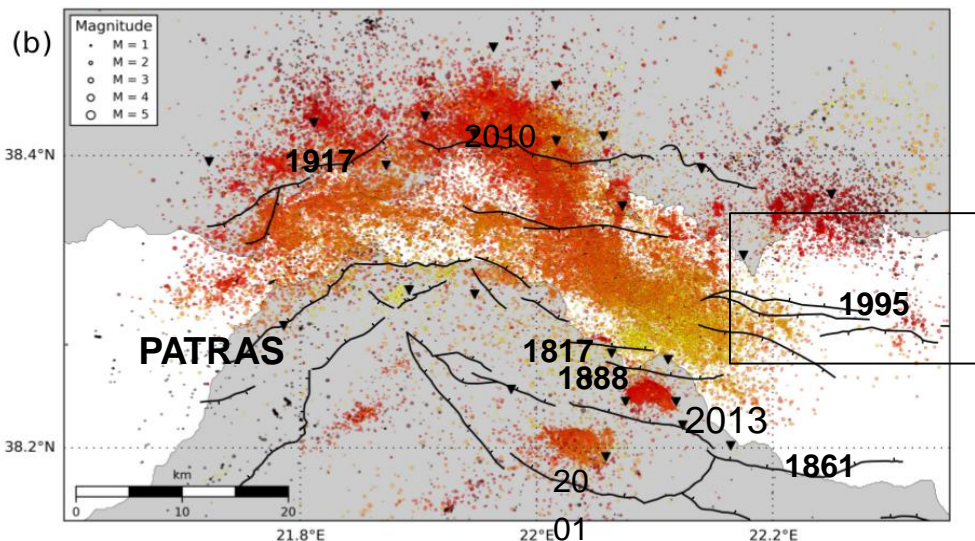
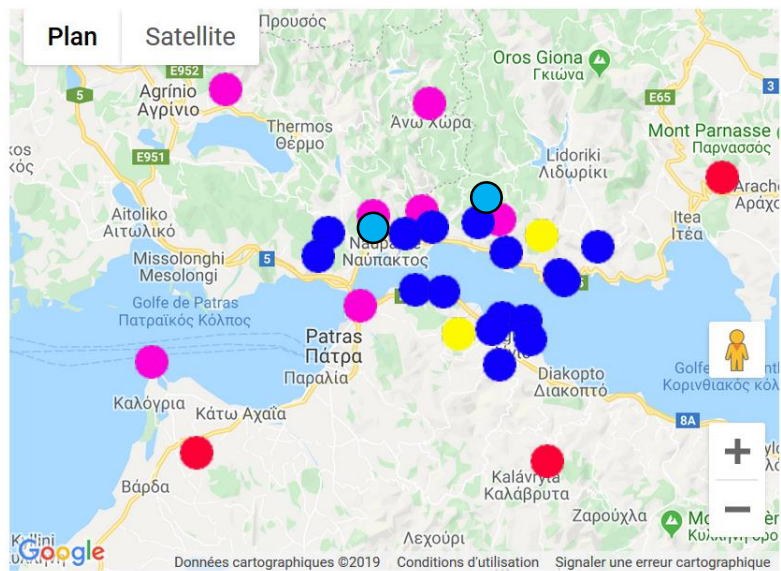
- Near Fault Observatory (NFO) , EPOS, depuis 2017
- Site instrumenté de l'INSU depuis 2017

**Observatoire = infrastructure de terrain + base de données**

**Quel statut/financement pour maintenir l'infrastructure de terrain?**

France:

ENS, EOST, GEOAZUR, IPGP, GM, ISTerre, IRSN, Ecole Mines, ...  
+ partenariat international: Grèce, Italie, Rep. Tchèque



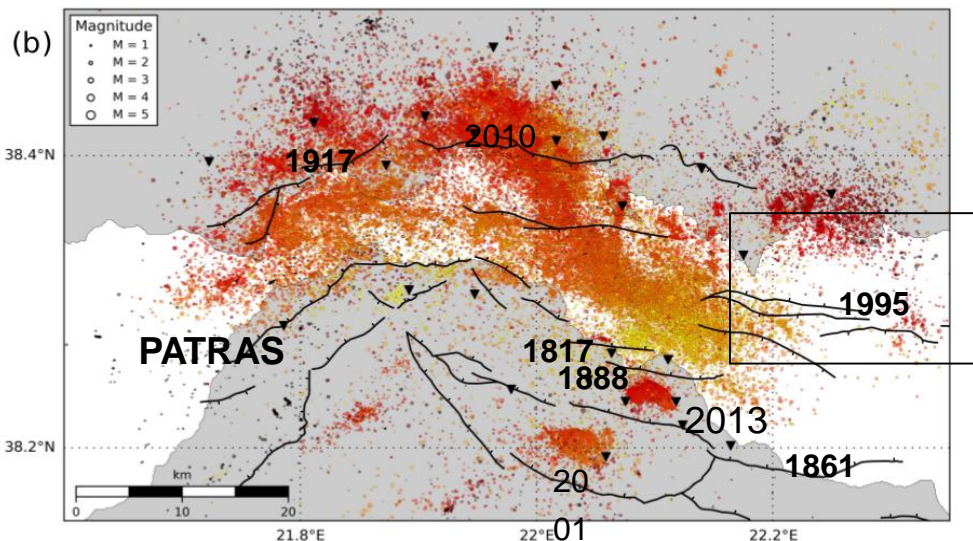
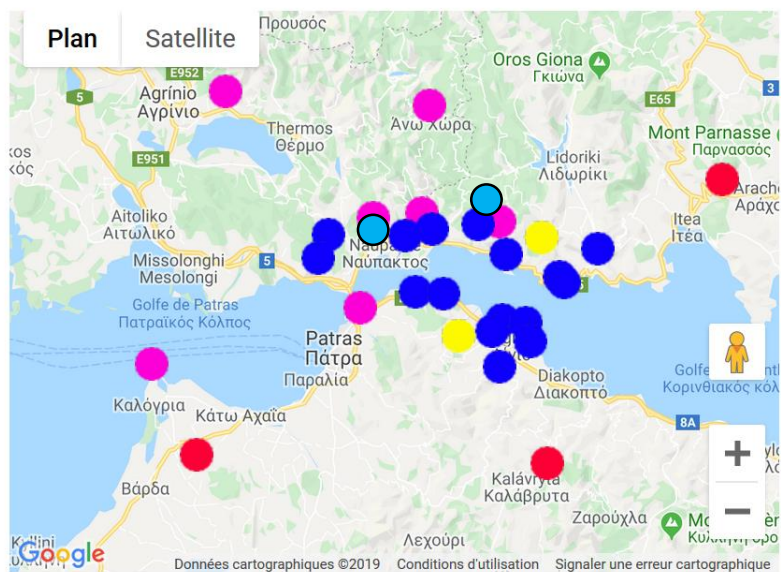
# Observatoire international de Corinthe - CRL

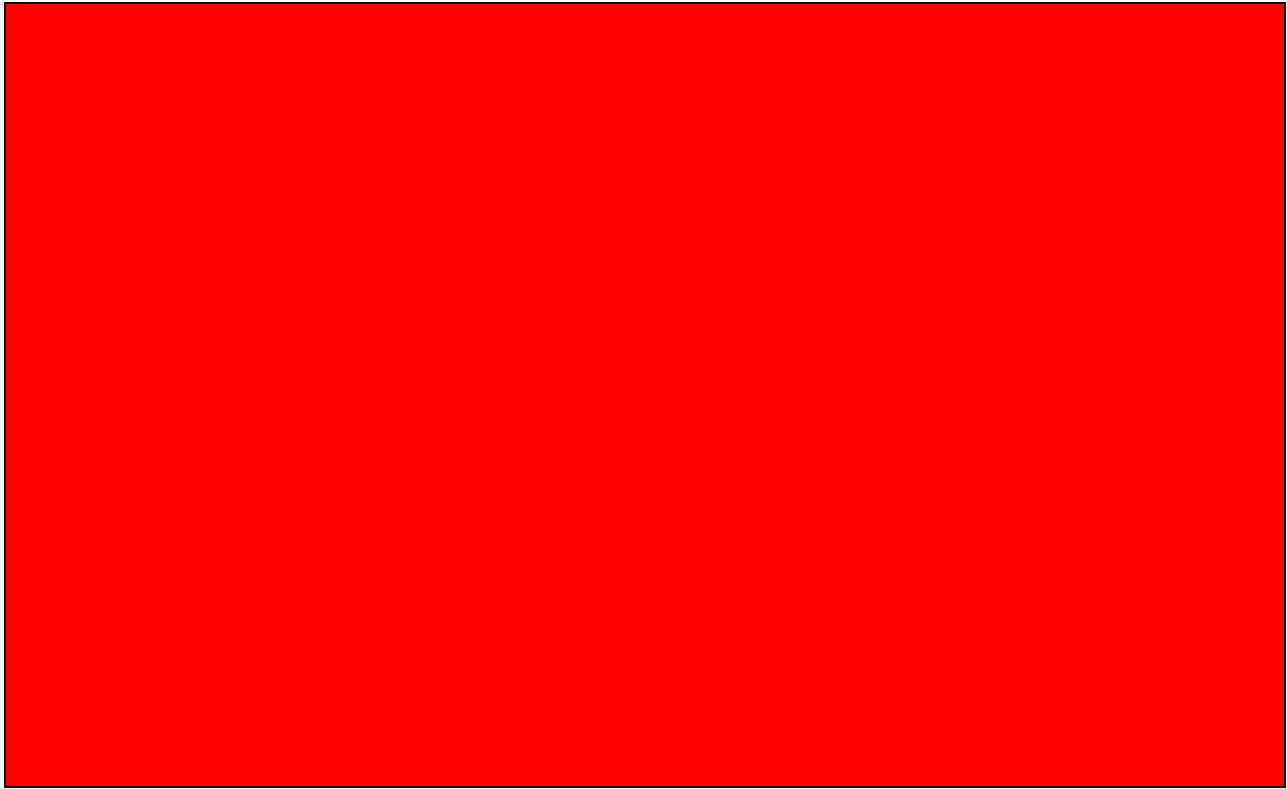
## Thématiques de recherche sur le long terme

- Mécanique multi-échelle d'un système de failles
- Dynamique spatio-temporelle de la sismicité: essaims, migration, repeaters
- Couplages sismique/asismiques
- Rôle des fluides
- Lien entre structure crustale, sismicité, et déformation/glislements asismiques
- Mécanique du rifting
- Aléa et risque sismique

...

## Continuité d'observation -> Maintien de l'infrastructure de terrain





# Observatoire de Corinthe – NFO

## Mécanique sismogène des failles:

Multidisciplinaire: Sismologie, géodésie, tectonique

## Structure + Dynamique

## Couplage sismique-asismique - rôle des fluides

- sismicité espace-temps
- multiplets - repeaters
- variation temporelle du milieu (corrélation de bruit, repeaters)
- tomographie 3D (réseaux actuels ou temporaires)
- recherche de transitoires asismiques (GPS, extensomètres)
- déformation GPS et INSAR : modèles élastiques
- lien sismicité / taux déformation
- modélisations hydromécaniques
- modélisation mécanique du rifting (cycles sismiques) – 2D- 3D
- cinématique/mécanique des différentes zones sismogènes
- aléa sismique dépendant du temps
- mouvements forts (analyses et prédictions)
- risques: Aigion, Patras

.....

## Instrumentation nouvelle

- Faille Rio-Patras / Psathopyrgos
- Réseaux temporaires tomo sismique
- instrumentation sous-marine
- MT – tomographie résistivité

Pascal Bernard

Pierre Briole

Lucile Bruhat

Louis de Barros

Fabian Bonilla

Anne Deschamps

Pierre Dublanchet

Blandine Gardonio

Stéphanie Gautier

Alexandrine Gesret

Maxime Godano

Agnès Helmstetter

Emilie Klein

Hélène Lyon-Caen

David Marsan

Jean-Arthur Olive

Alexis Rigo

Clara Duverger

Cécile Doubre

Françoise Courboux

Claudio Satriano

Christel Tiberi

Sophie Lambotte

Olivier Lengliné

Sébastien Hock

Oona Scotti

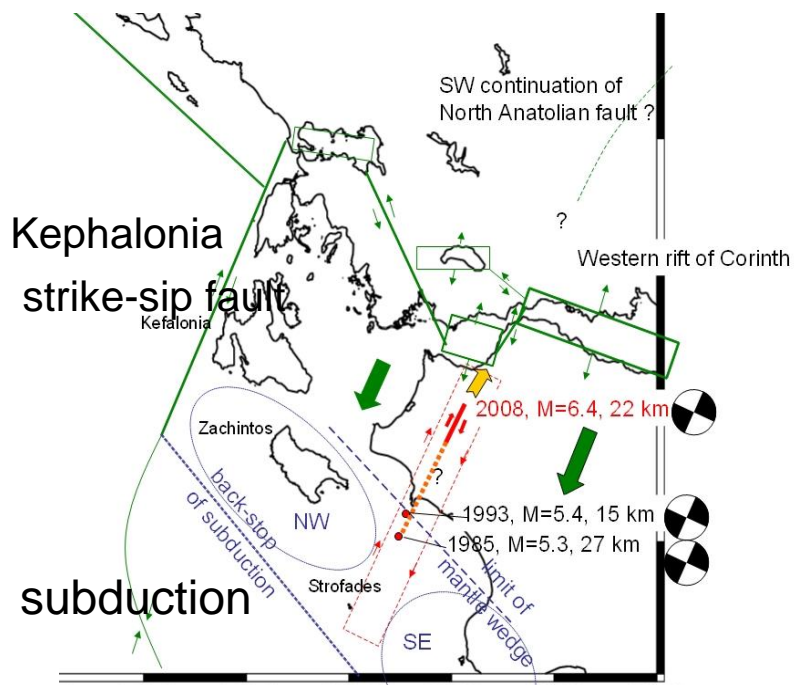
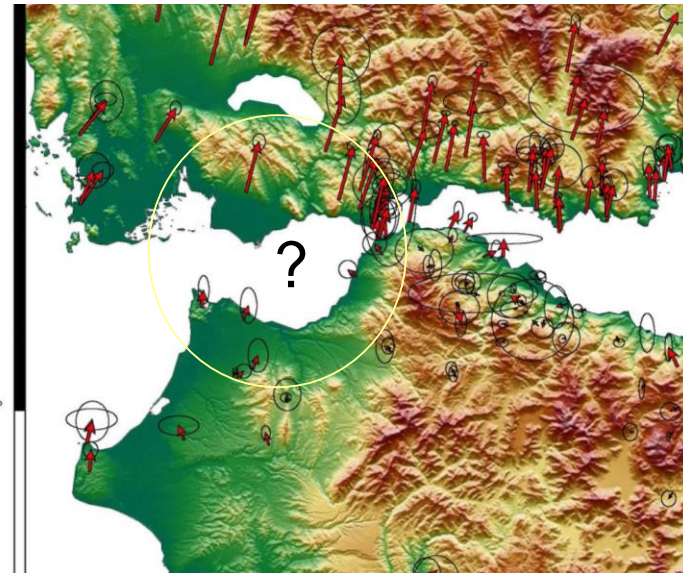
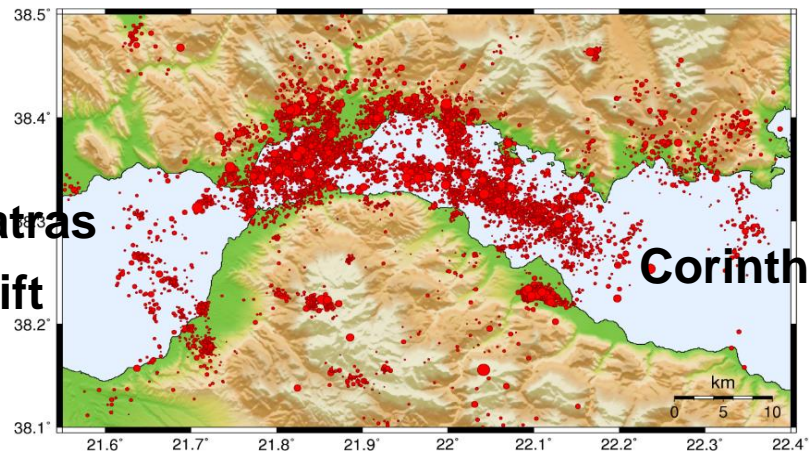
Marc Noble

.....

+NKUA,NOA,UPatras,CUP,...



# How does the rift change towards west?



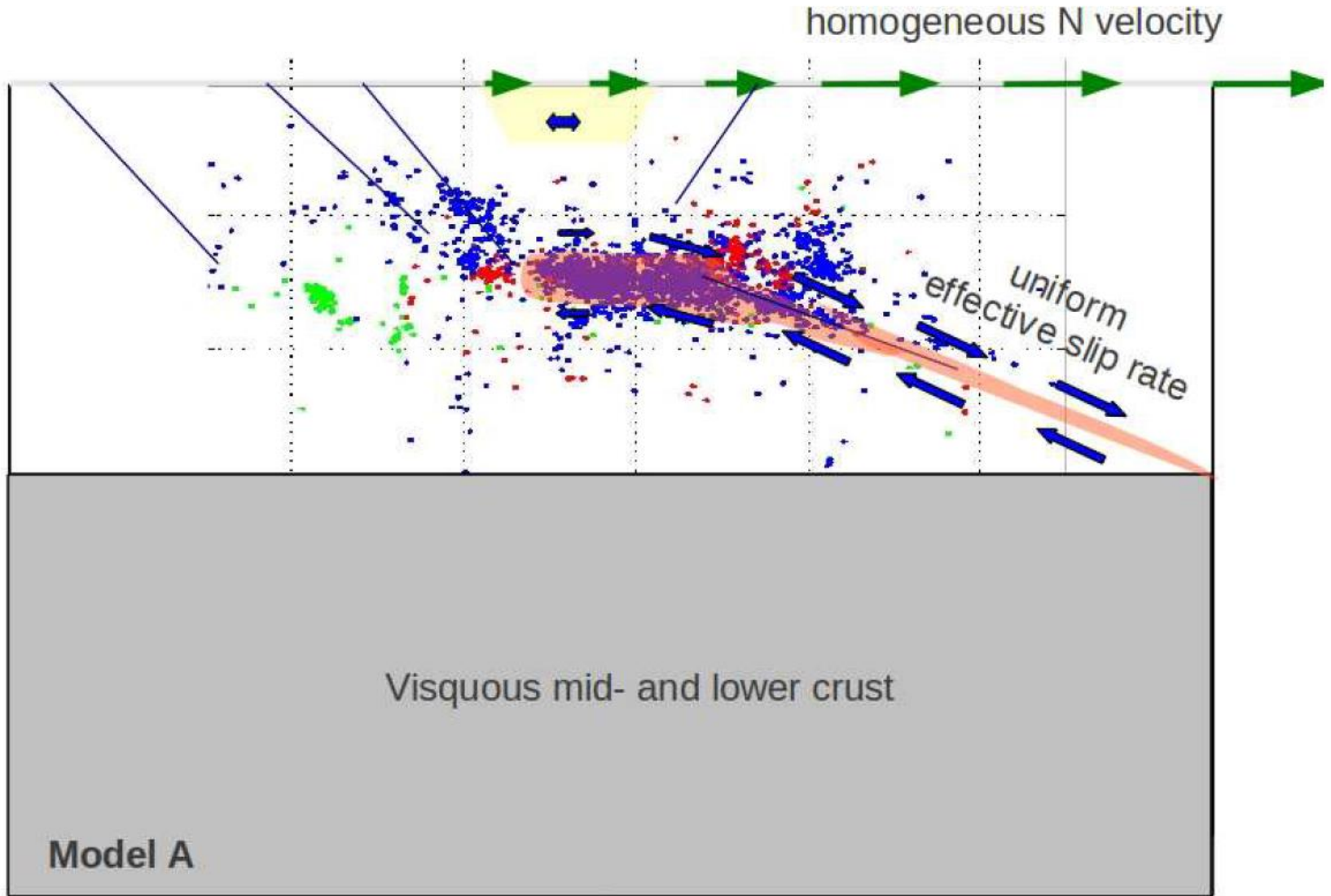
## Western Corinth **Near Fault Observatory (NFO-EPOS)** :

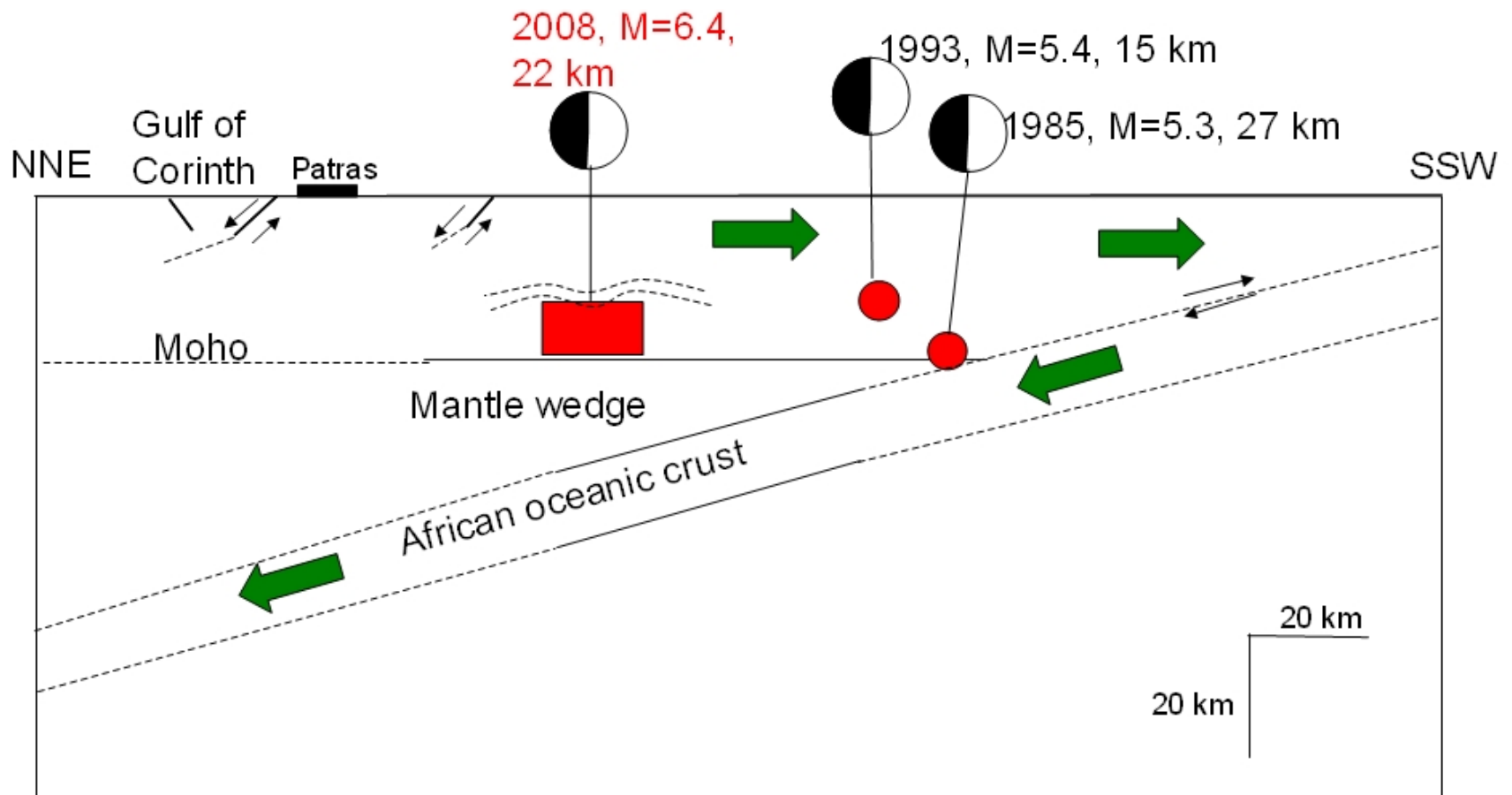
**Multidisciplinary research infrastructure**

**Analysis and modeling of multiscale , coupled seismic/aseismic processes on fault systems**

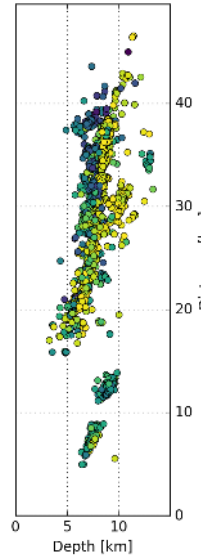
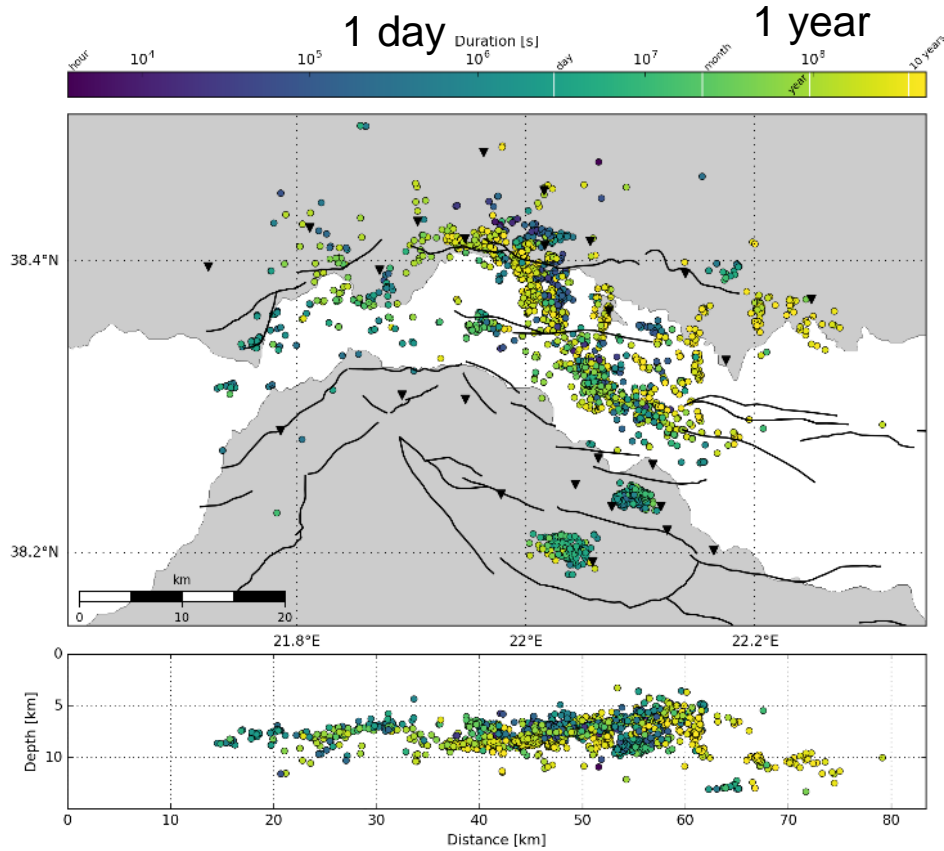
- Enlarge to Patras Rift , to the West
- Connection to subduction and Kephalonia fault system?

Detachment model with constant opening rate (remote) and decreasing friction





# Long-lasting multiplets



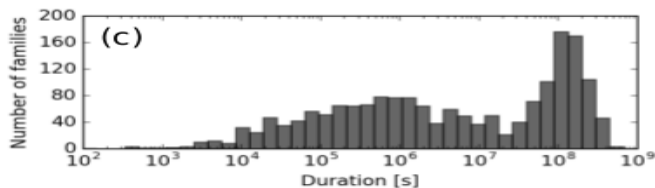
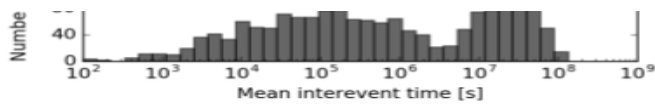
**Short duration:**  
Swarms,  
pore pressure diffusion

**Long Duration:**  
Creep,  
Edge of major asperities

(Godano et al., 2015)

## Long duration multiplets:

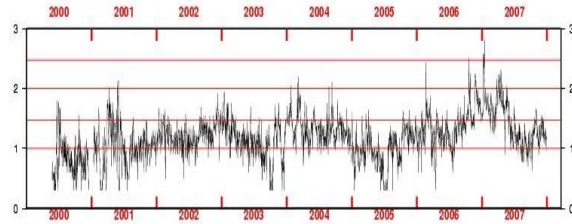
Multiplets with  $N > 10$  events



- 1995 fault zone
- deep Agion fault zone
- deep Lambiri/Psathopyrgos fault zone

# Stability of rift opening rate versus variability of micro-seismicity: different strain sources ?

Log(seismicity rate) /day



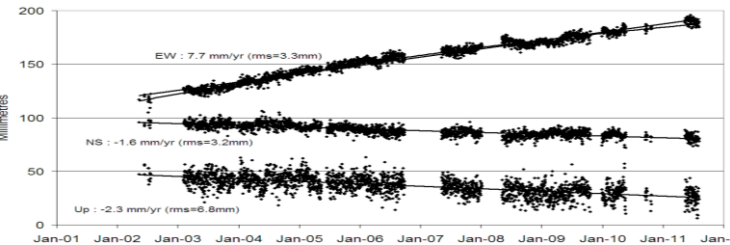
Seismicity:  
Fluctuating  
stressing rate

contradiction

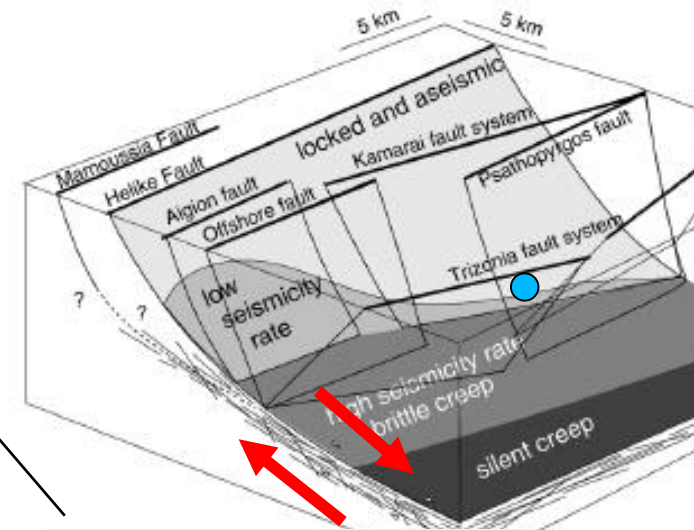


Displacement (N) (mm)

CRL - Permanent station TRIZ (Trizonia) - 2257 samples - Corrected



GPS:  
constant  
stressing rate



Shearing of the  
« detachment »

cGPS (TRI)

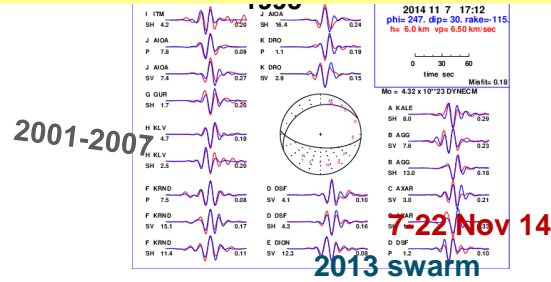
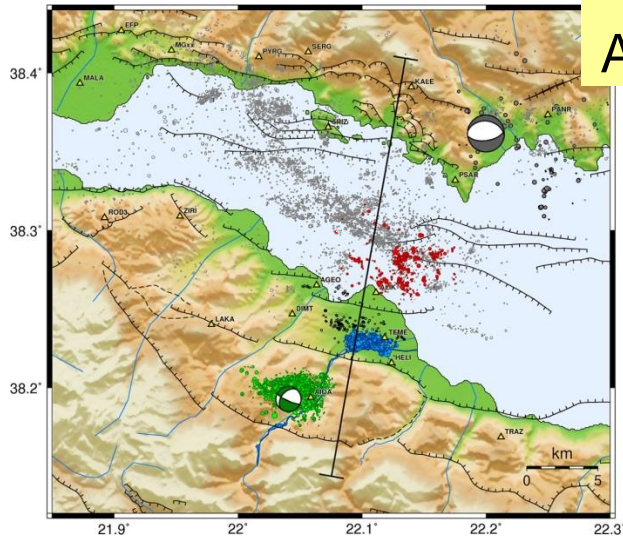
Briole et al. 2013

The fluctuating micro-seismicity and the stationary GPS rate are unlikely to have the same source:

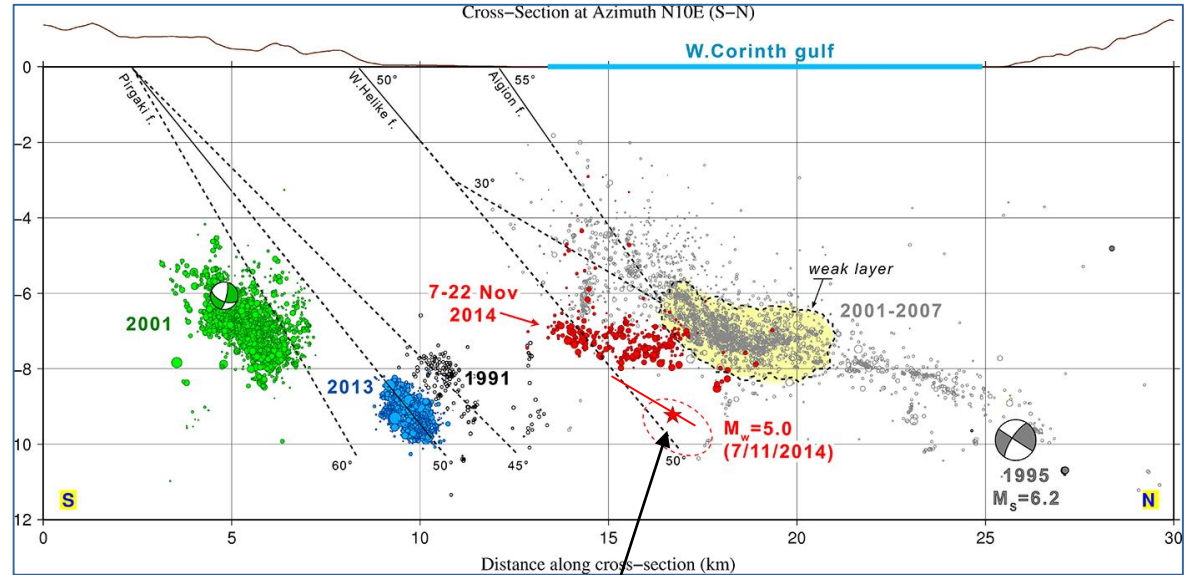
The main source of GPS rate may not be the shearing of the seismic layer

# The 7 nov 2014 earthquake, M=5

## At the root of the Helike fault?



2001

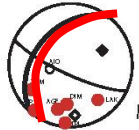
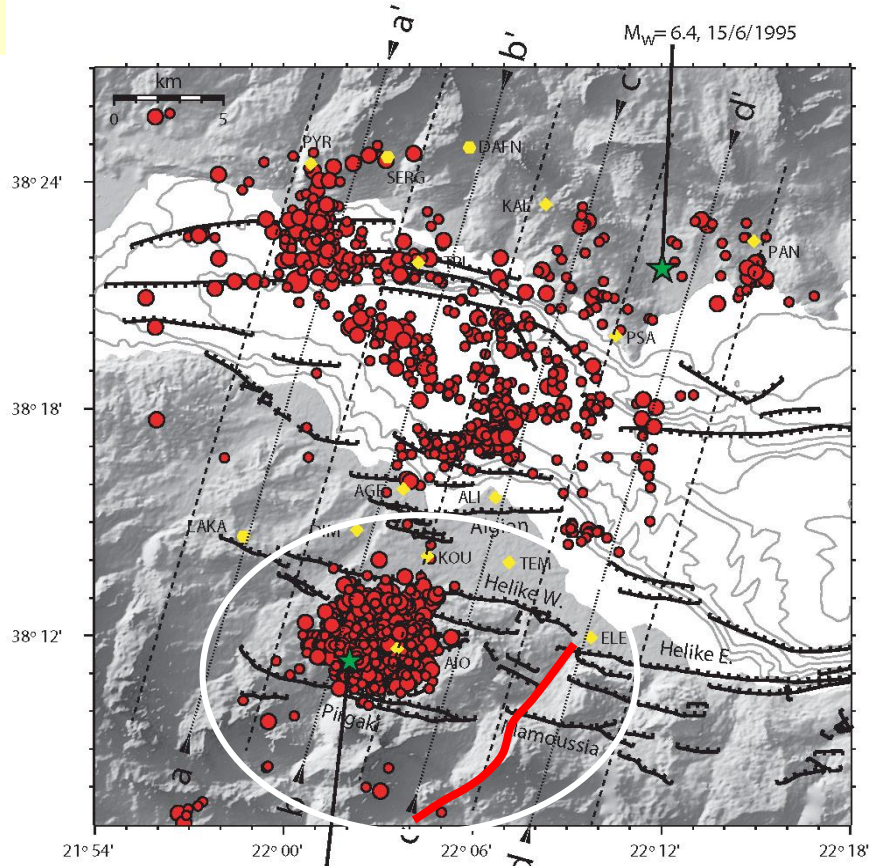


The crust is brittle and seismically active **beneath** the main seismic layer

# 2001 seismic swarm

Re-activation of an ancient fault, dip 45° NW

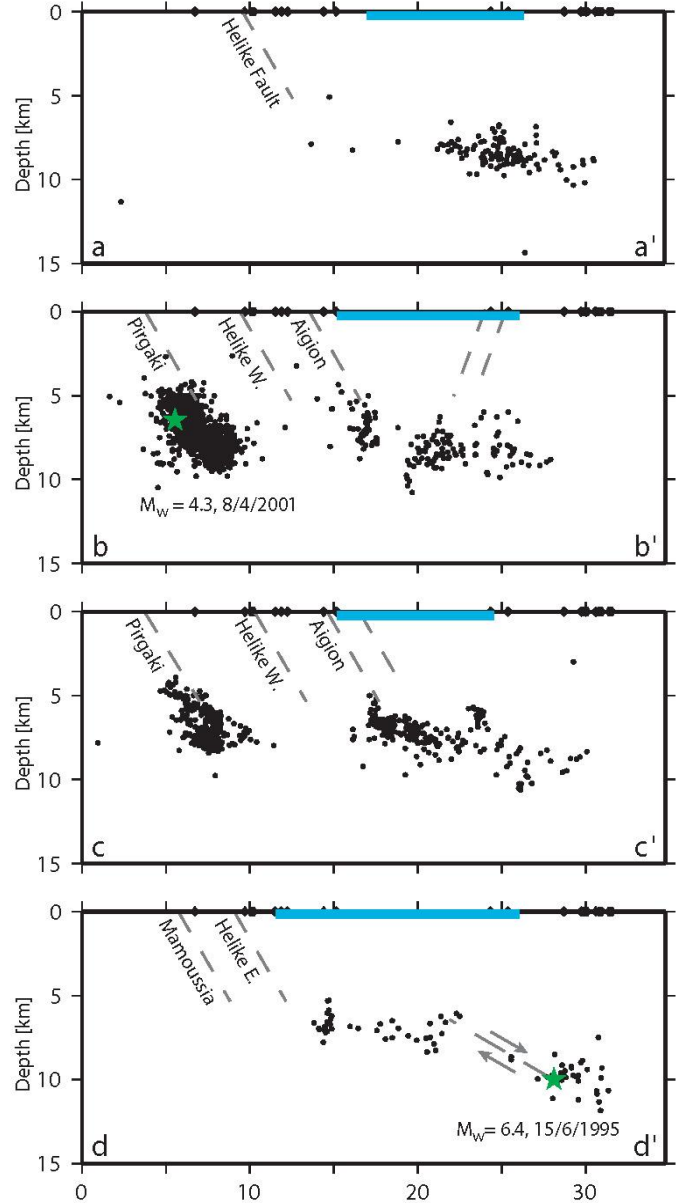
3 months -  $M_{max}=4.2$



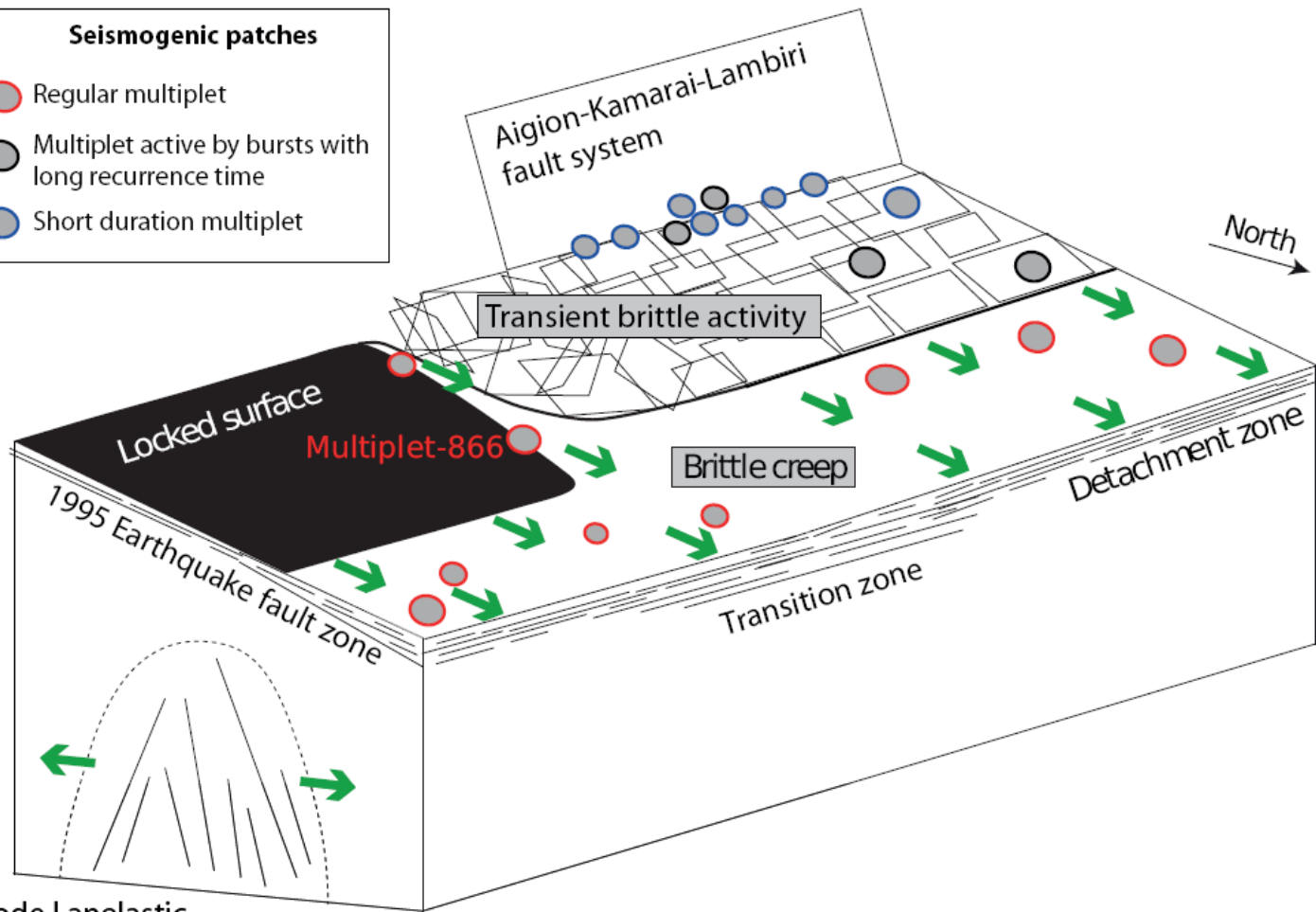
**Kerinitis fault**

$M_w = 4.3, 8/4/2001$

$M_w$ : ● 1-2   ● 2-3   ● 3-4   ● 4-5



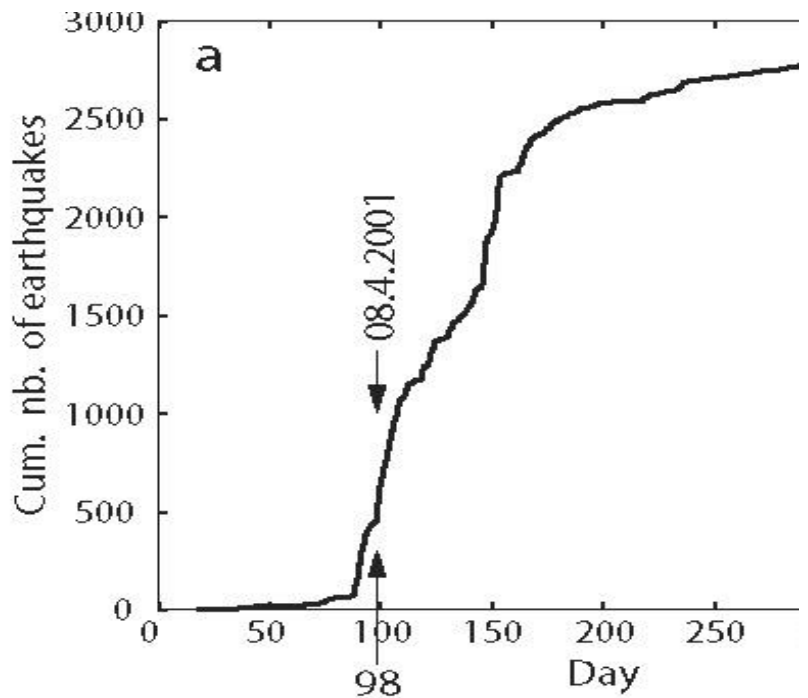
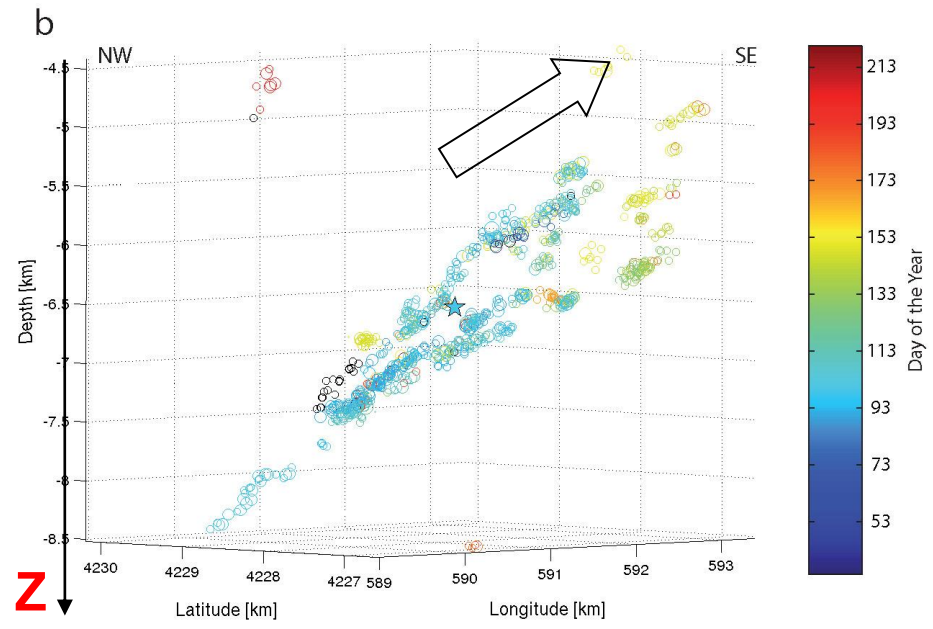
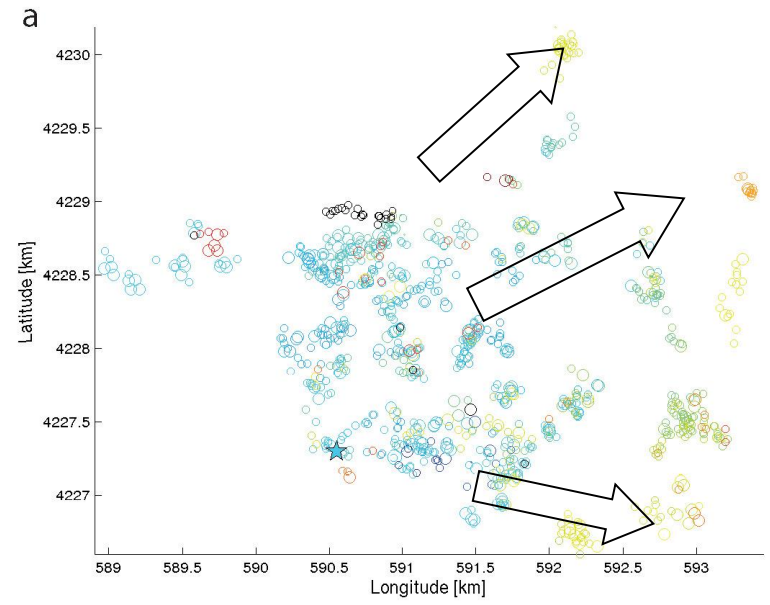
- Seismogenic patches**
- Regular multiplet
  - Multiplet active by bursts with long recurrence time
  - Short duration multiplet



Mode I anelastic  
horizontal opening



# 2001 seismic swarm: relocated multiplets



3 – 4 months:

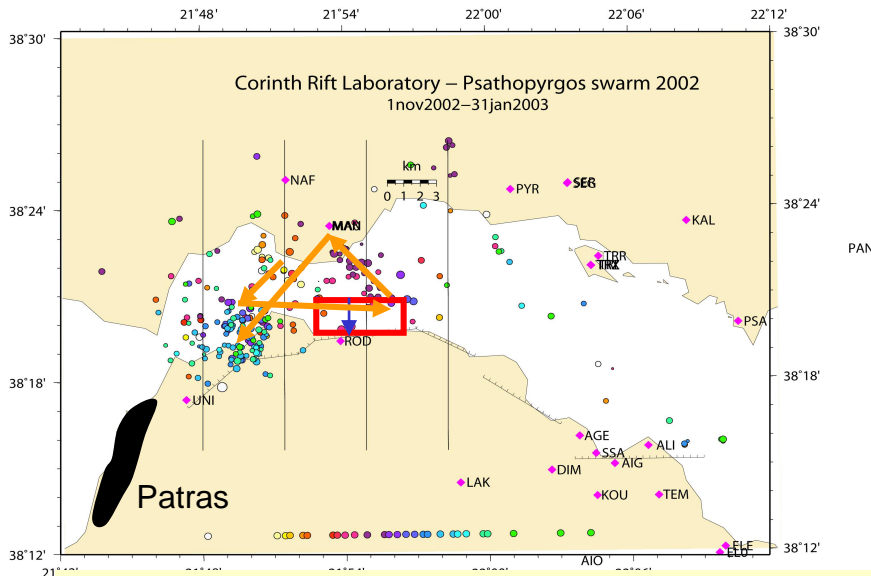
Upward migration on the Kerinitis fault, 20-30 m/day

pore pressure pulse?

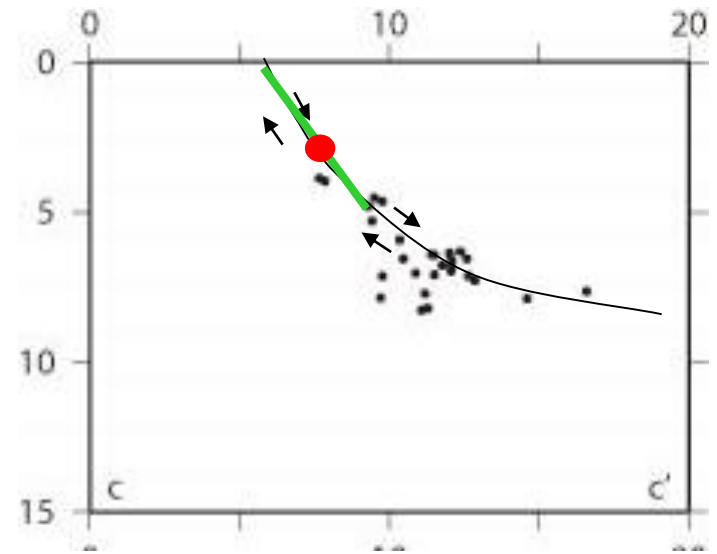
- Diffusivity  $0.1 \text{ m}^2/\text{s}$
- Hydraulic conductivity  $10^{-5} \text{ m/s}$
- Permeability  $7 \cdot 10^{-13} \text{ m}^2$

+ coupled slow slip ?

# 2002 : seismic swarm and transient creep on the Psathopyrgos fault



Migration velocity 100 m/day to 1 km/day



Aseismic slip, 30 minutes on

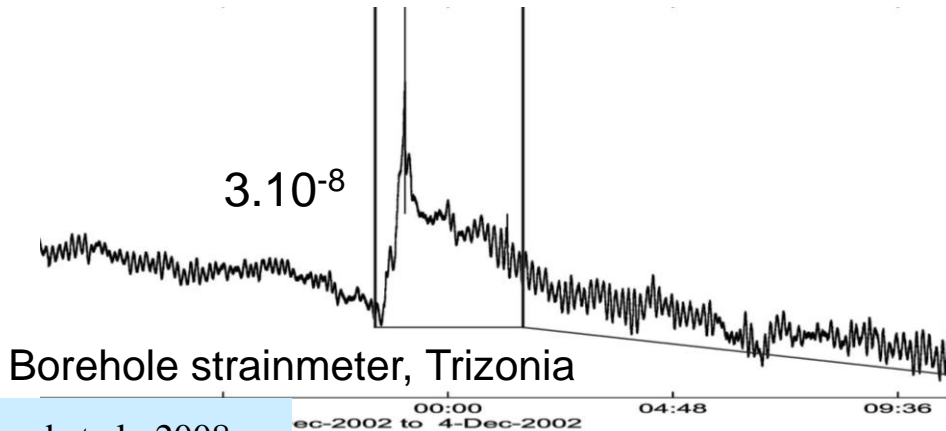
Psathopyrgos fault :

**SLIP = 10 cm - Mw equiv = 5**

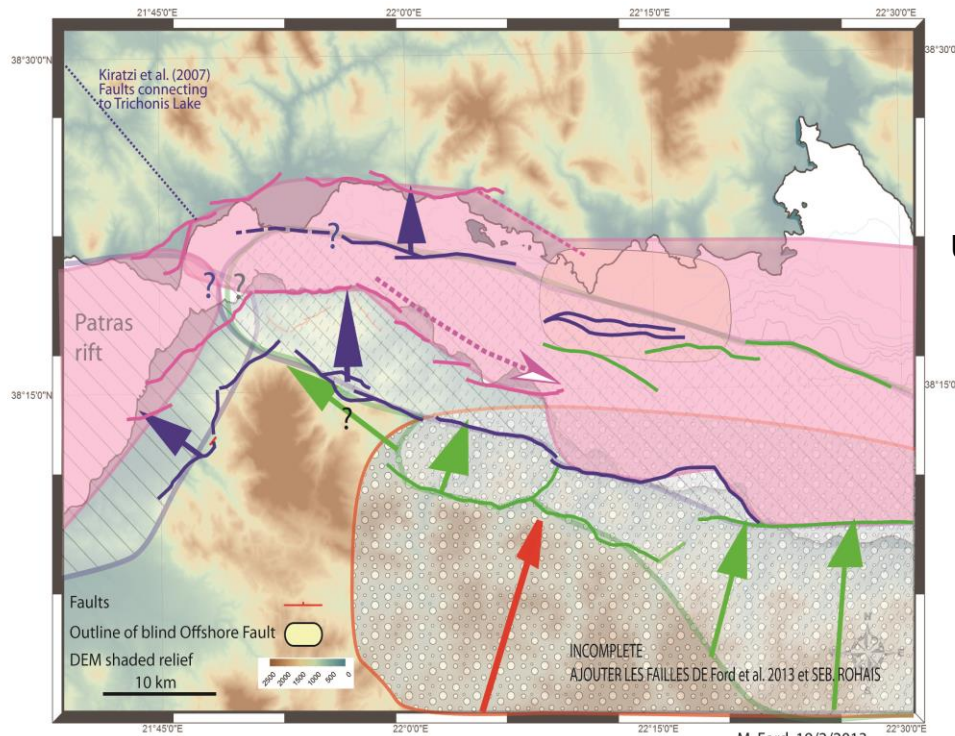
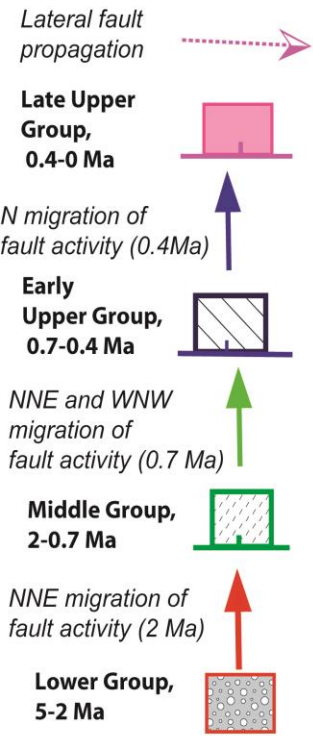
Strain peak:

M=3.5 earthquake,

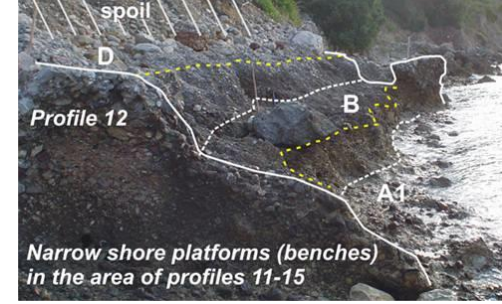
very shallow depth 2.5 km



# Geological studies

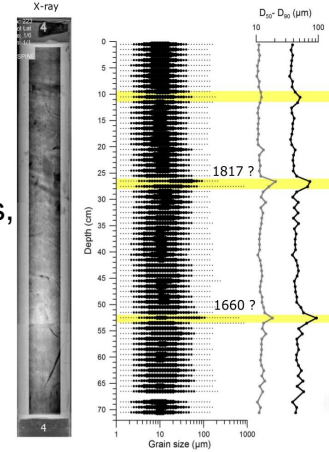


Ford, 2013



Uplifted paleo-shorelines, Palyvos et al., 2008

Cores of offshore sediments, Beckers et al., 2013



Evolution and present activity of fault system: migration of depocenter, uplifted fan deltas, dating of fault activity and marine terraces, offshore sedimentology (cores, seismics), ...

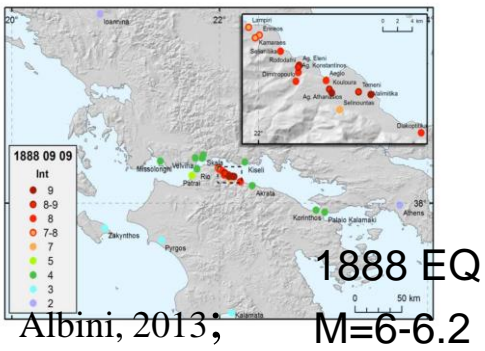
complex inherited 3D faulted structure, results in a complex mechanical and seismic interplay of old and new faults



Helike fault and fan deltas

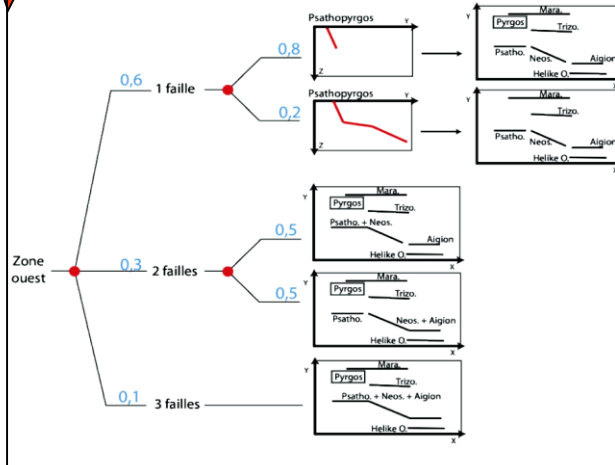
# Earthquake rupture forecast for M>6

Boiselet et al. , 2014

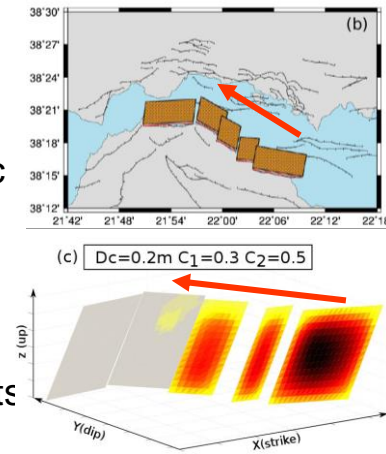


- Fault slip rates
- Strain from geodesy
- Paleoseismicity
- Historical seismicity
- Instrumental seismicity
- + uncertainties

Logic tree:  
- Seismicity models  
- fault geometries

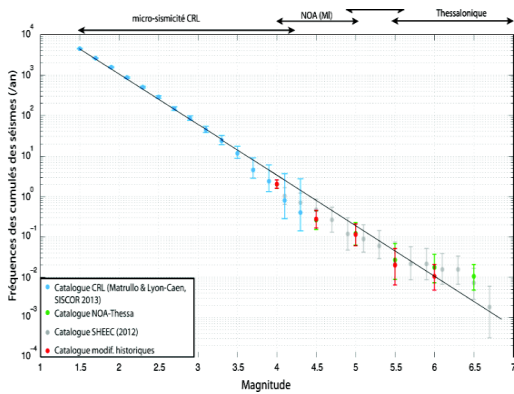


Dynamic ruptures across fault segments

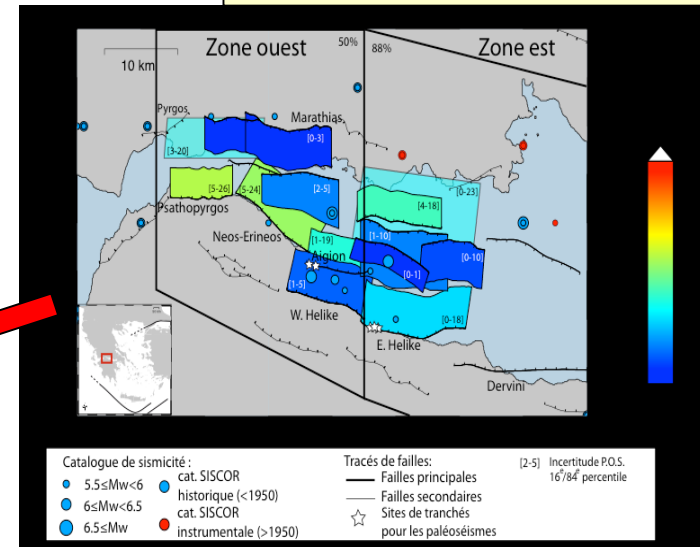


Durand et al. 2014

ERF forecast for individual faults



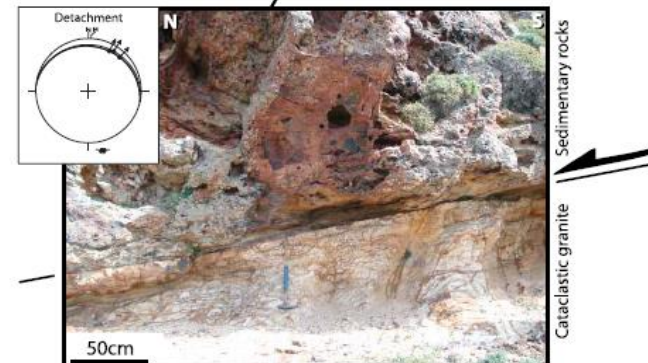
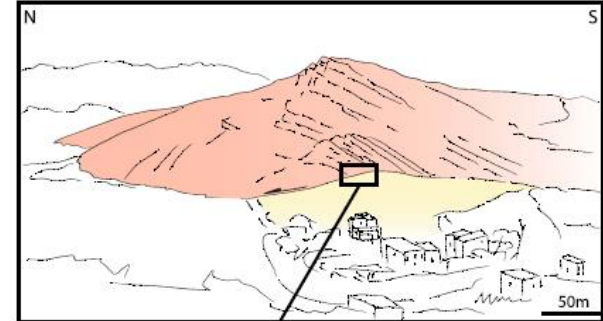
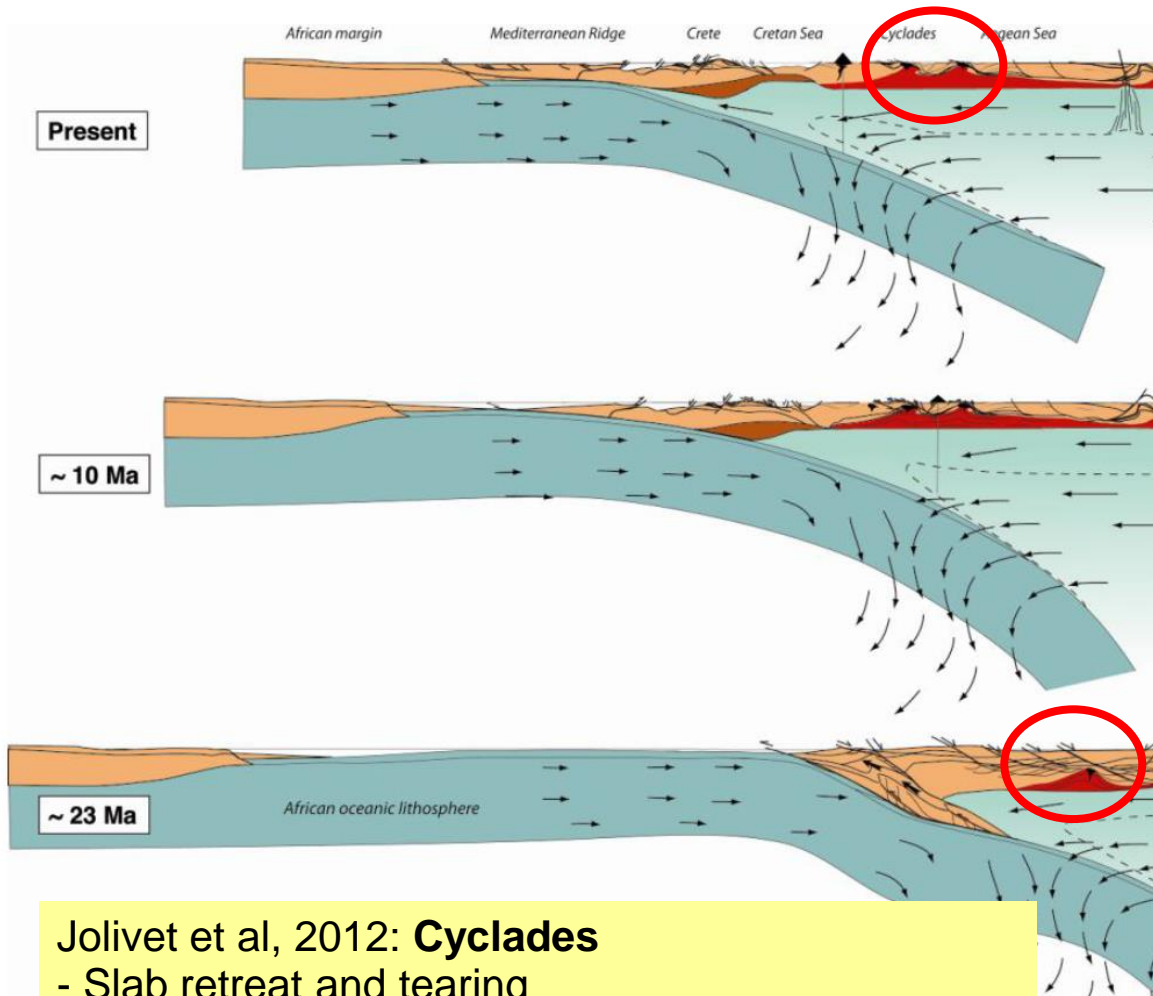
**P = 45 - 75 %  
pour M>6  
0 - 30 ans**



# Aegean tectonics – large scale extension

CS012

LECOMTE ET AL.: KINEMATICS OF MYKONOS DETACHMENT



## Jolivet et al, 2012: **Cyclades**

- Slab retreat and tearing
- Crustal scale detachment
- Exhumation of core complex, from 8-15 km depth

## **Western Corinth rift:**

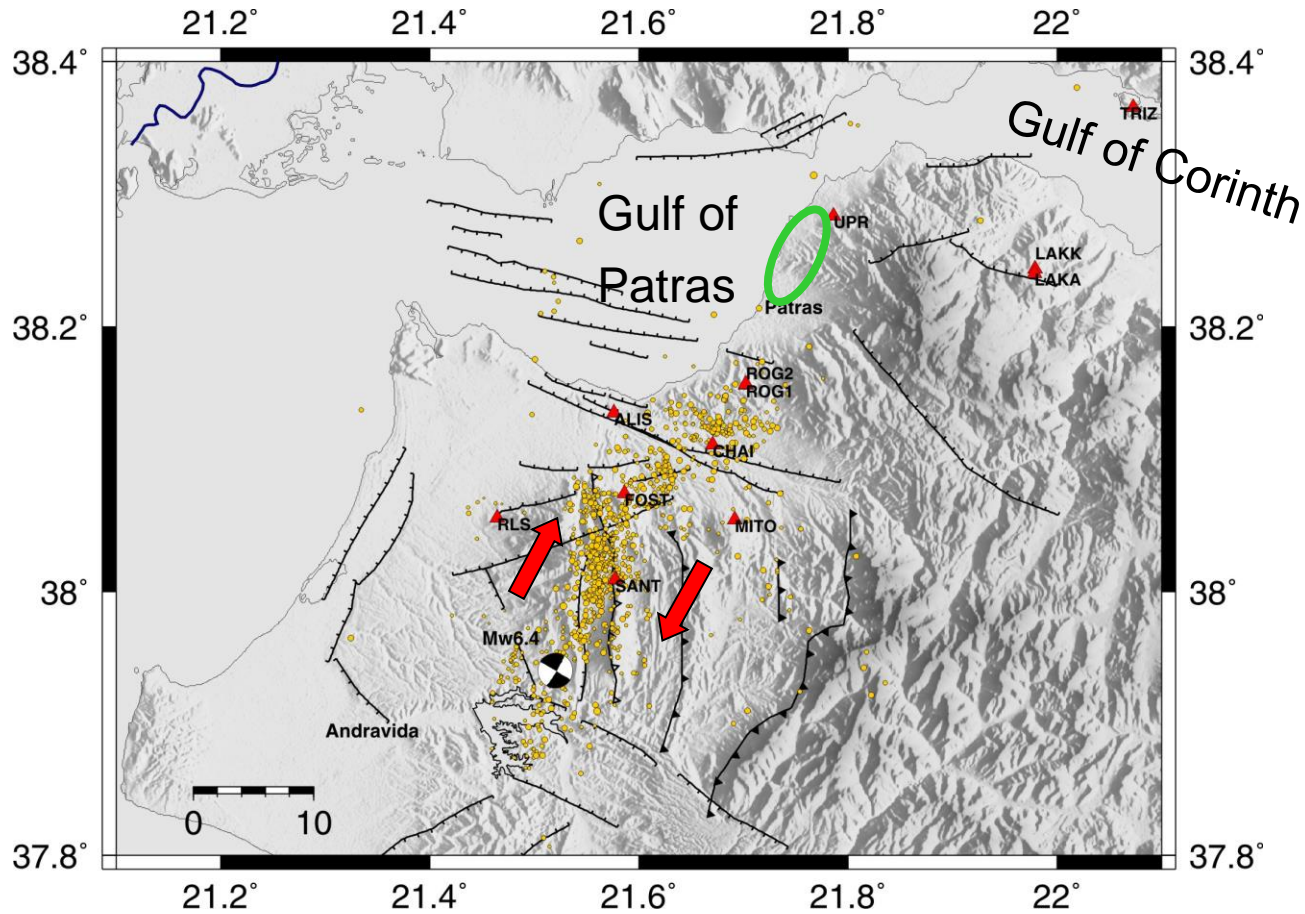
similar to early stage of Cyclade detachment?

LeComte et al. 2010  
Cyclades:  
Mykonos detachment

# The western end of the Corinth rift:

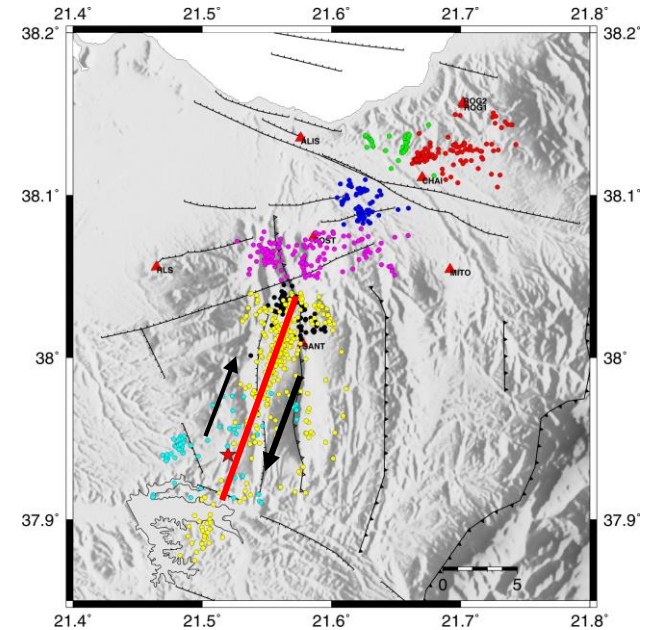
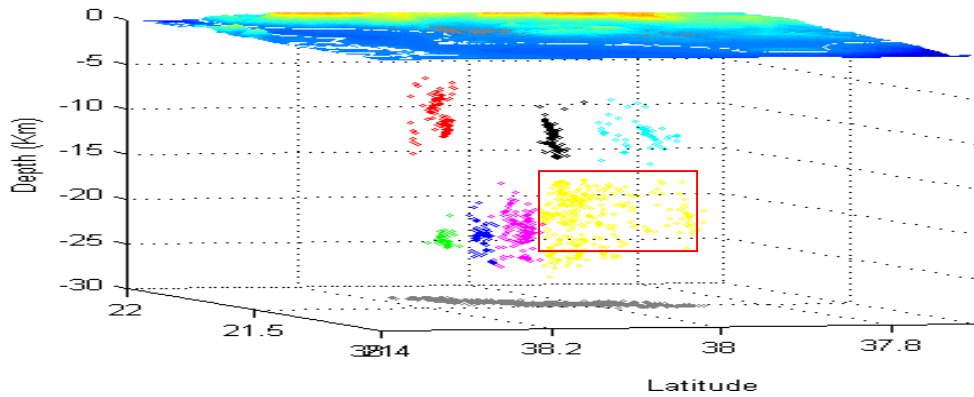
the SW connection to the subduction

Insights from the M=6.2, Achaia-Ilia, 2008 earthquake



# The M=6.2, Achaia-Ilia (Andravida), 2008 earthquake

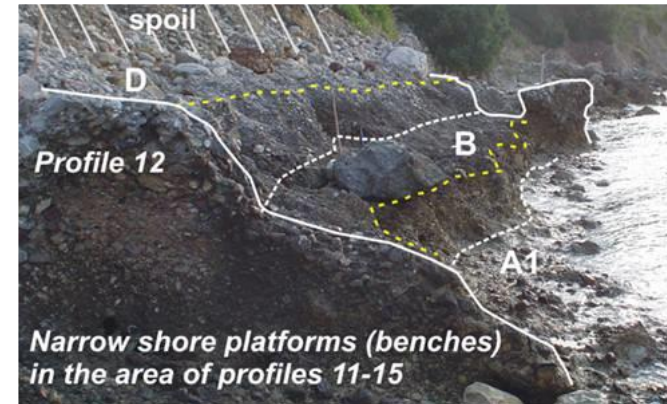
A deep strike slip event on  
a unmature fault



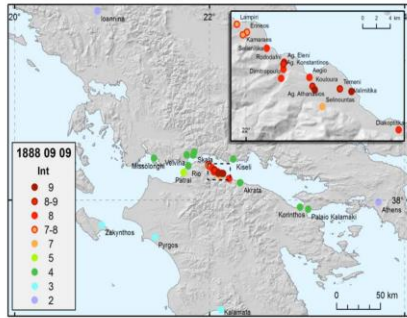
Serpetsidaki et al. 2013

# Past large earthquakes for fault activity and seismic hazard assesment:

Work in progress

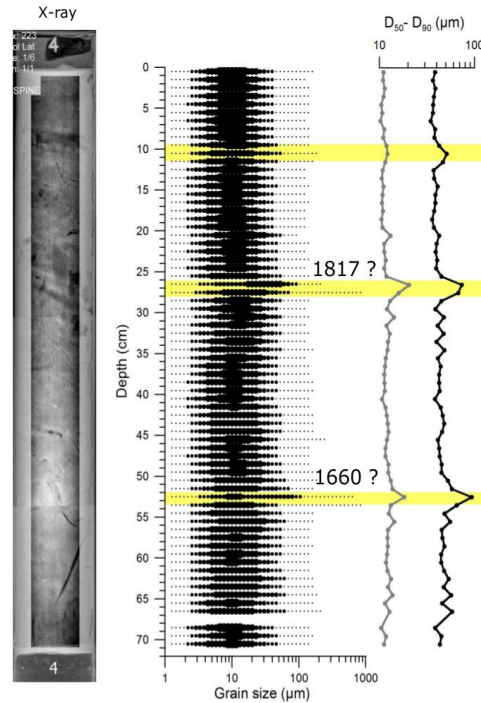


Uplifted paleo-shorelines:  
5 event, 0.5 m each  
in 2000 years  
Palyvos et al. 2008



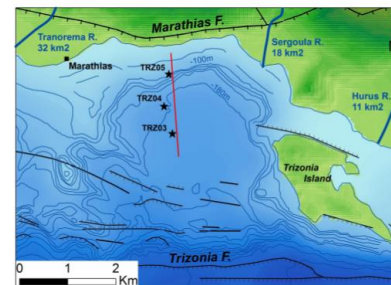
Intensity maps  
of historical  
earthquakes

Albini, 2013



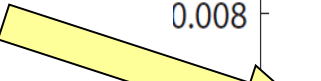
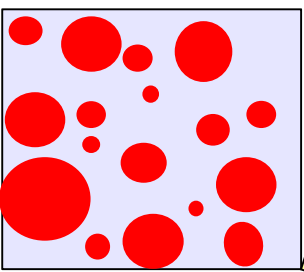
Short gravity core – grain size variations

Beckers et al., 2013





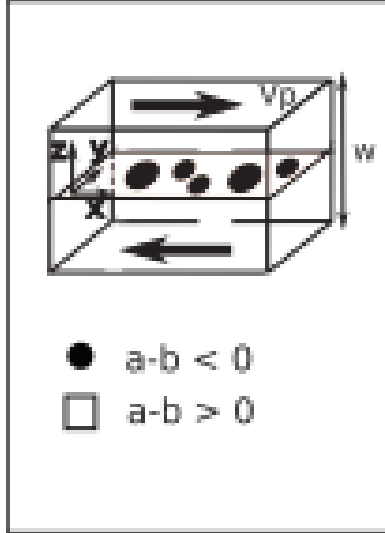
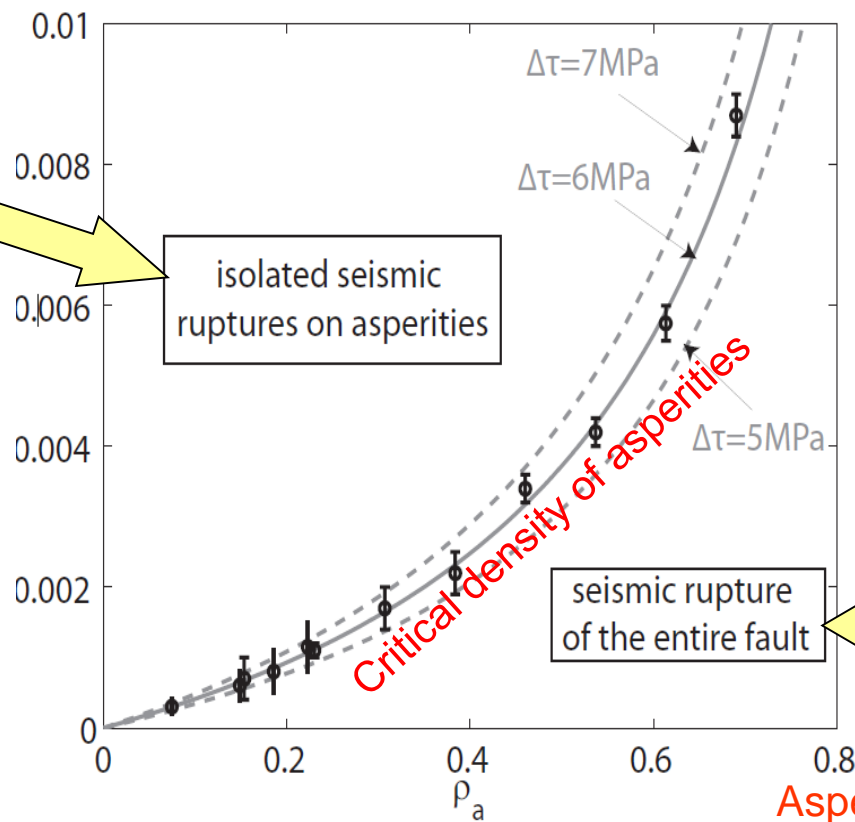
# Models of Interacting Mechanical asperities with R&S friction: threshold for system stability



Delayed interaction

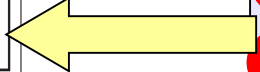
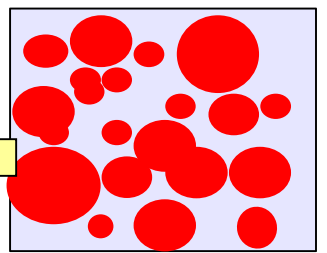
Effective Friction parameters

(a-b) (velocity strenghtening)



Dublanchet, 2013

Dynamic rupture



Asperity density (velocity weakening)

Dublanchet et al., 2013

Stability threshold depends on:  
 asperity density and friction characteristics of inter-asperity « barriers »

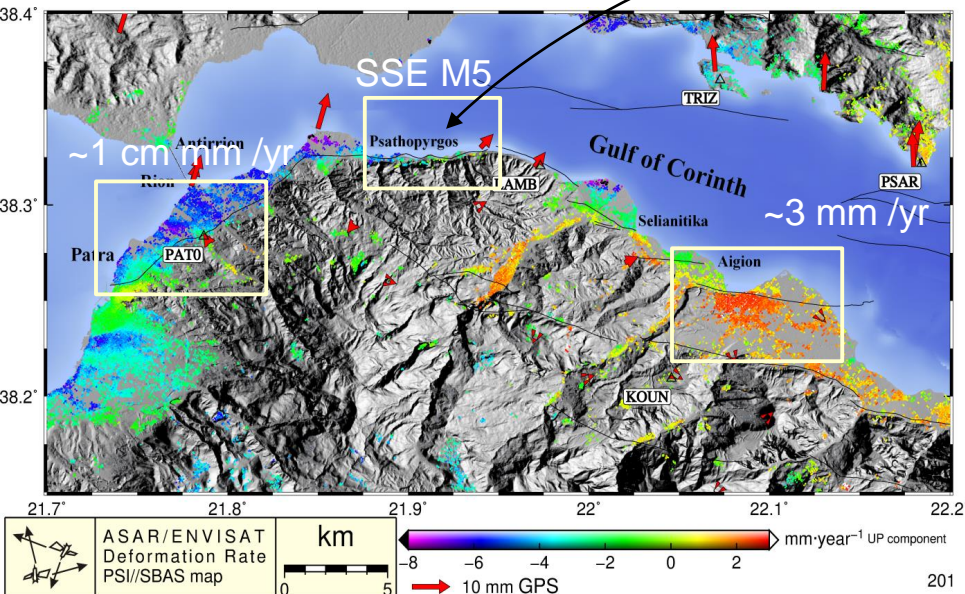
# Aseismic creep on major faults

from GPS, InSAR, and borehole strainmeters

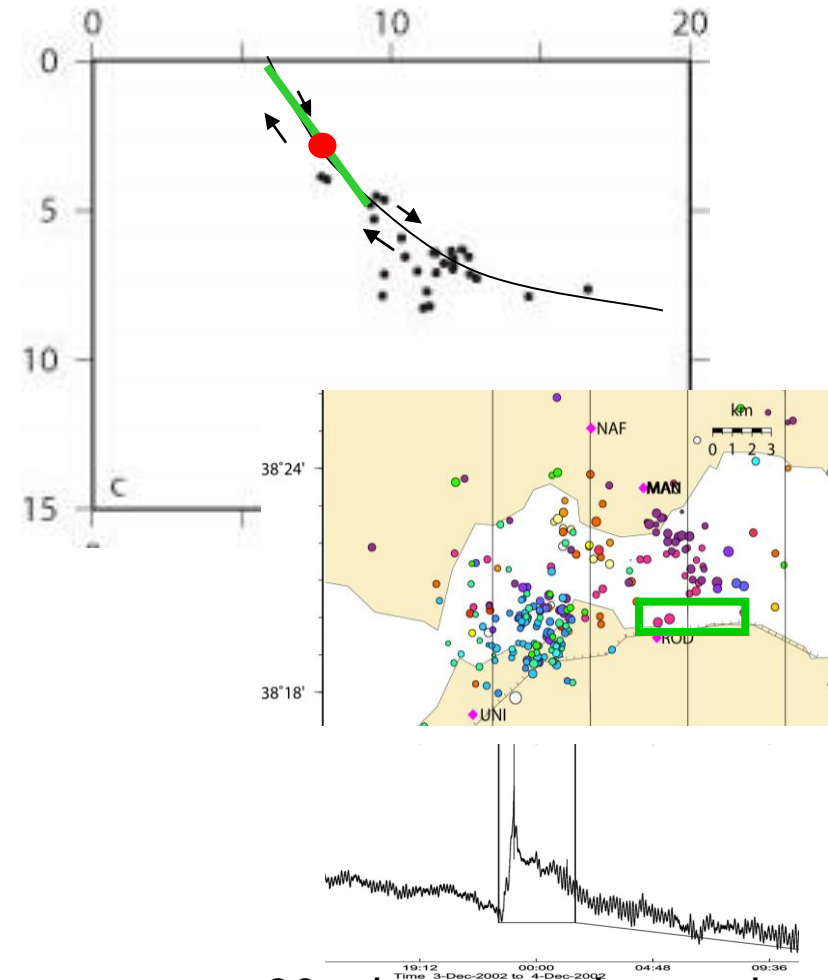
Psathopyrgos (transient creep), Rio-Patras, and Aigion faults

Bernard et al. 2007

2002 swarm  
and slow slip

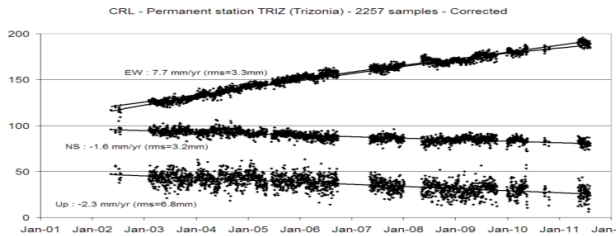


Elias and Briole, 2013



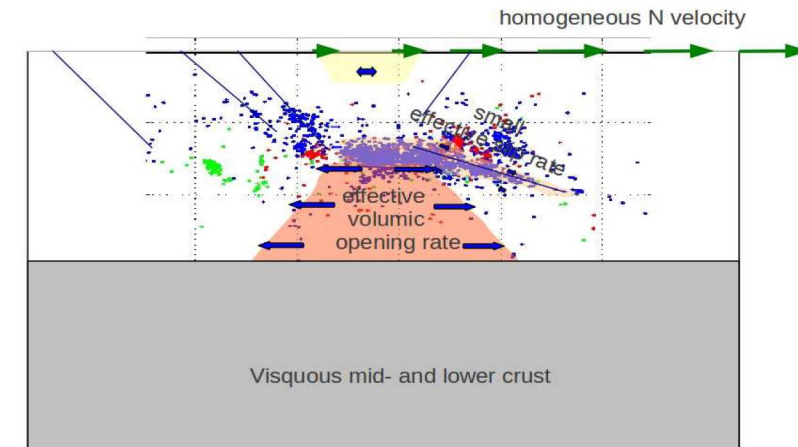
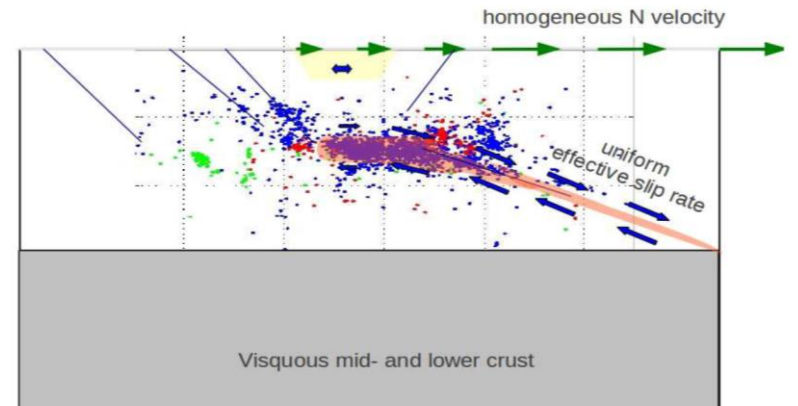
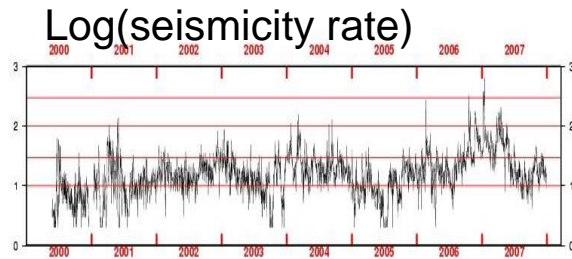
30 minutes strain transient  
on dilatometer,  $3 \cdot 10^{-8}$

# Rifting process



cGPS: **constant** stressing rate

Seismicity: **Fluctuating** stressing rate



Lambotte et al. 2013

The « detachment » is immature, growing downwards,

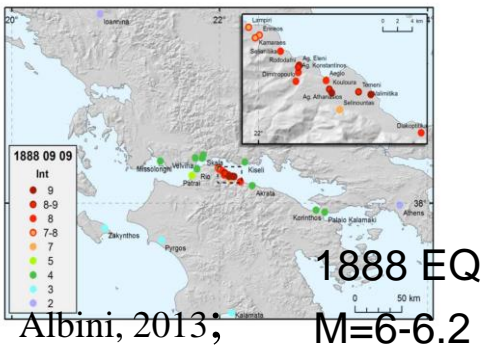
- not yet efficiently connected to the visquous mid-crust

-Another source of anelastic strain, below the seismic layer ... mode I ?

-Requires 3D mechanical modeling

# Earthquake rupture forecast for M>6

Boiselet et al. , 2014

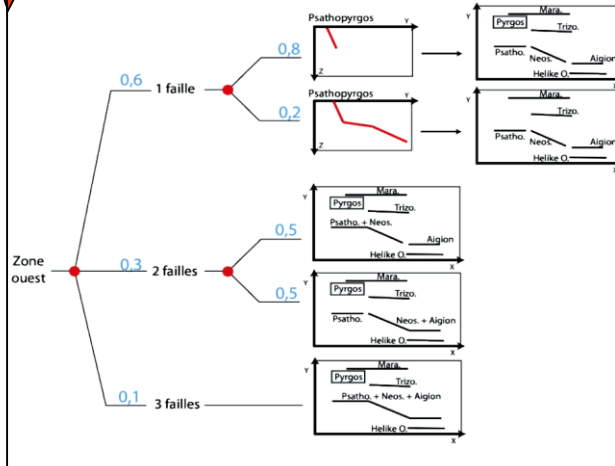


Albini, 2013;

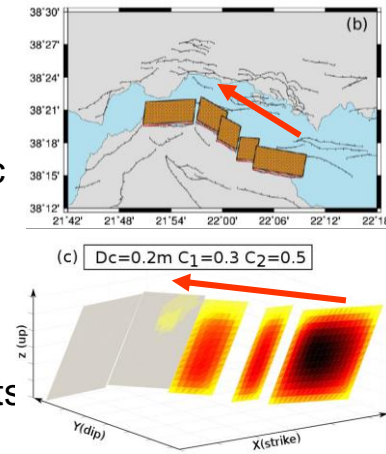
- Fault slip rates
- Strain from geodesy
- Paleoseismicity
- Historical seismicity
- Instrumental seismicity
- + uncertainties**

Logic tree:

- Seismicity models
- fault geometries

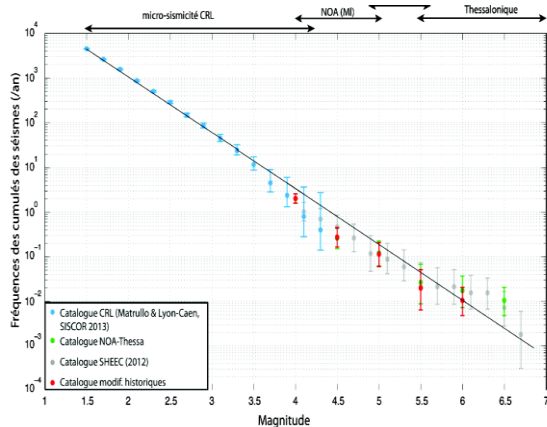


Dynamic ruptures across fault segments

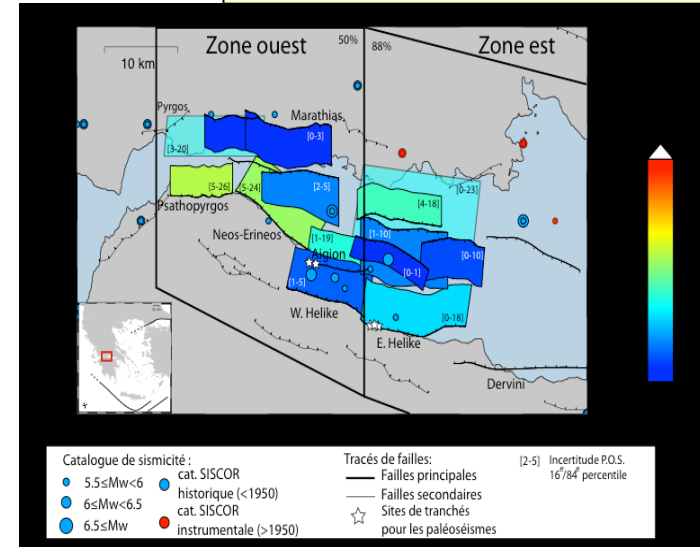


Durand et al. 2014

ERF forecast for individual faults

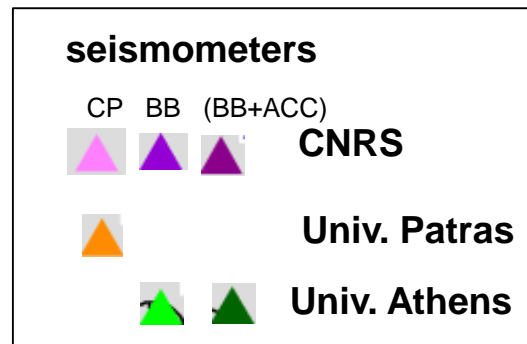
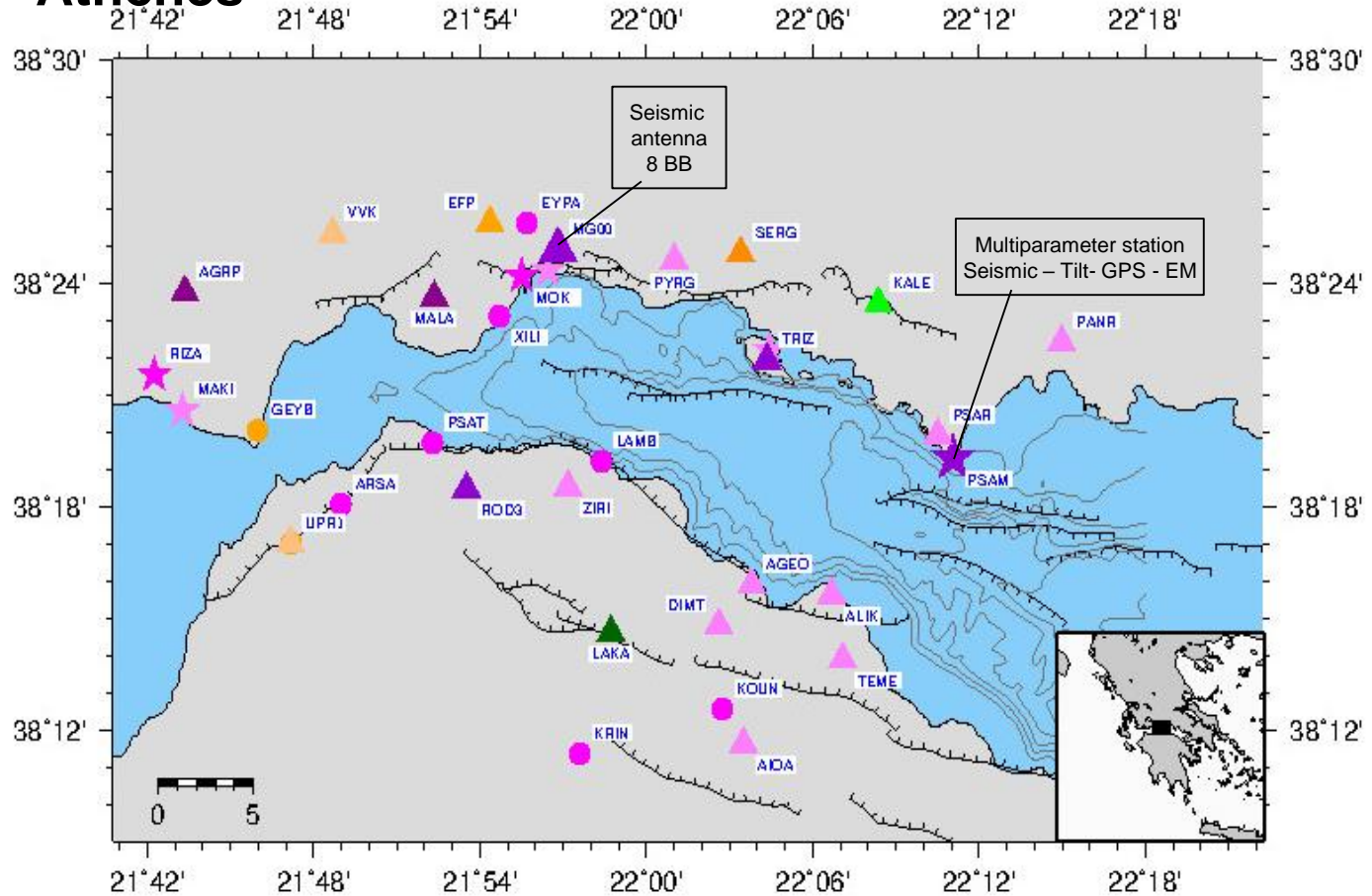


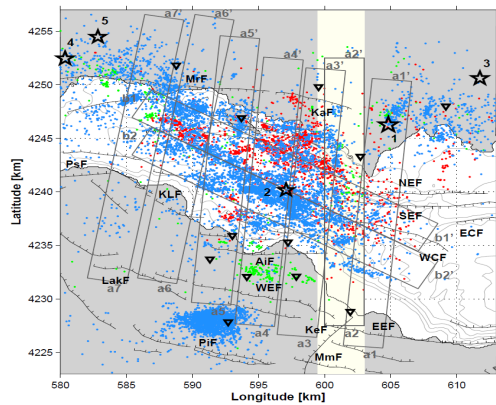
**P= 45 - 75 %  
pour M>6  
0 - 30 ans**



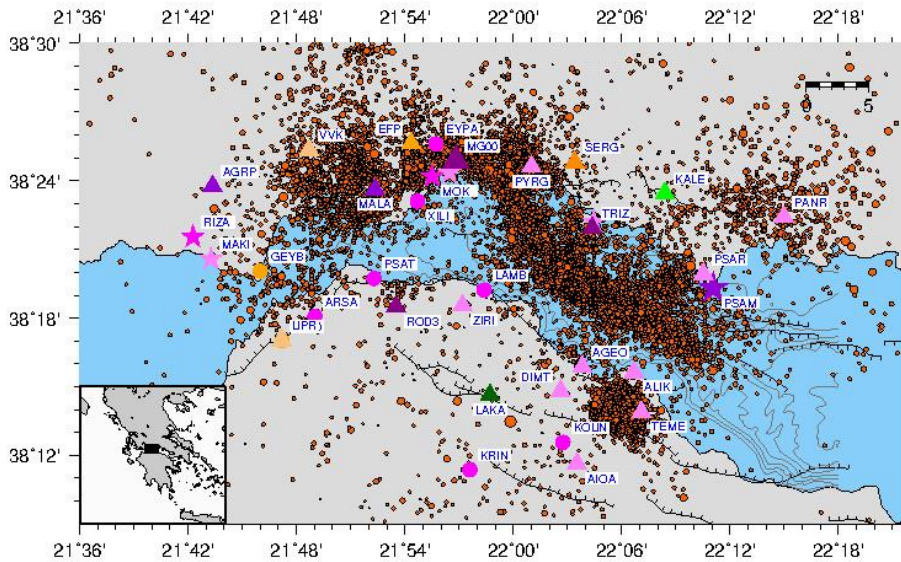
# RESEAU PERMANENT CRL – Télémétrie Univ.

## Athènes





2000-2006



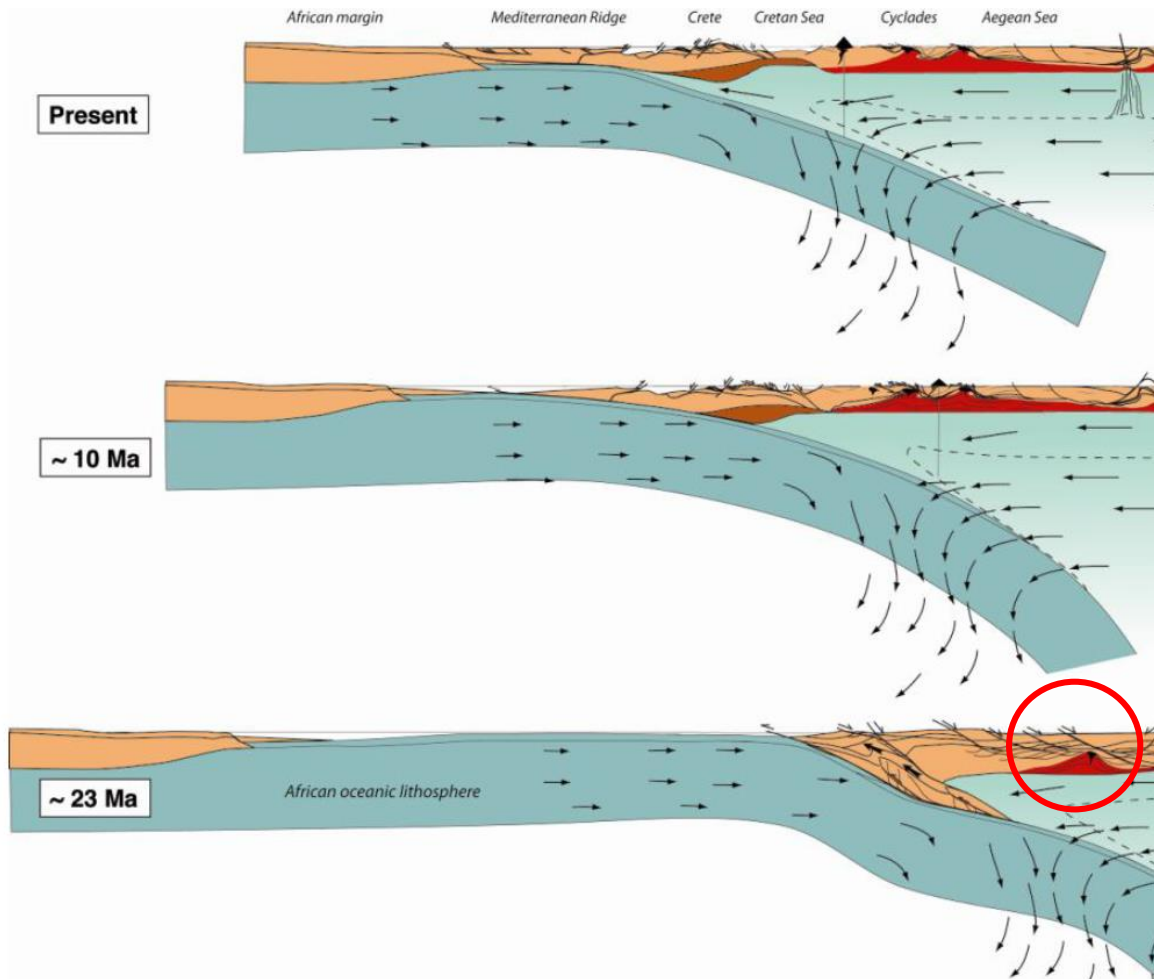
2010-2013

Réseau CRL étendu

[ftp://geoazur.unice.fr/pub/outgoing/deschamp/crl/film\\_sismib\\_2000\\_2013.lent.gif](ftp://geoazur.unice.fr/pub/outgoing/deschamp/crl/film_sismib_2000_2013.lent.gif)

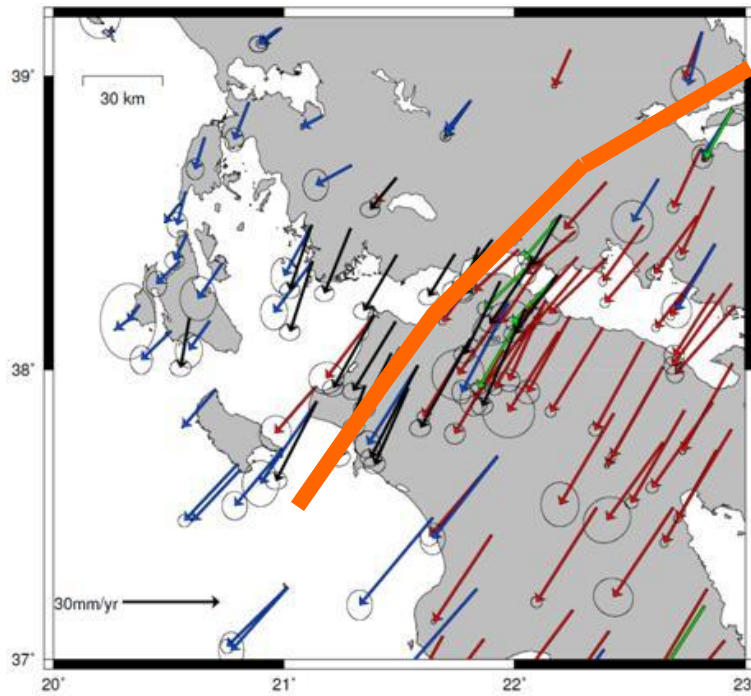
# The western end of the Corinth rift:

## The link to the subduction



The western end of the Corinth rift:

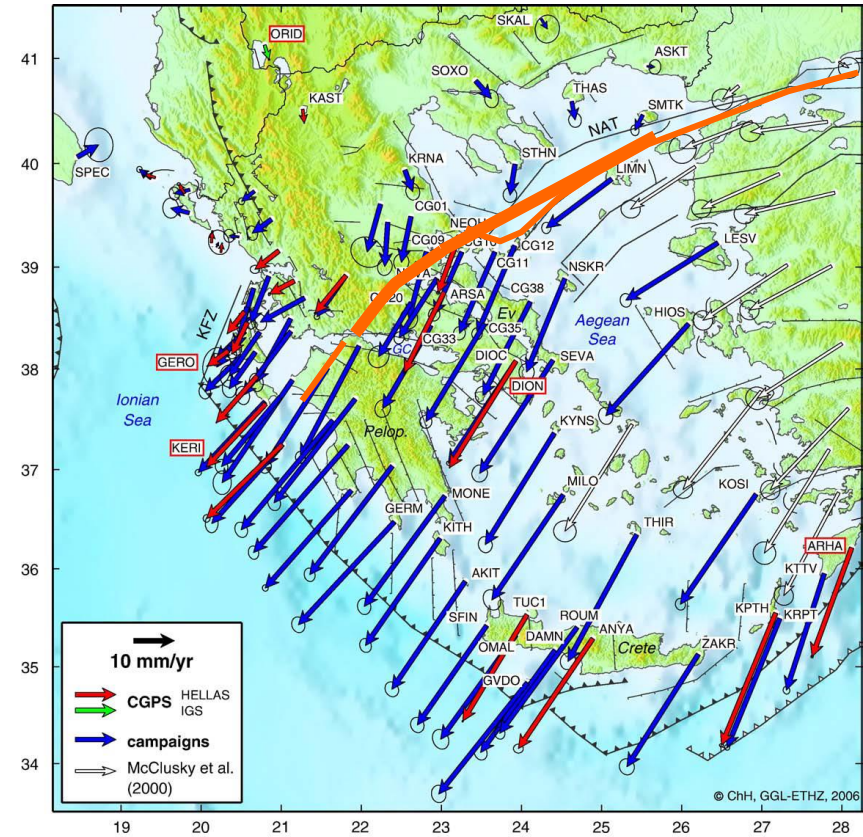
The link to the North Anatolian fault system



GPS :

Charara, 2010

Floyd et al., 2010



GPS 1993-2003

Hollenstein et al., 2008



# The western end of the Corinth rift:

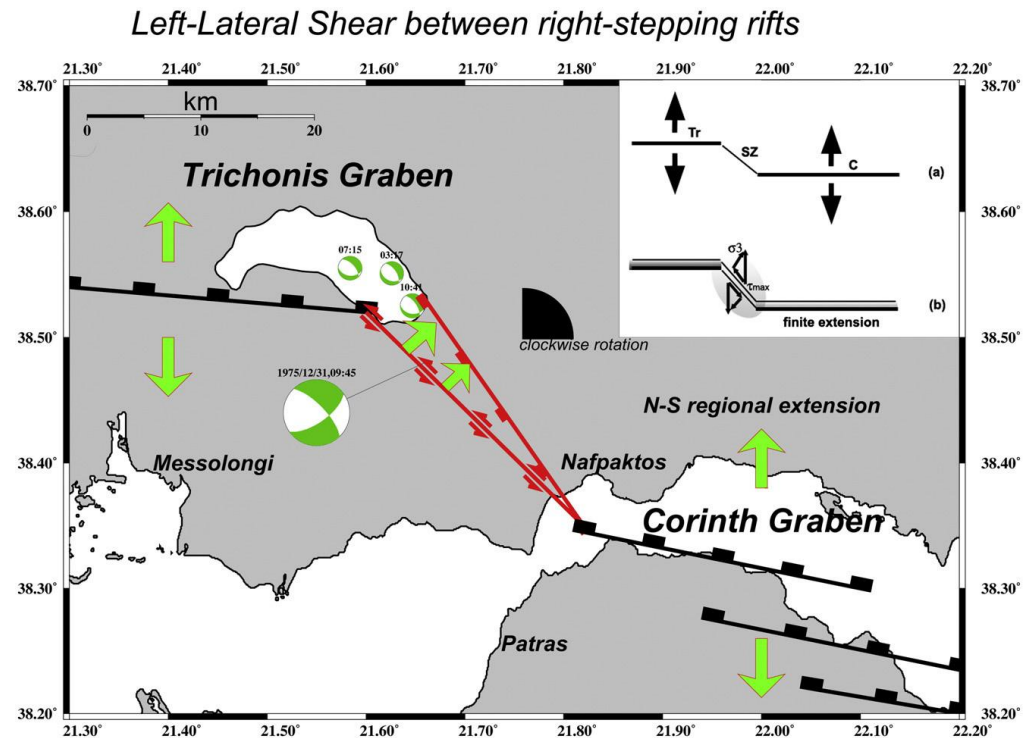
The NW connection to the Trichonis rift

Insights from the M=5, 2006 sequence

## Seismicity:

1975: M=6.0

2007: swarm, M=5.0-5.2



Seismicity:

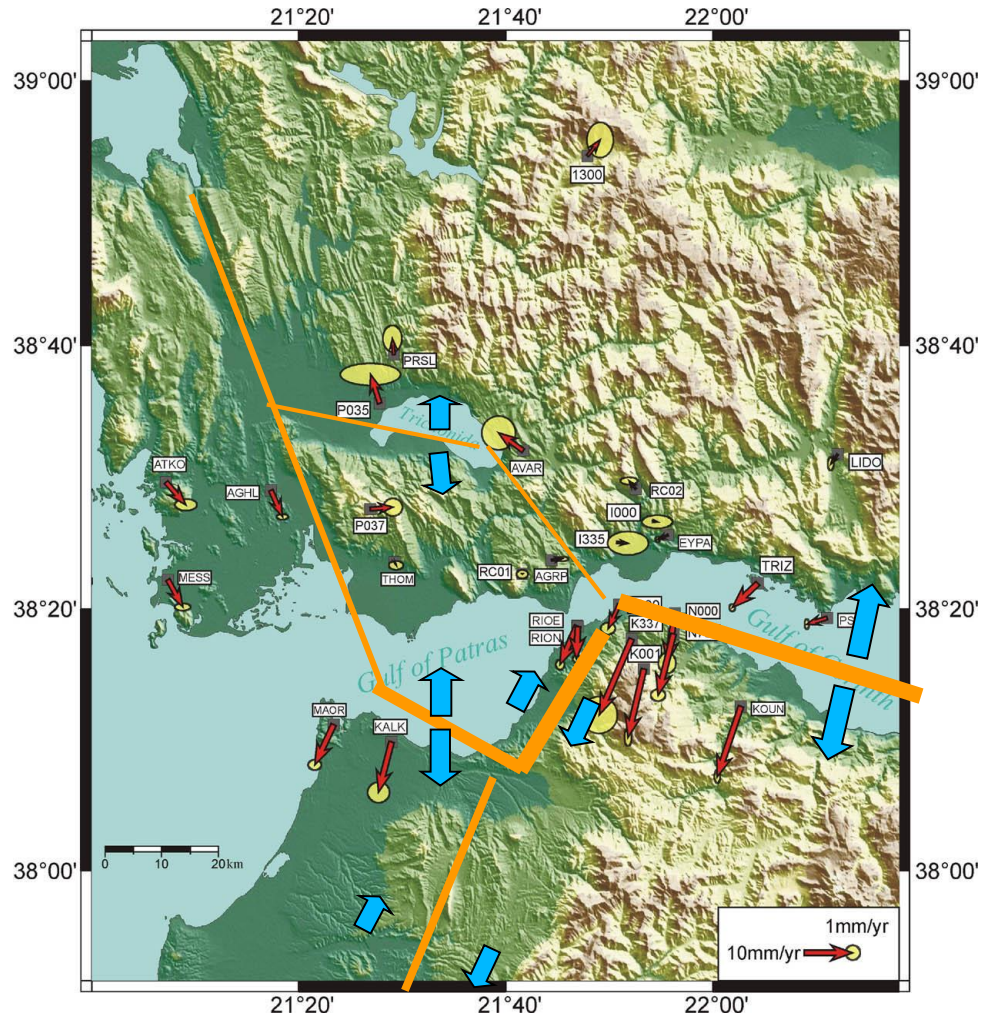
Kiratzí et al., 2008

Faults:

Doutsos et al., 1987; Lekkas & Papanikolaou, 1997; Goldworthy et al., 2002

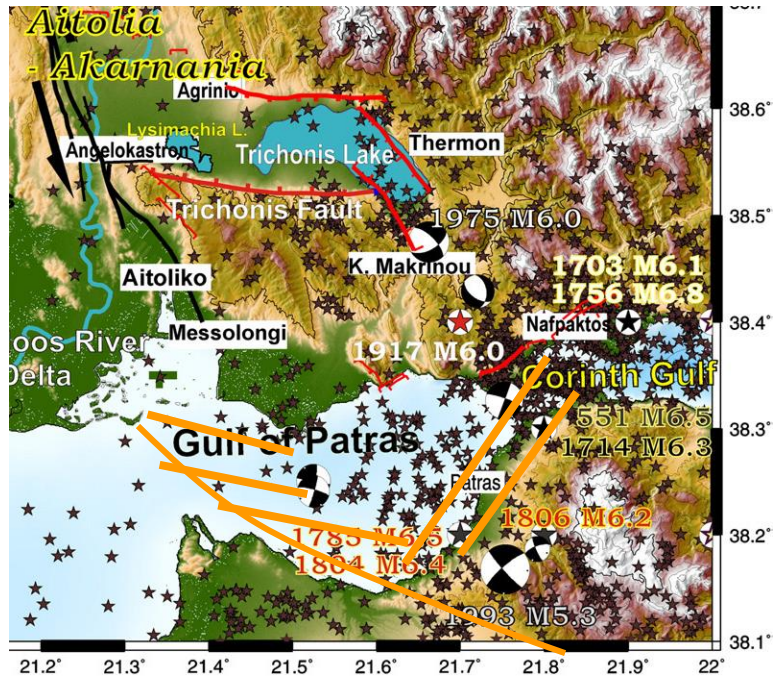
# The western end of the Corinth rift:

clues from GPS....



# The western end of the Corinth rift:

the western connection through the rift of Patras...



Kiratzi et al. 2008

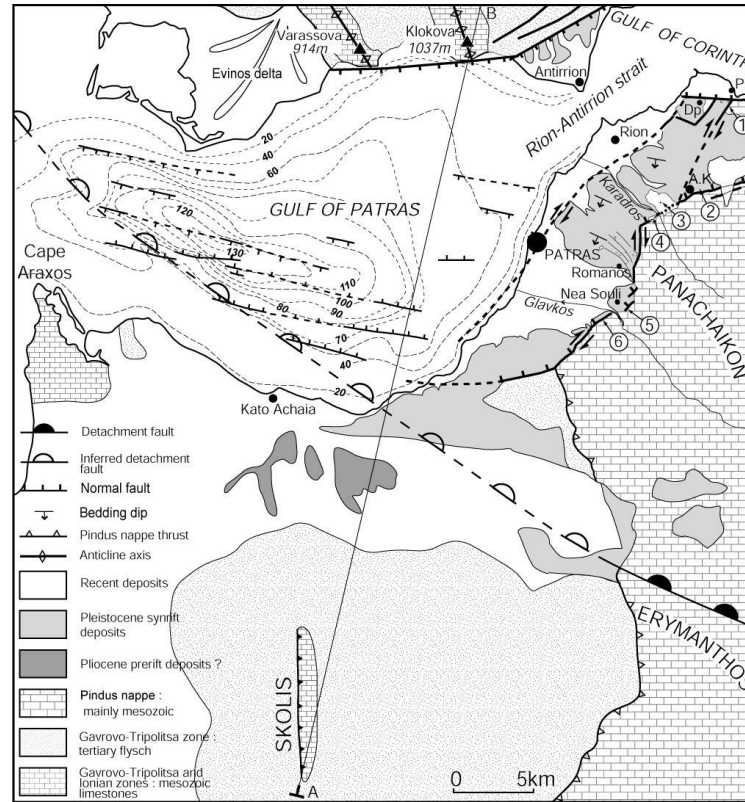
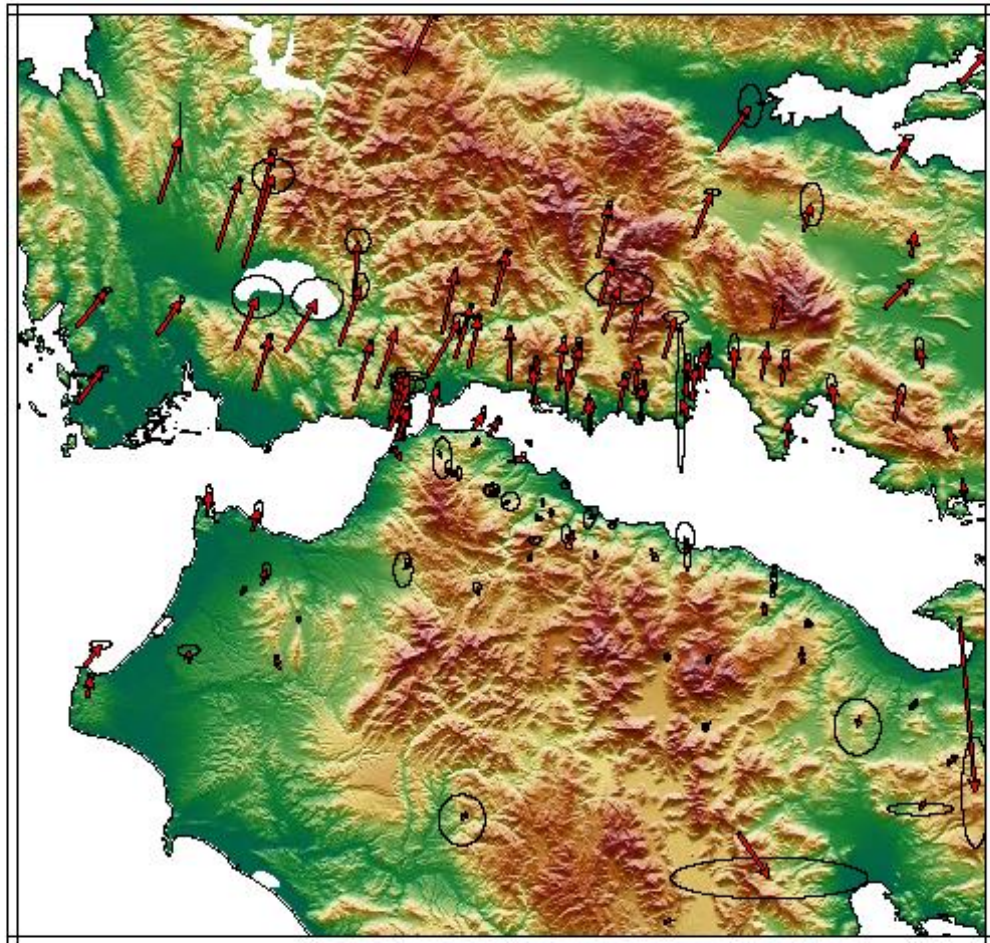
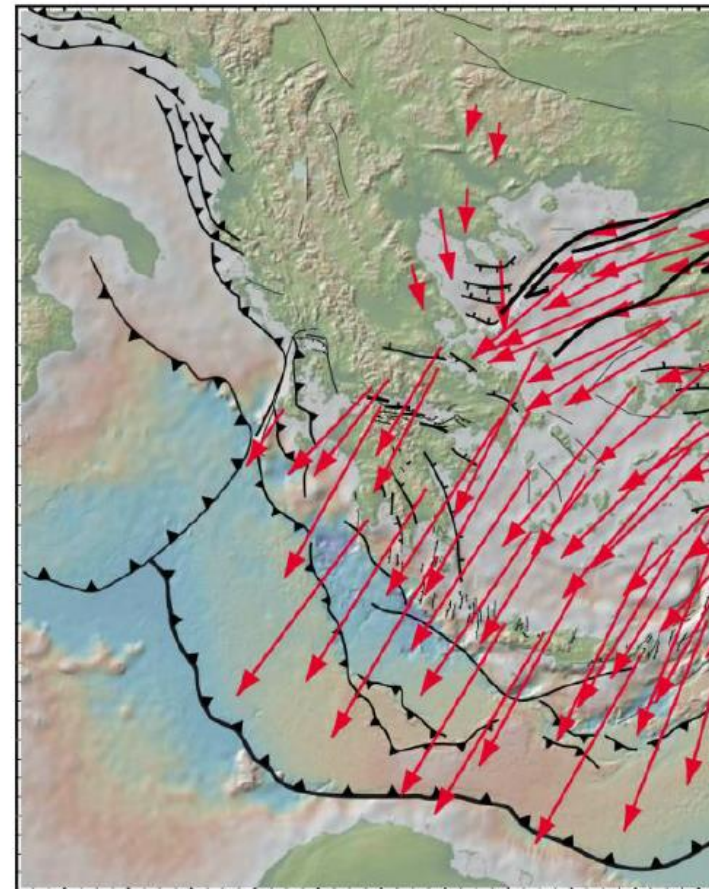


Fig.10-10. Simplified geological map of the surroundings of the Gulf of Patras. Geology after British Petroleum (1971), Dufore (1975), and personal data. Bathymetry after Ferentinos et al. (1985). Numbers 1 to 6 locate the areas described in the text. A-B: location of the section of Fig.III-69.

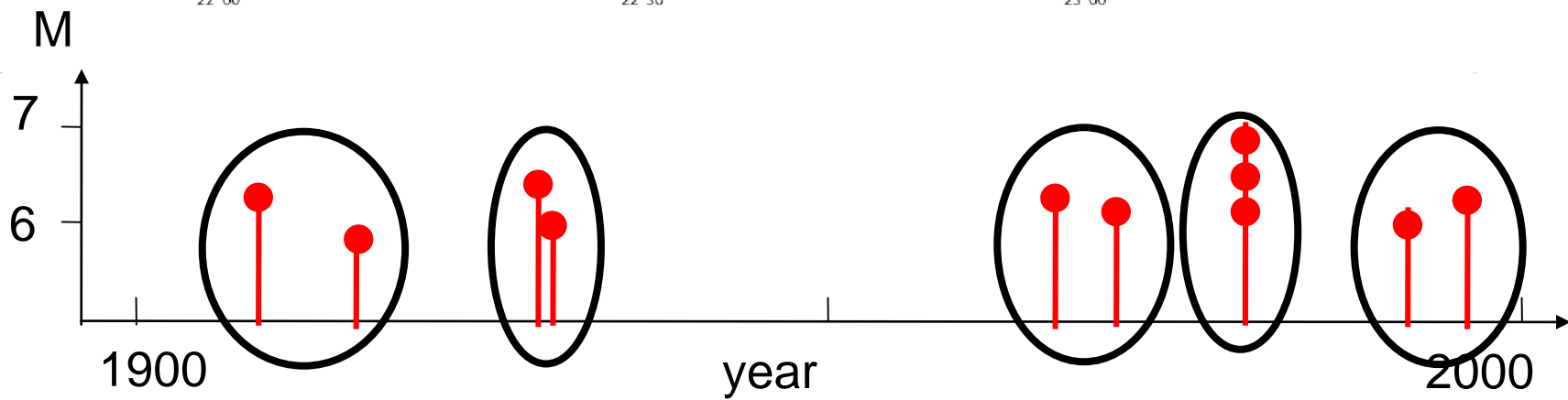
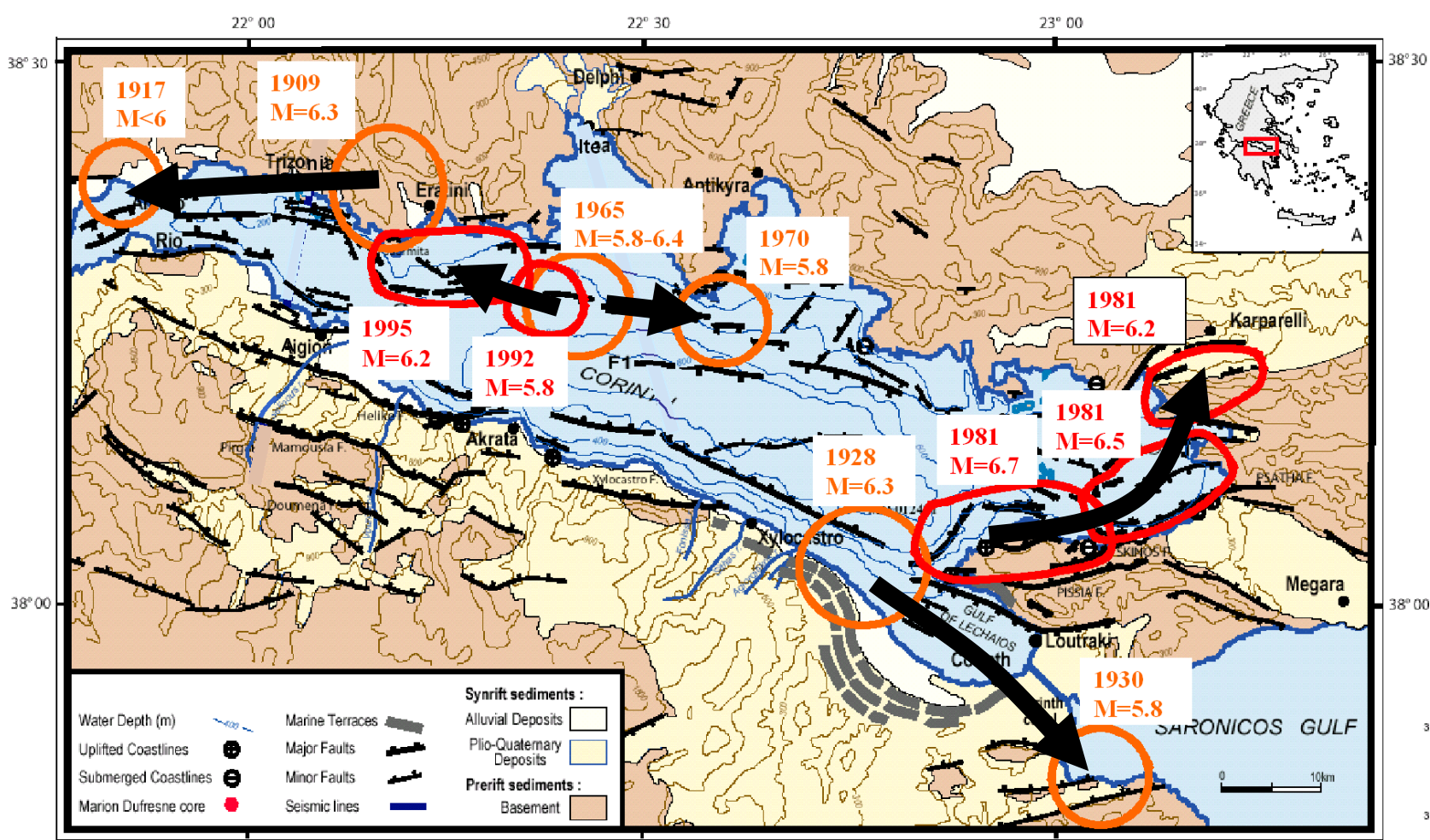
Flotté, 2005



GPS, Briole et al 2013

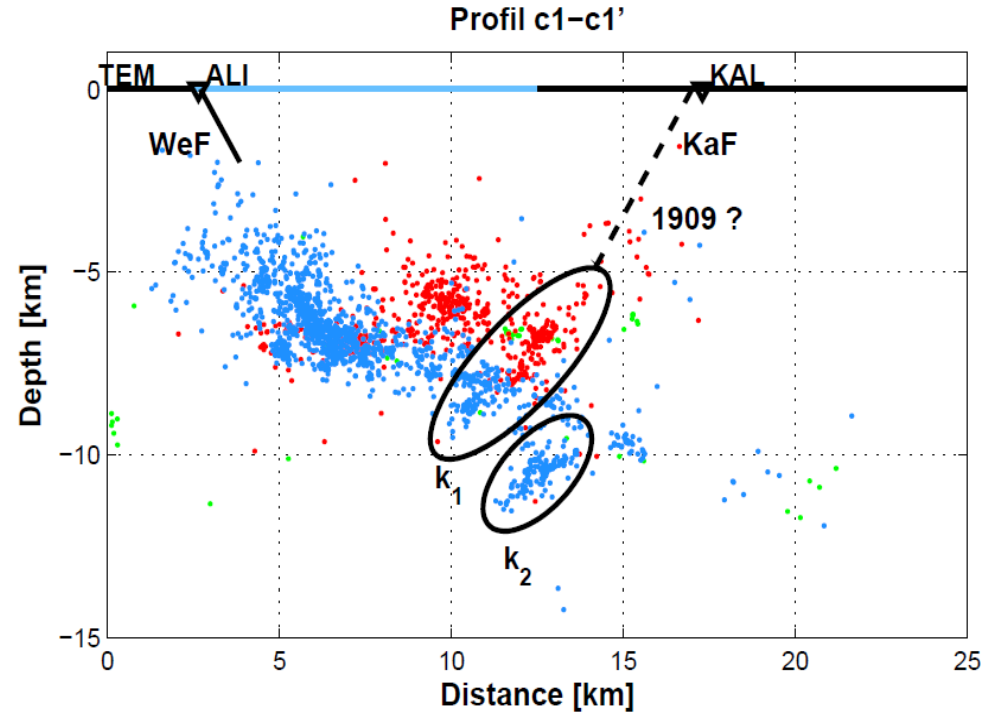
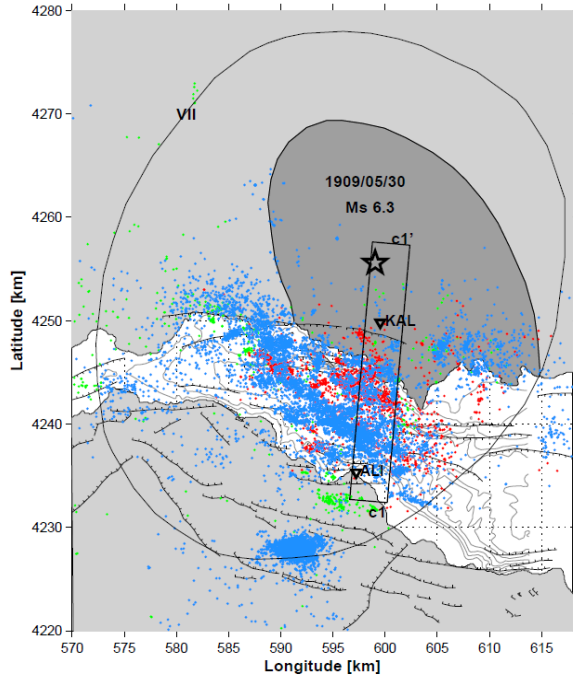


GPS d'après Reilinger et al. (2010).



# The Fokida Earthquake, 1909, M= 6.- 6.5

- on the Kalithea fault?



Quelles sont les zones bloquées:

### **Analyse temporelle**

-GPS - sismicité: fluctuations différentes

- strainmeter:

les courtes crises n'ont pas de déformation marquée

Donc l'activation sismique ne se produit pas sur la zone source de l'ouverture du rift.

### **Analyse spatiale:**

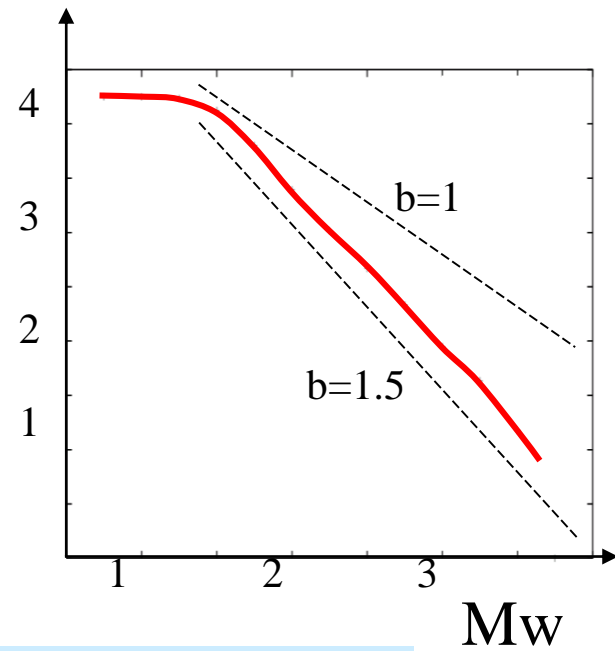
Les zones bloquées des asperités sismiques de 2010 et 1995 ne semblent pas le GPS



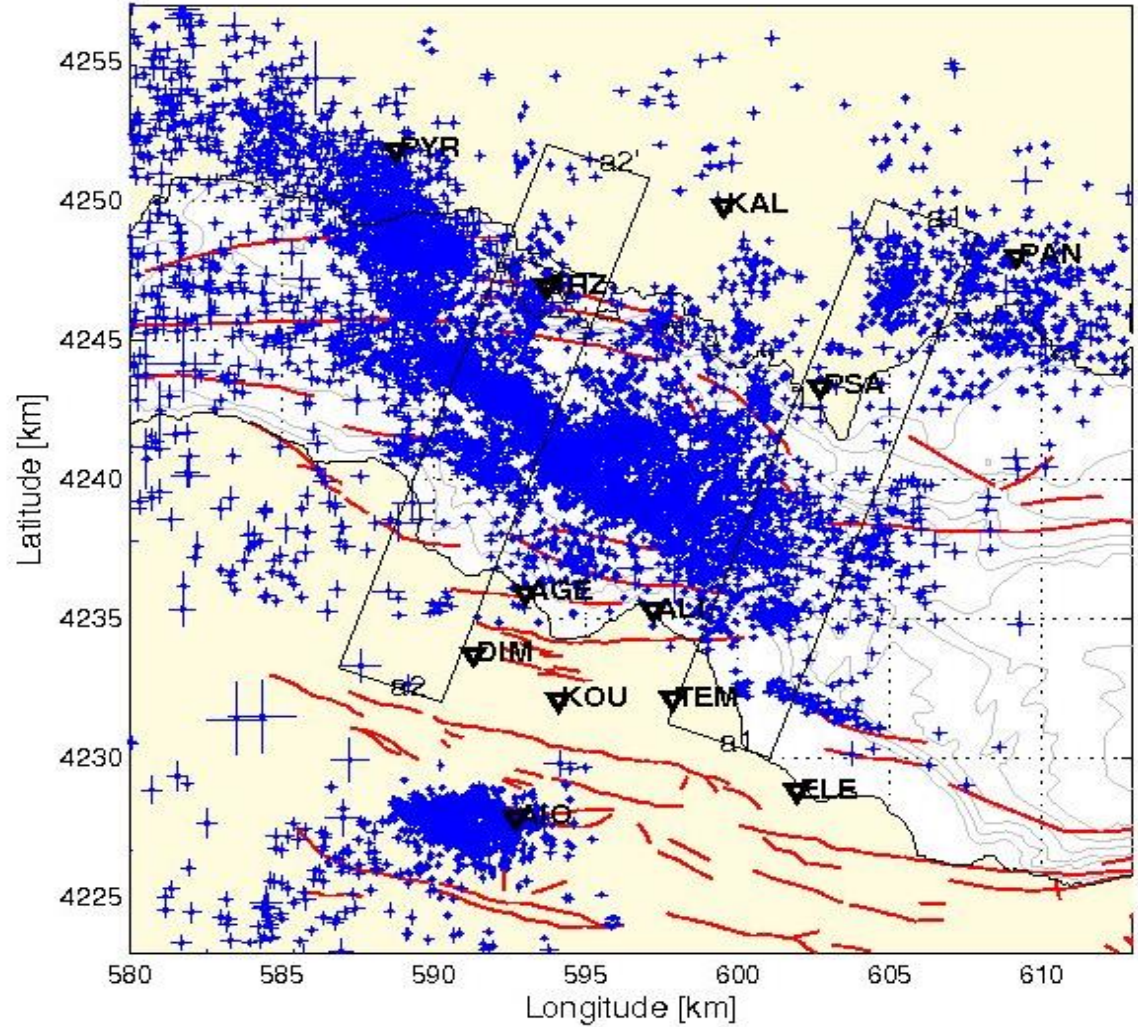


# Seismicity 2000-2007

$\log(N)$



Pacchiani, 2006



Lambotte, 2008

Space clustering

Fault segments

1 km – 10 km

Time Clustering :

Pulsations

year-month-day

3 large crisis:

2001

2004

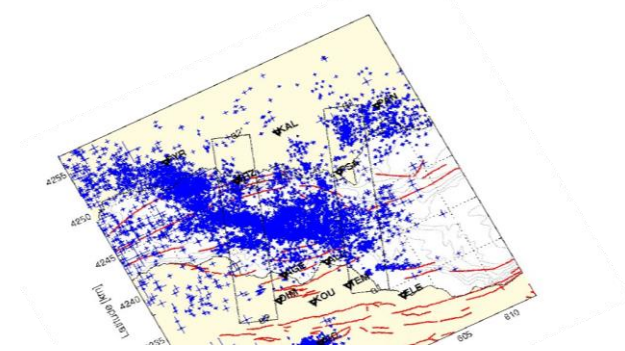
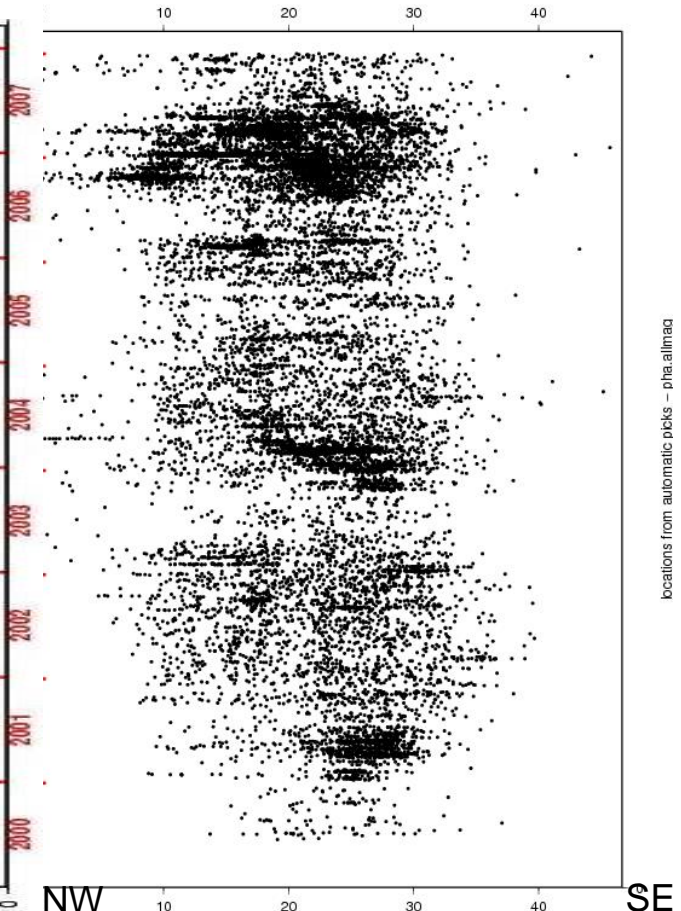
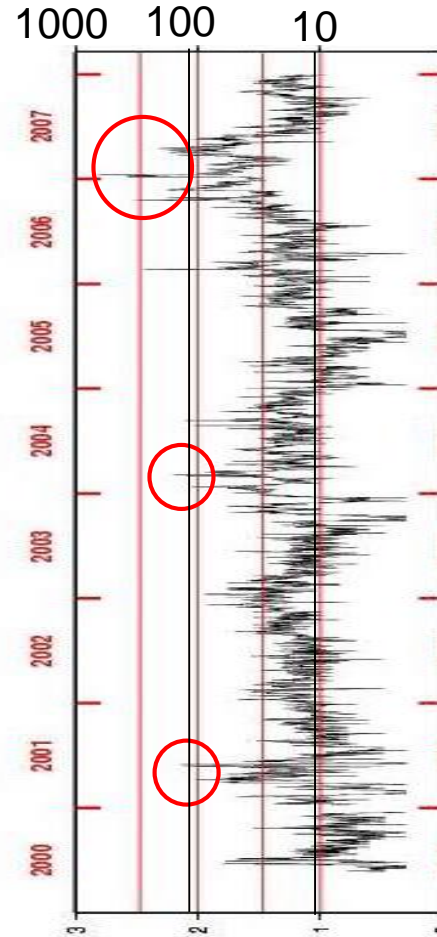
2007

Time  
(year)

seismicity rate  
/day

2008

2000

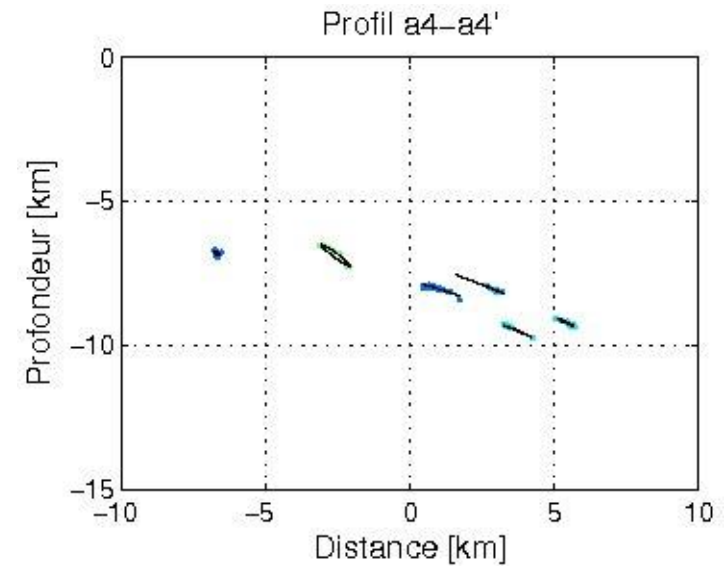
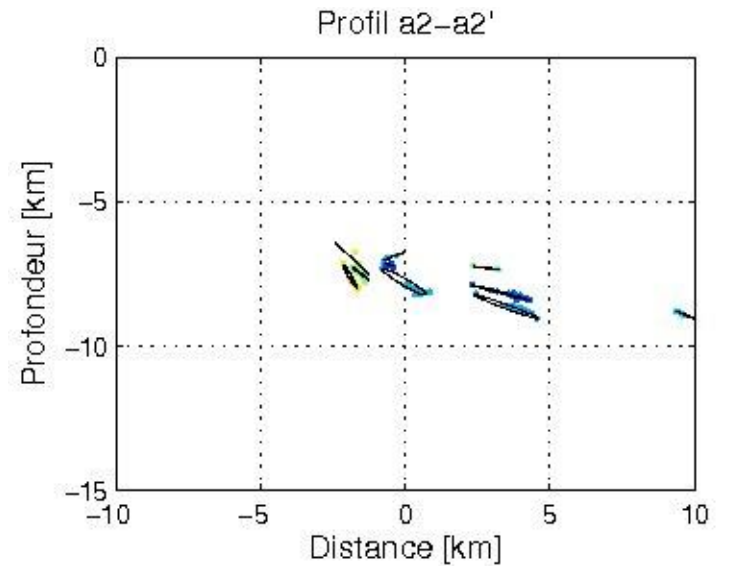
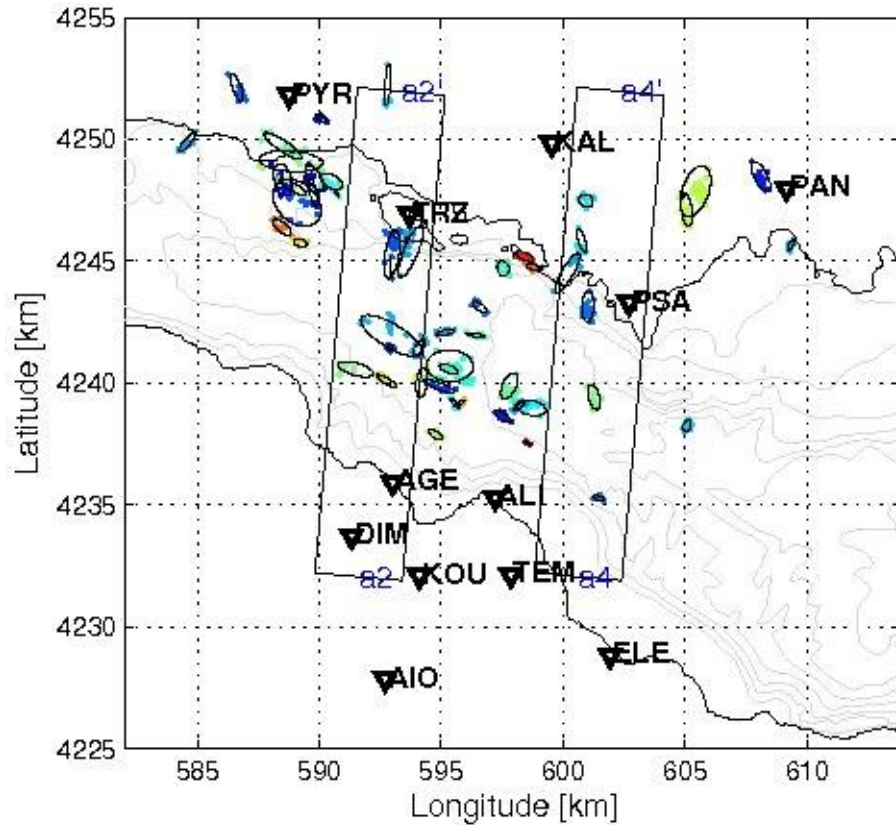


locations from automatic picks - pha.allmag

distance in km from (lat38.45 lon21.9) in az125

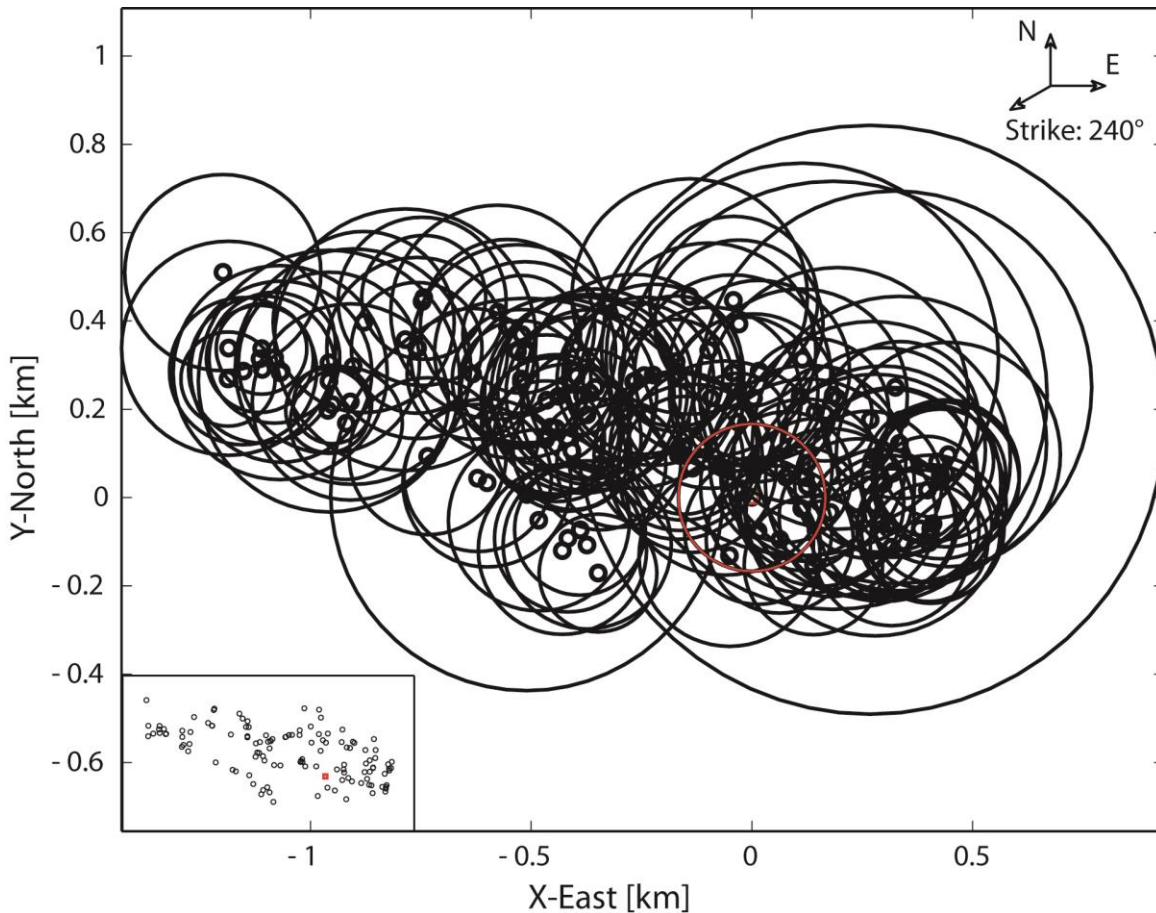
Multiplet fine structure:

Kilometric, north dipping fault segments



Lambotte et al 2009

# spatial distribution of events in one multiplet



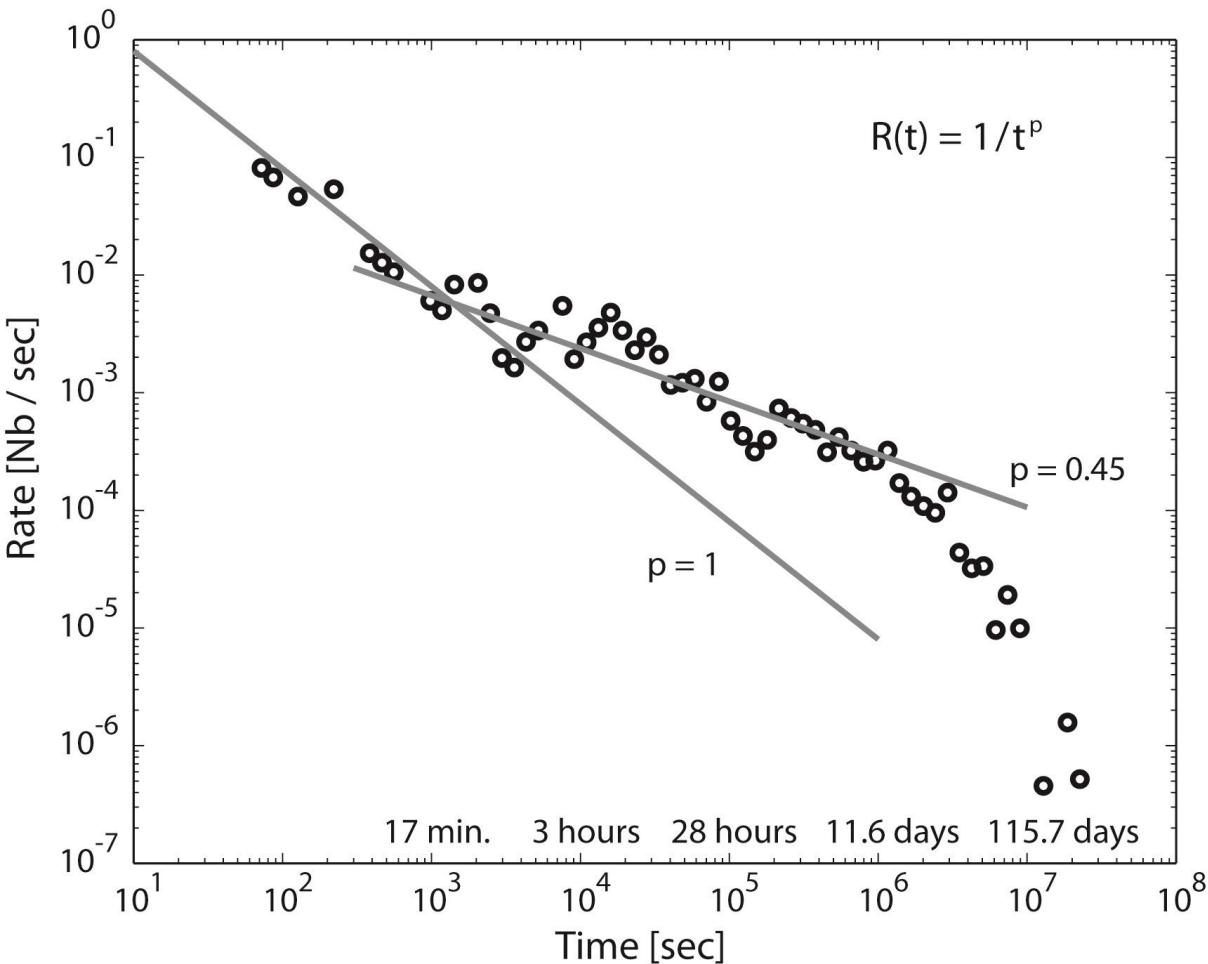
Multiple rupture of the same asperities

- Multiplet seismicity is forced/controlled by creep

- Creeping area must be larger than the seismic area

- Creep may be triggered/assisted by pore pressure

# Decay rate stacked for all multiplets



No standard Omori law:

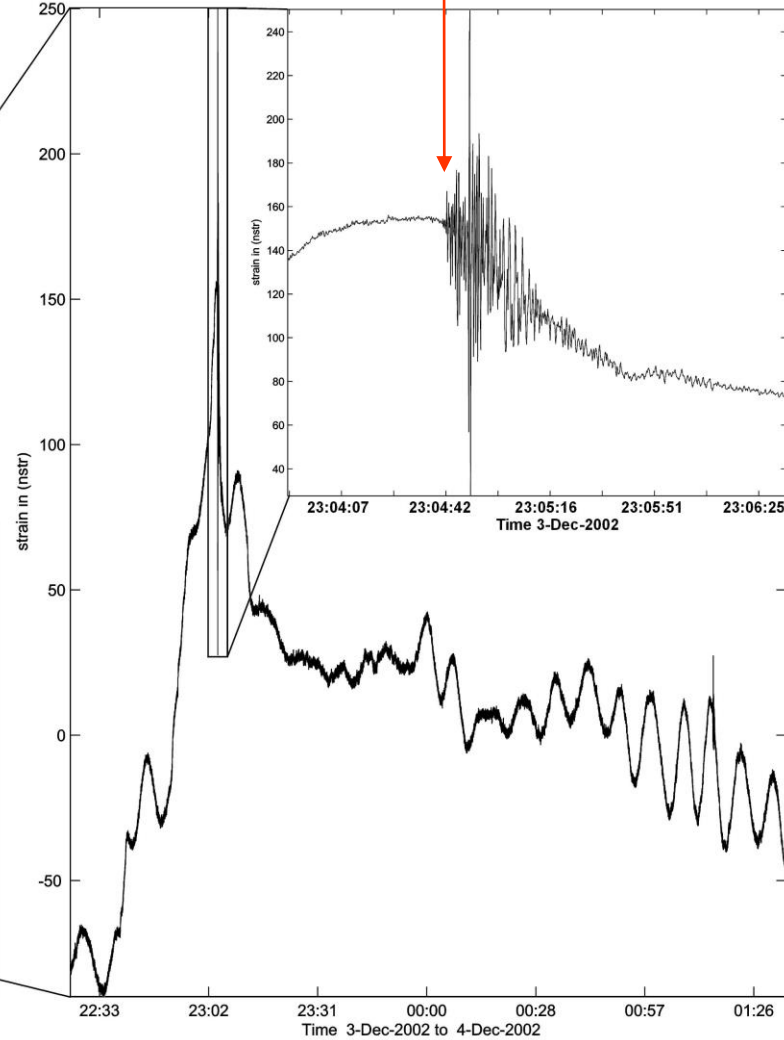
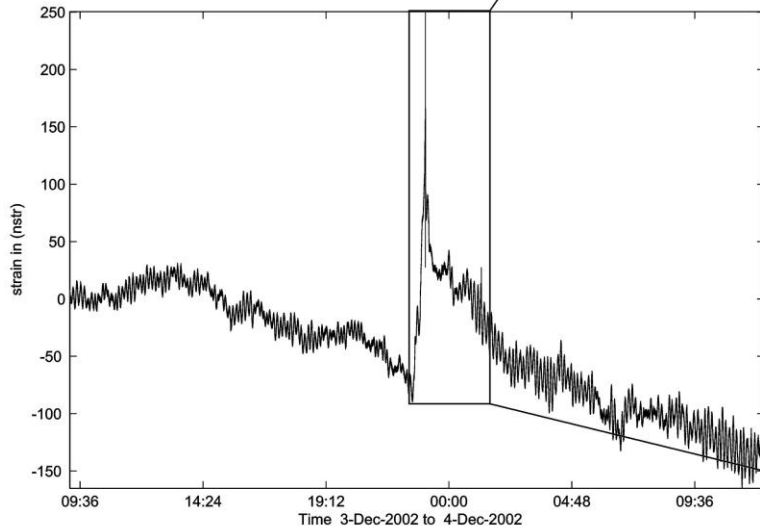
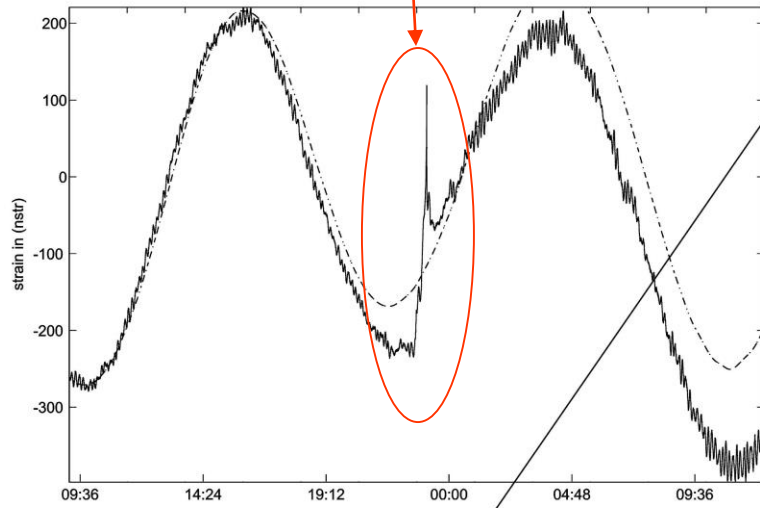
persistent process  
after 100 – 1000 s

- Multiplet seismicity  
forced by creep

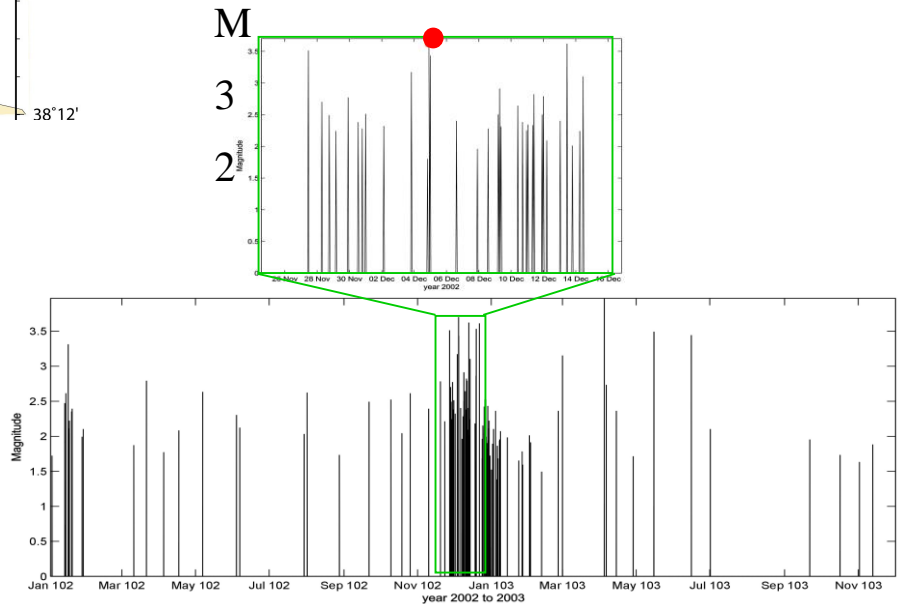
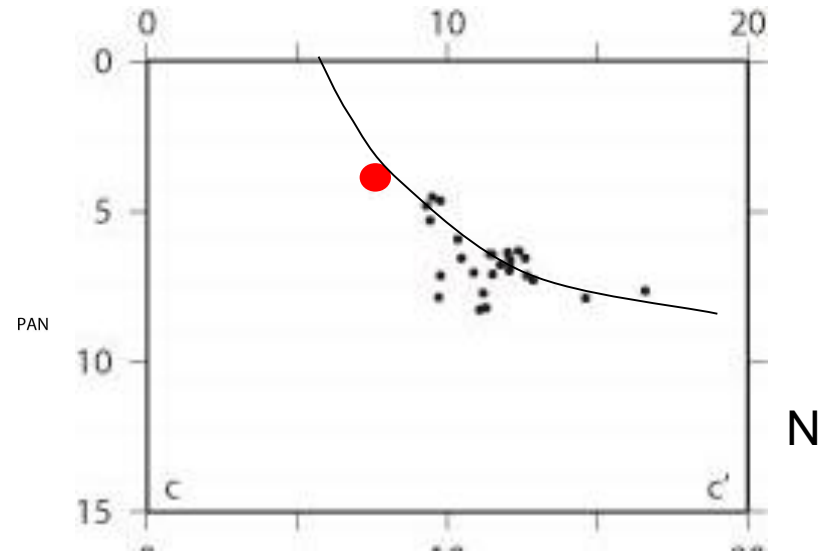
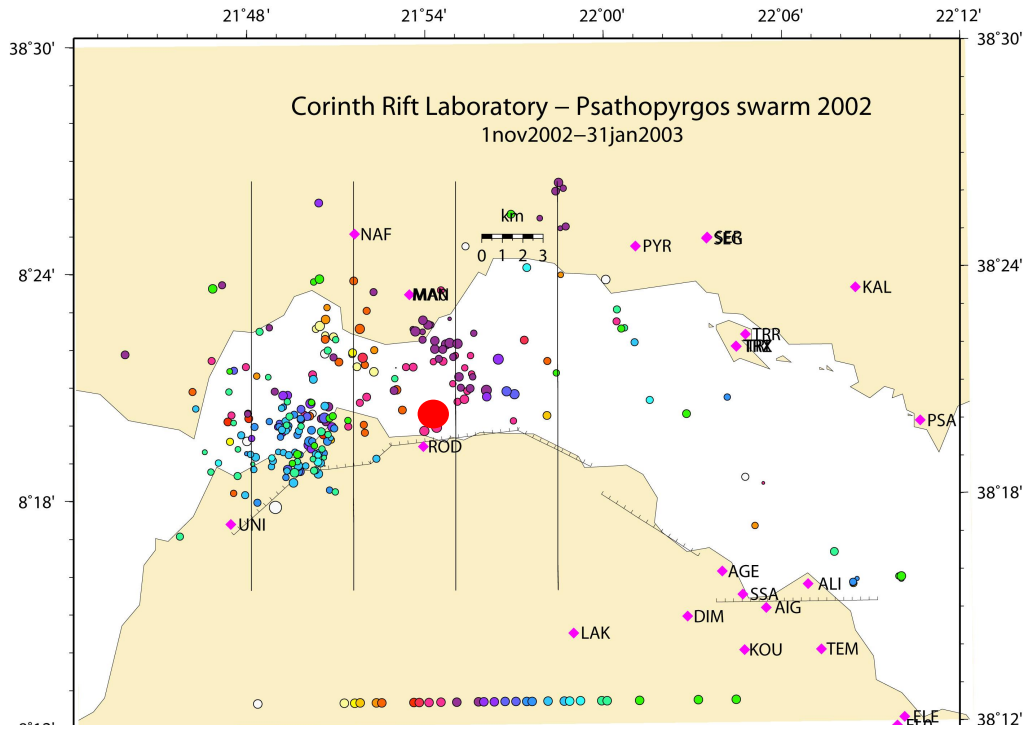
- Creep triggered by  
fluid flow

Transient strain

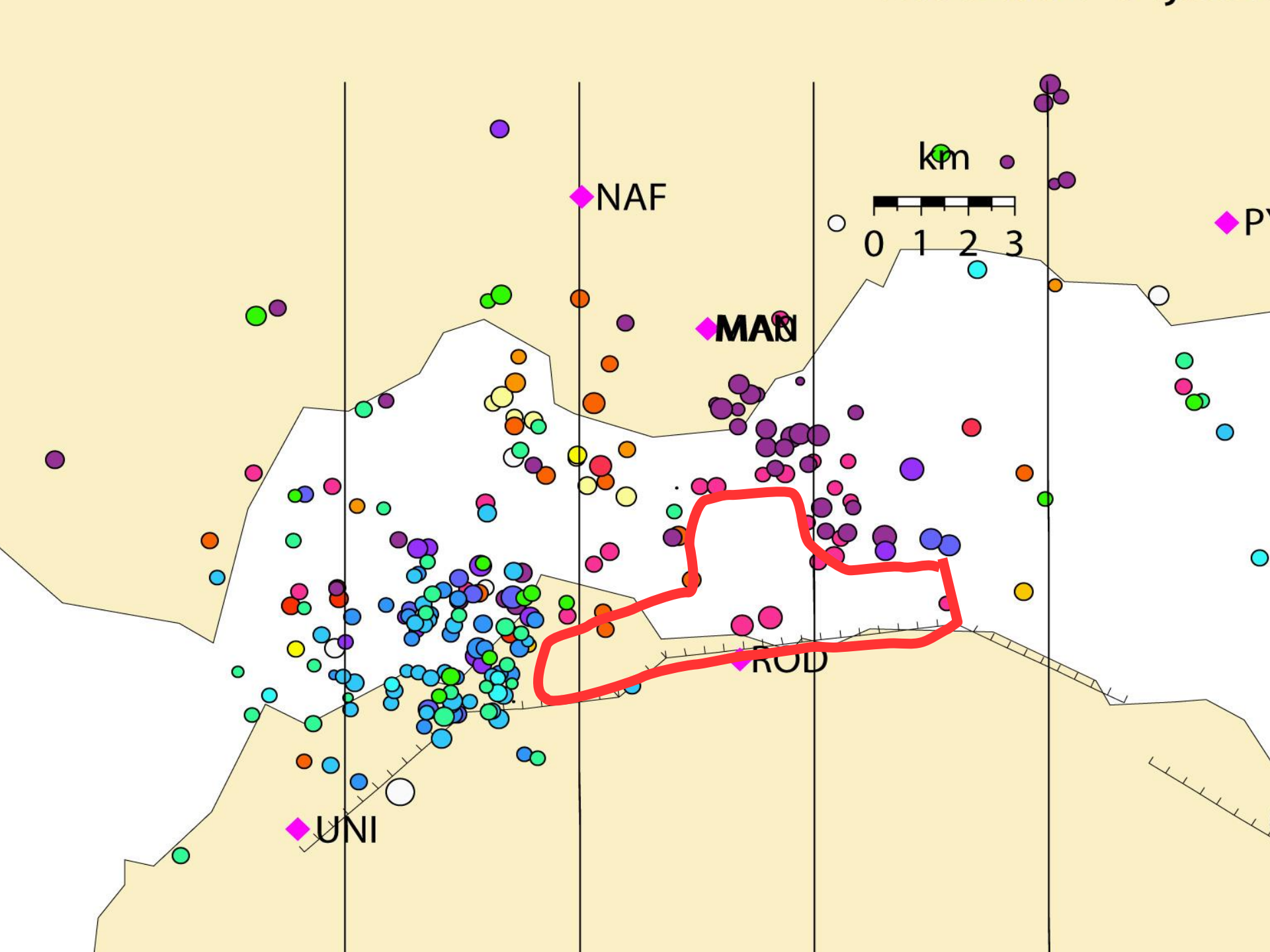
3 dec 2002 23:04 –  
M=3.5



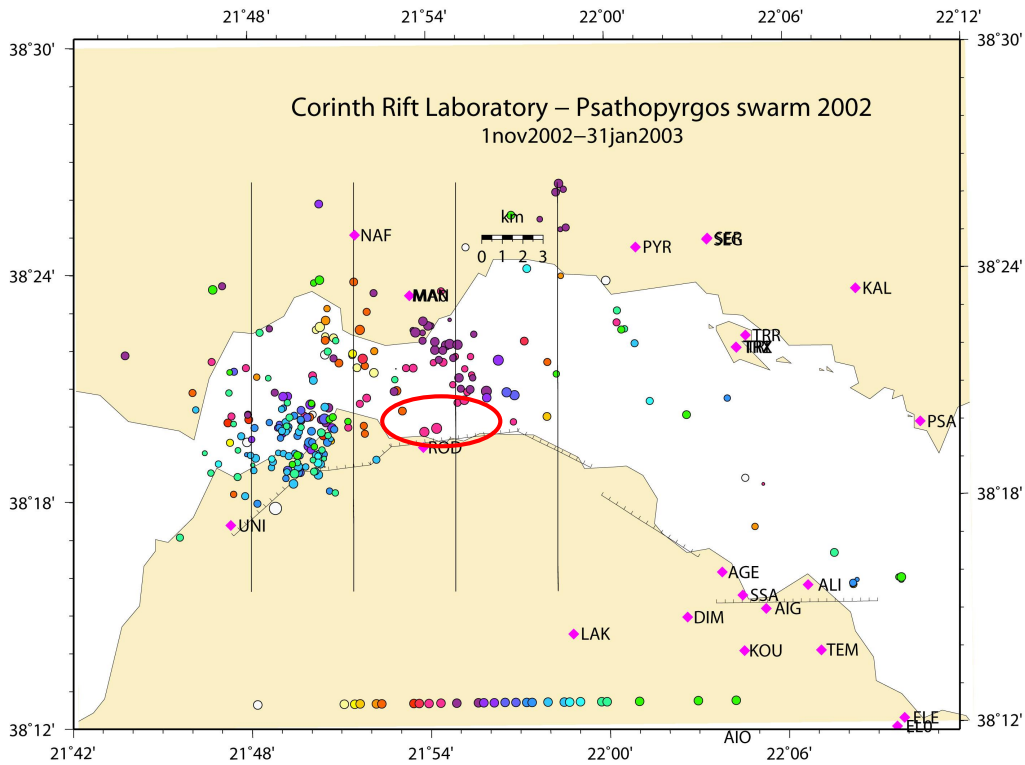
# Psathopyrgos fault seismic swarm - Nov-dec 2002



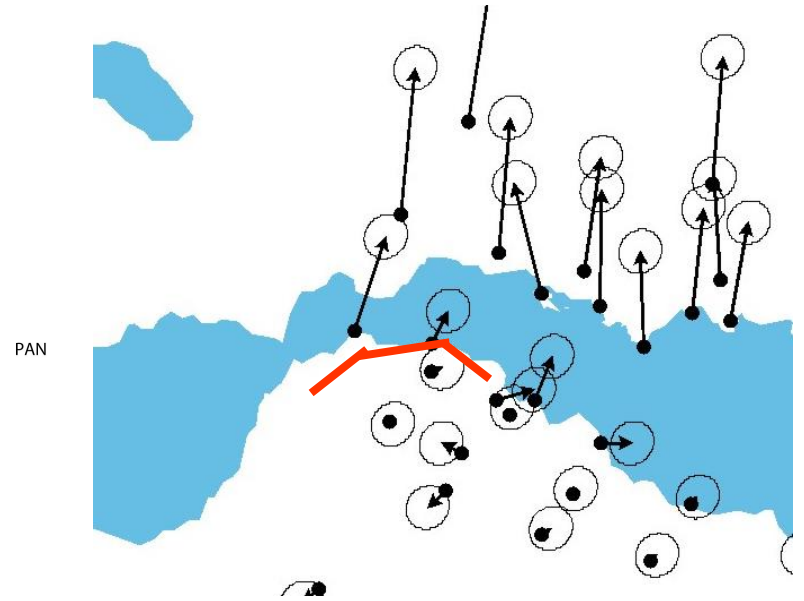
2002-2003







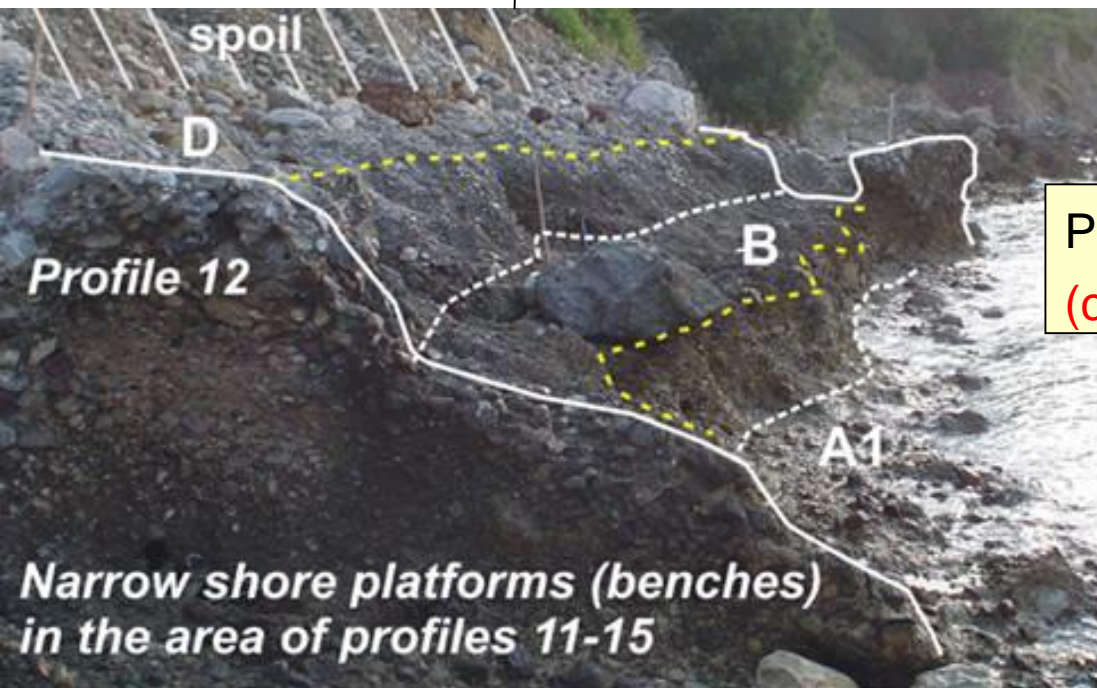
**Nov 02 – jan 03:  
Creep from seismic swarm  
and strain transient**



**Avalone et al. 2004**

**1990-2001:  
Creep from GPS  
1.5 cm/yr**

# Psathopyrgos fault



Paleoshoreline: 5 rapid uplift episodes (coseismic?) in 2000 years

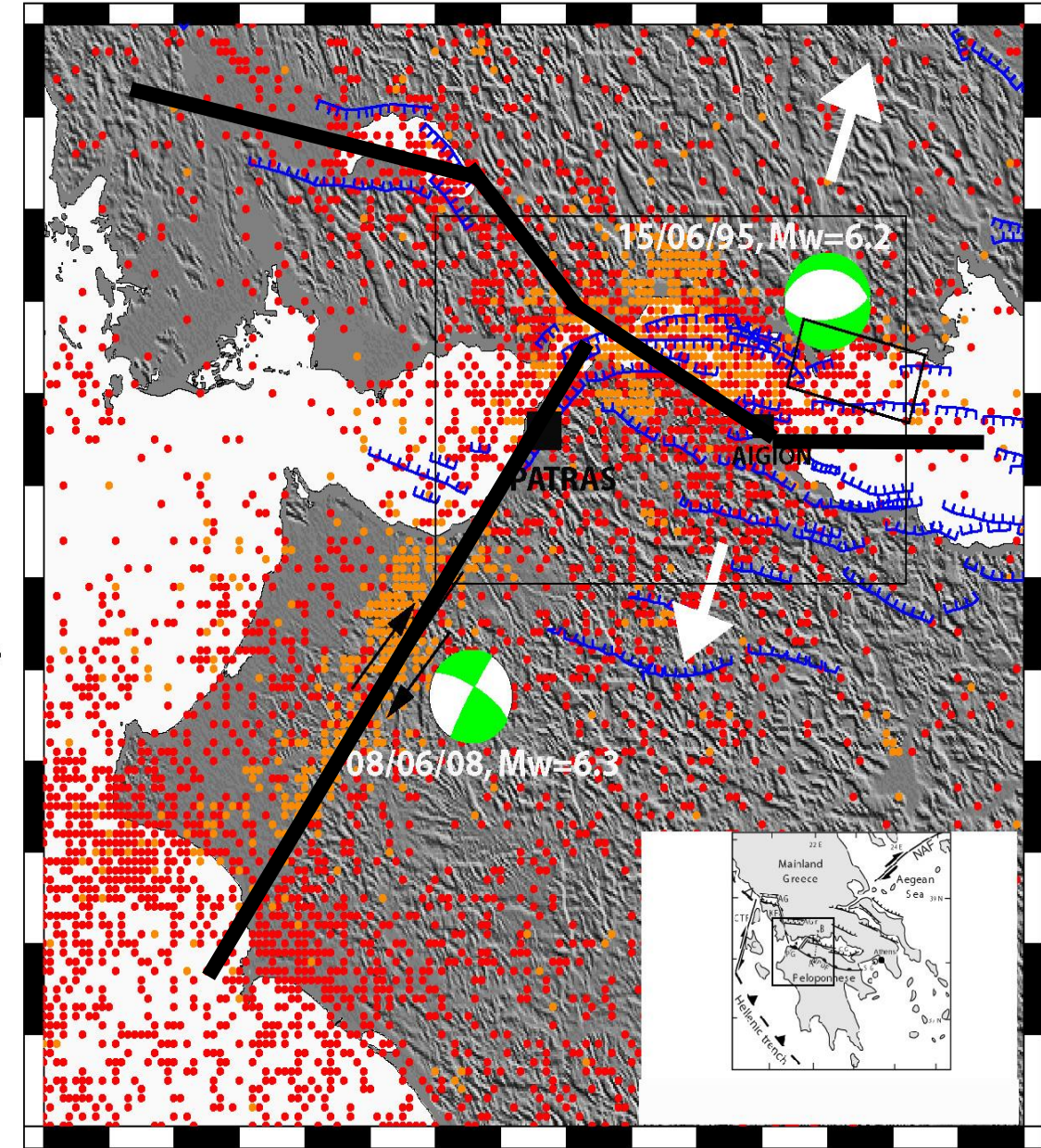
Pantosti et al. 2007

21°

22°

38°

38°



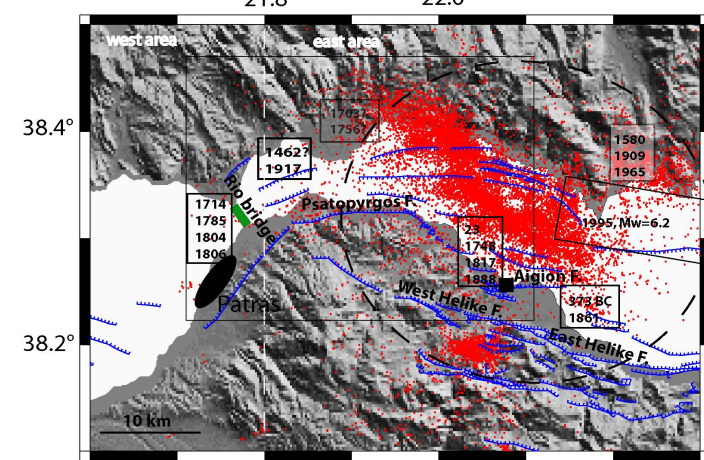
21°

22°

# CONCLUSION: normal fault activity

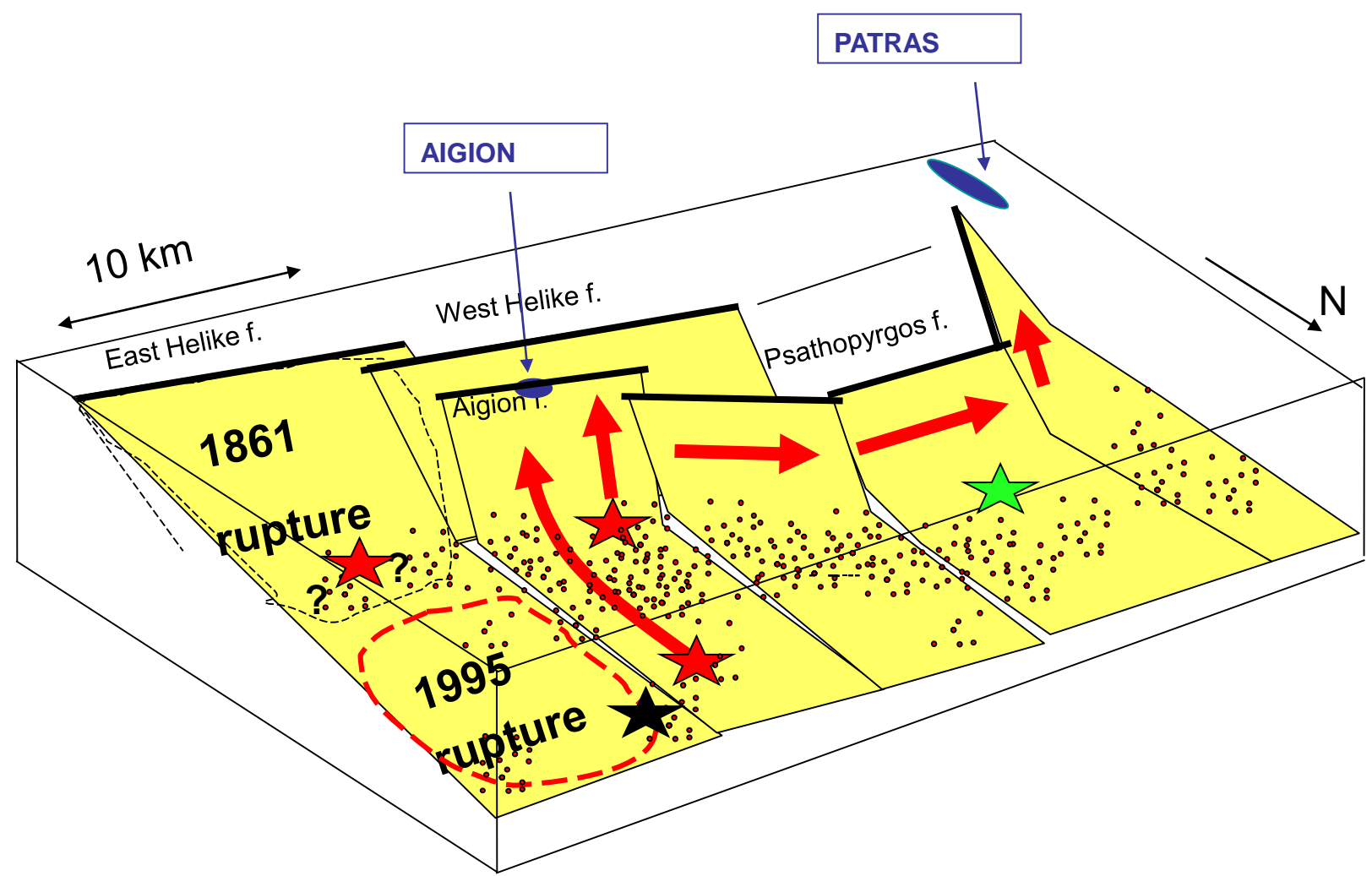
## at various space and time scales

- new fault system, oriented NW-SE, consistent with microseismicity: **what is its cause? when did it start?**
- large earthquakes: cycle 150-300 years; end of seismic cycle for several faults - **what probability for large rupture? Possibility of cascade events?**
- Steady fast creep or shear in 1-2 km thick layer, on small non coplanar faults at depth, loading the locked segments - **but does it relaxes strain on the deep segments?**
- microseismicity: migrating swarms (days, months) related to major creep events on faults, and strong east-west contrast: **is it related to fluid pulses, and why?**
- major transient creep (hour to weeks) on the Psathopyrgos fault: **phase of cyclic relaxation, or start of major failure?**

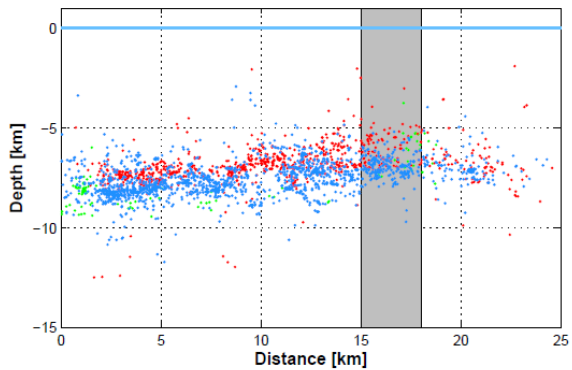




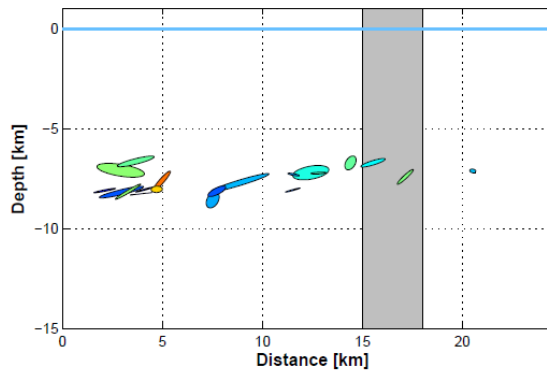
# EXPECTED EARTHQUAKES 0 - 30 YEARS ?



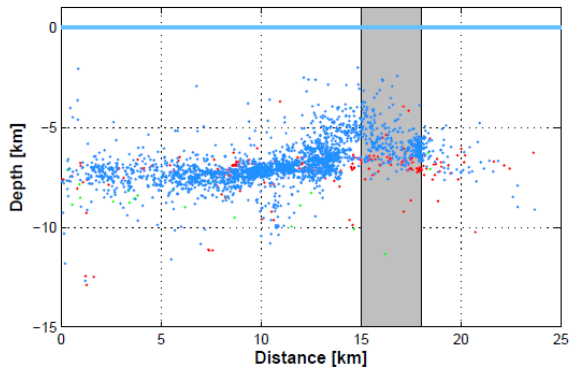
Profil b1-b1'



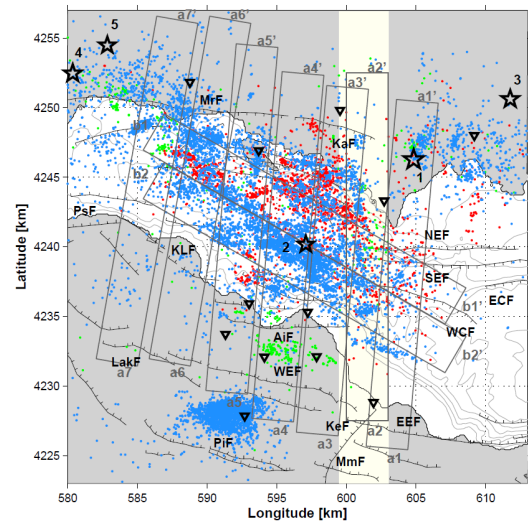
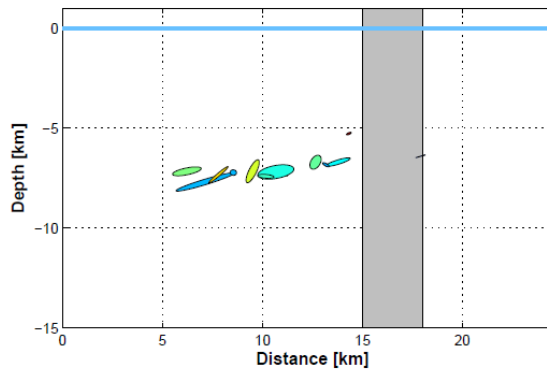
Profil b1-b1'

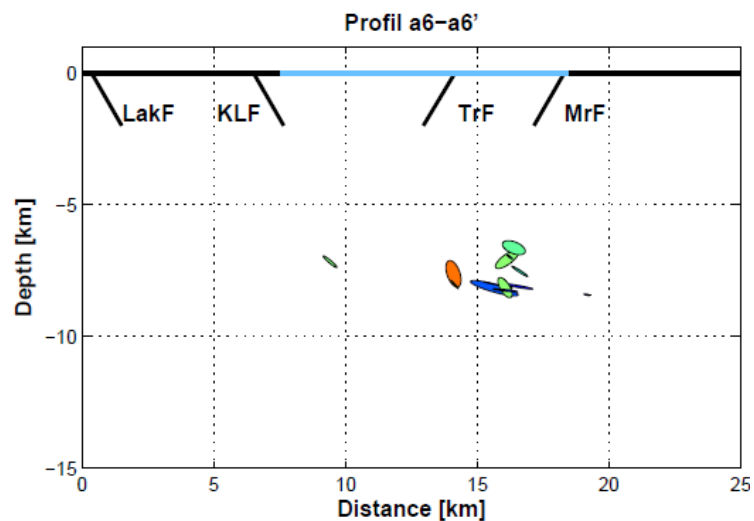
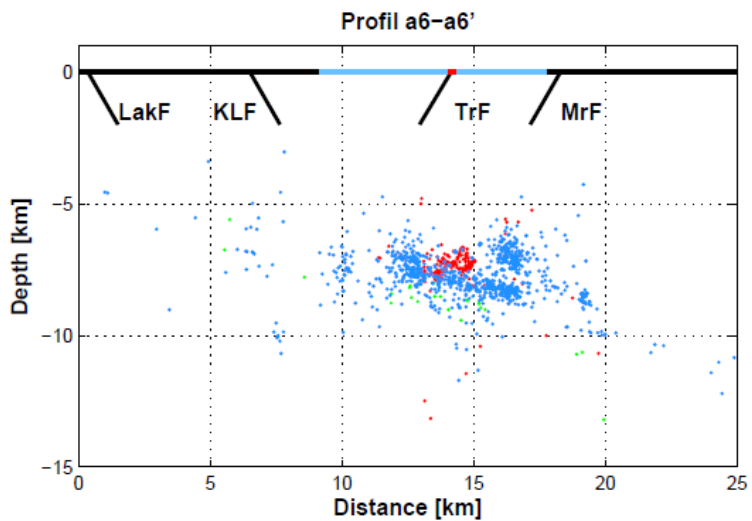
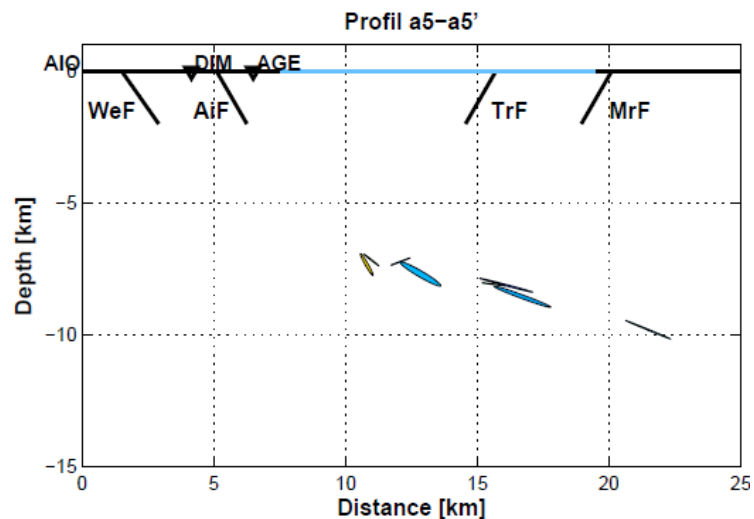
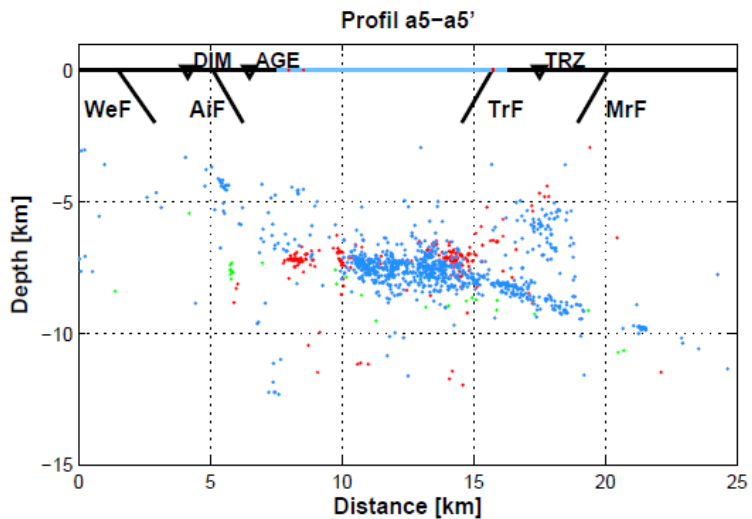


Profil b2-b2'

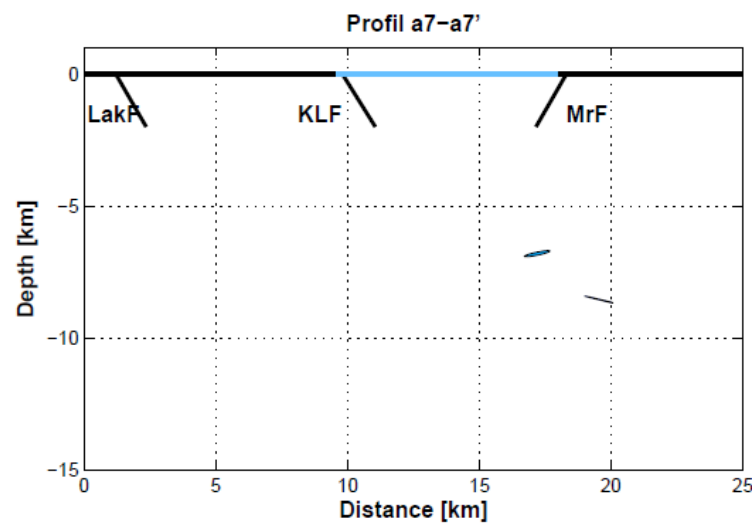
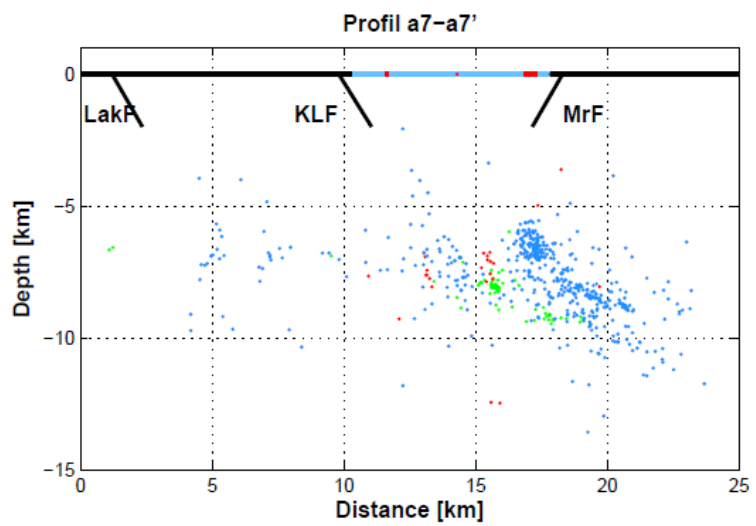


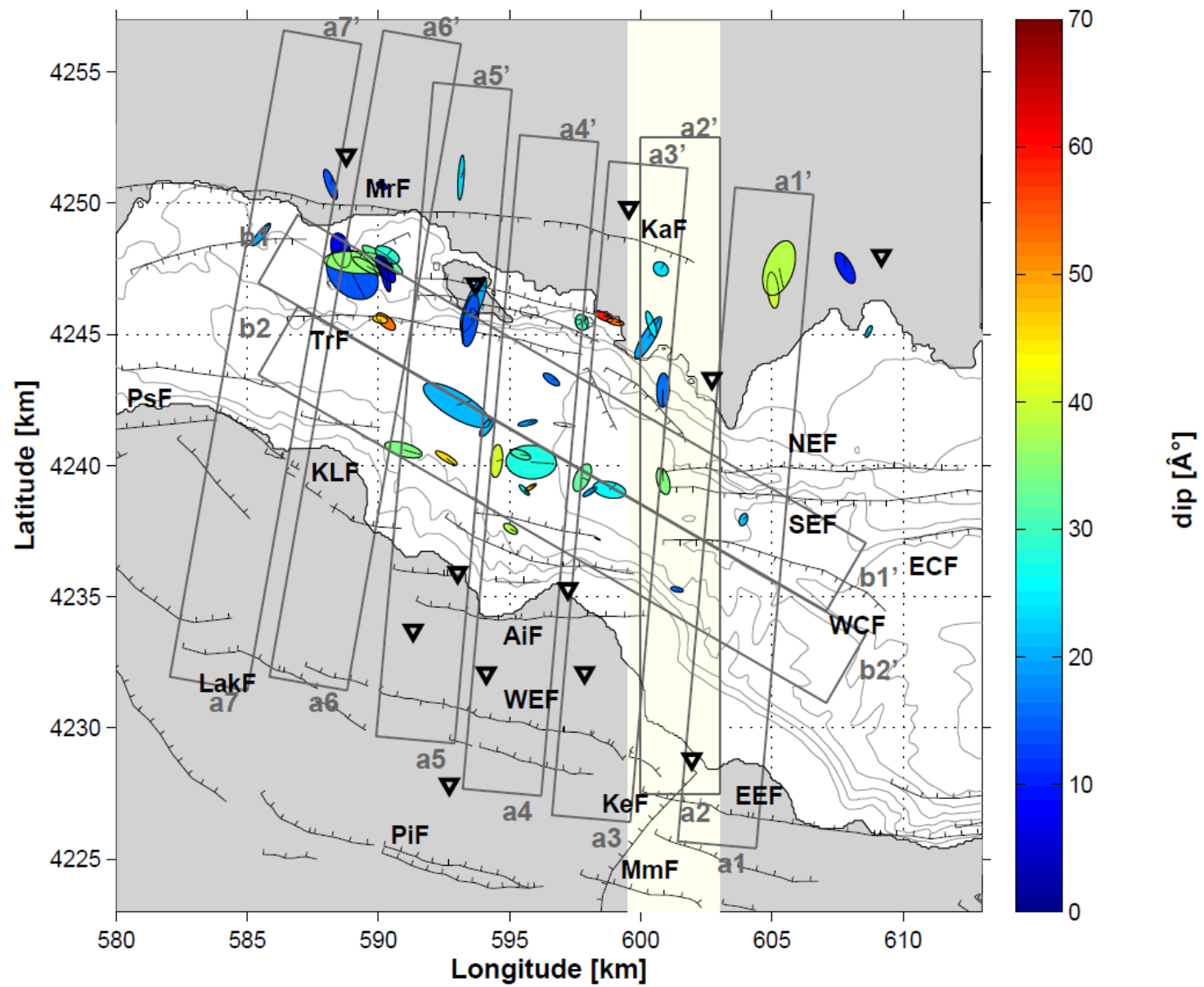
Profil b2-b2'

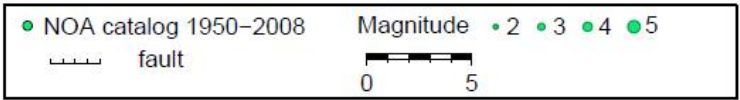
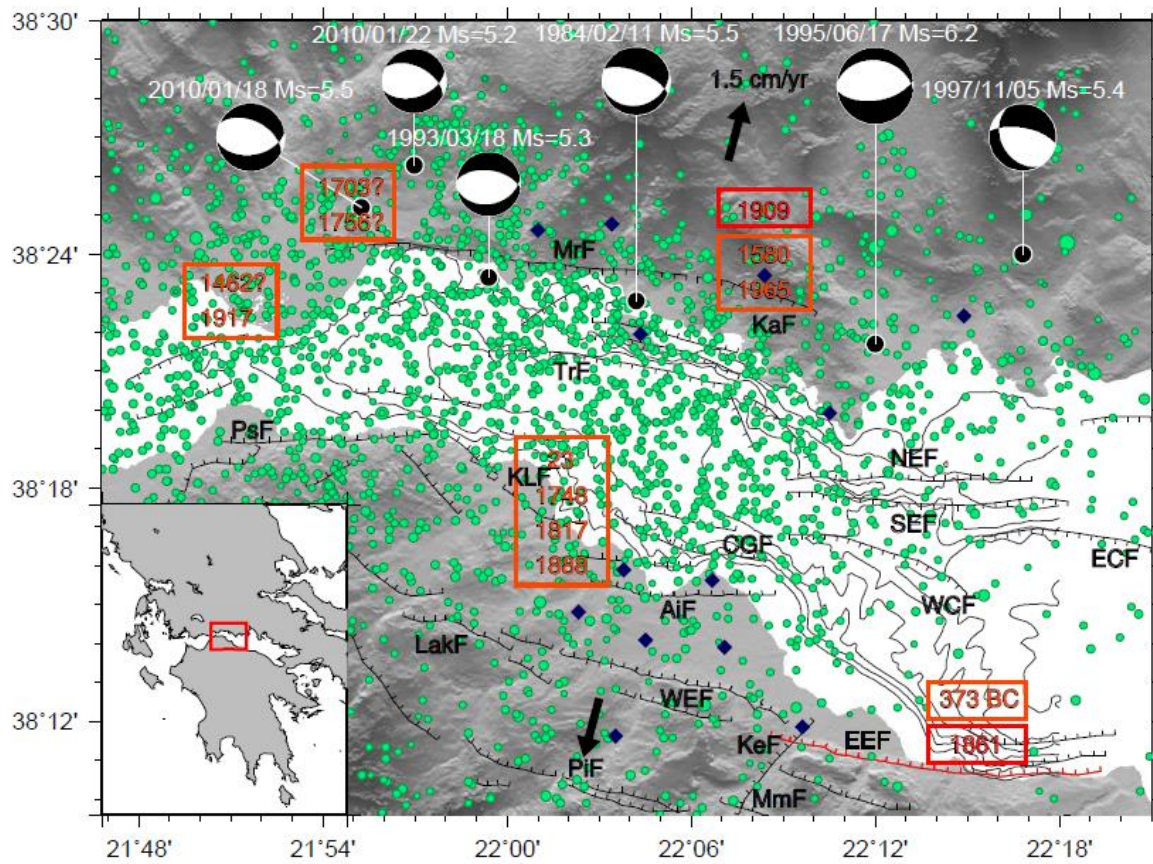


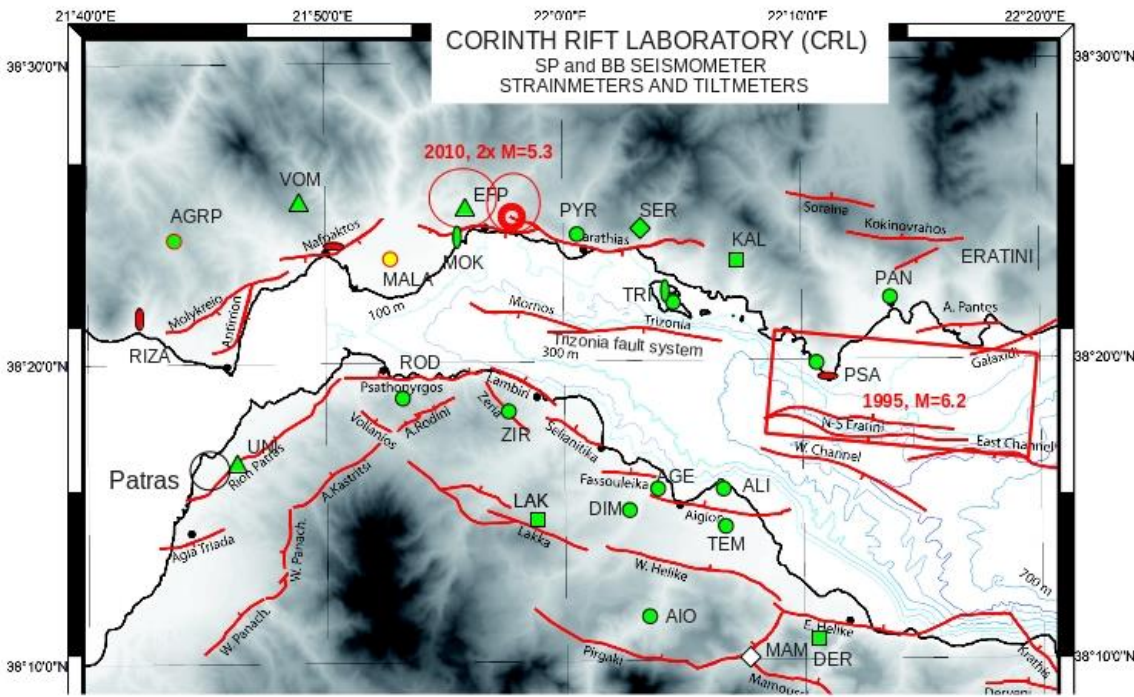










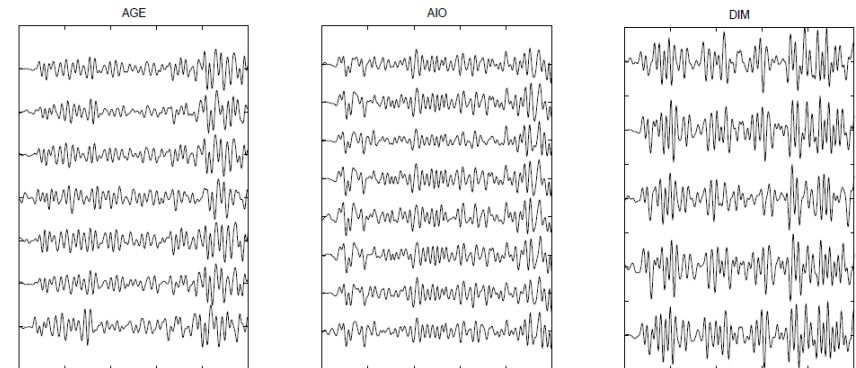


Permanent seismic array since 2000

- Double difference location
- Multiplets: improved locations through cross-correlations

Lambotte et al. 2013

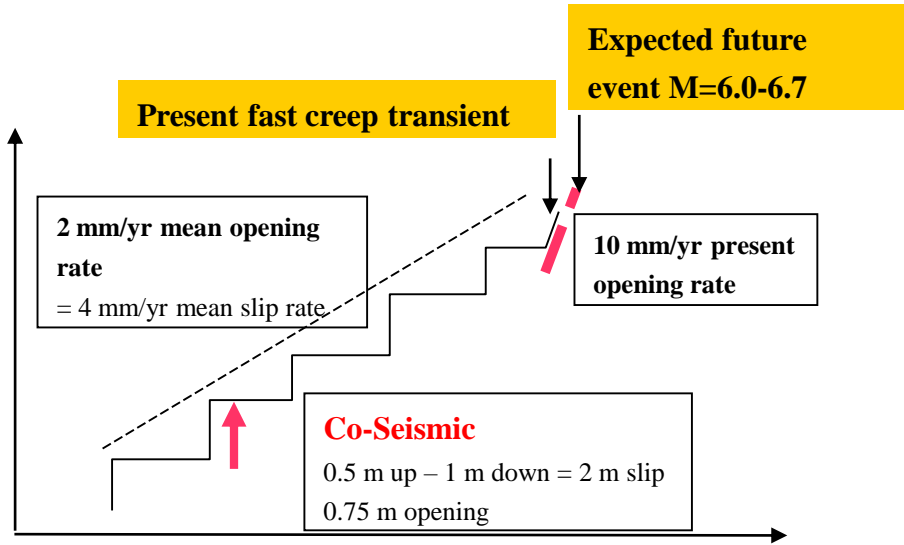
- NKUA
- ◆ NOA
- ▲ PATRAS
- ◆ PRAGUE
- FRANCE CNRS
- in operation
- planned in SISCOR (2012)
- strainmeter
- Planned strainmeter
- Planned tiltmeter



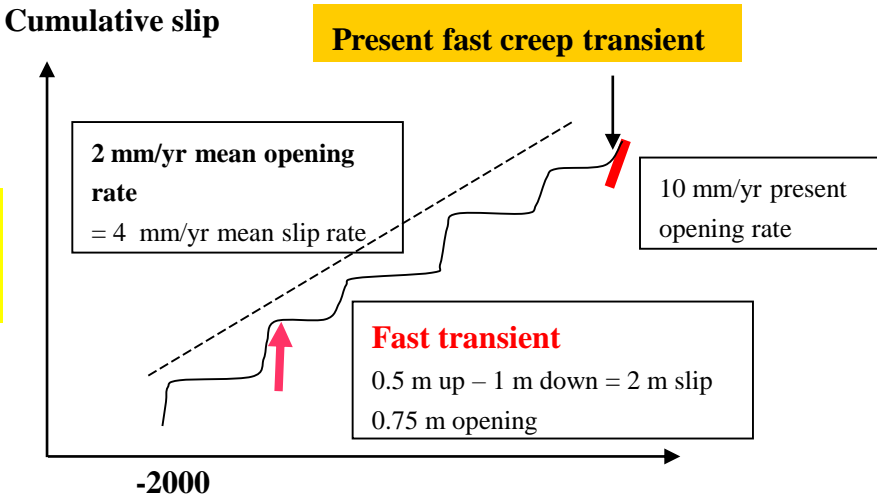
Structure and dynamics of the detachment?

Connection of the normal faults to the detachment?

**Dual model:**  
 Alternating creep and seismic;  
 velocity weakening and  
 strengthening



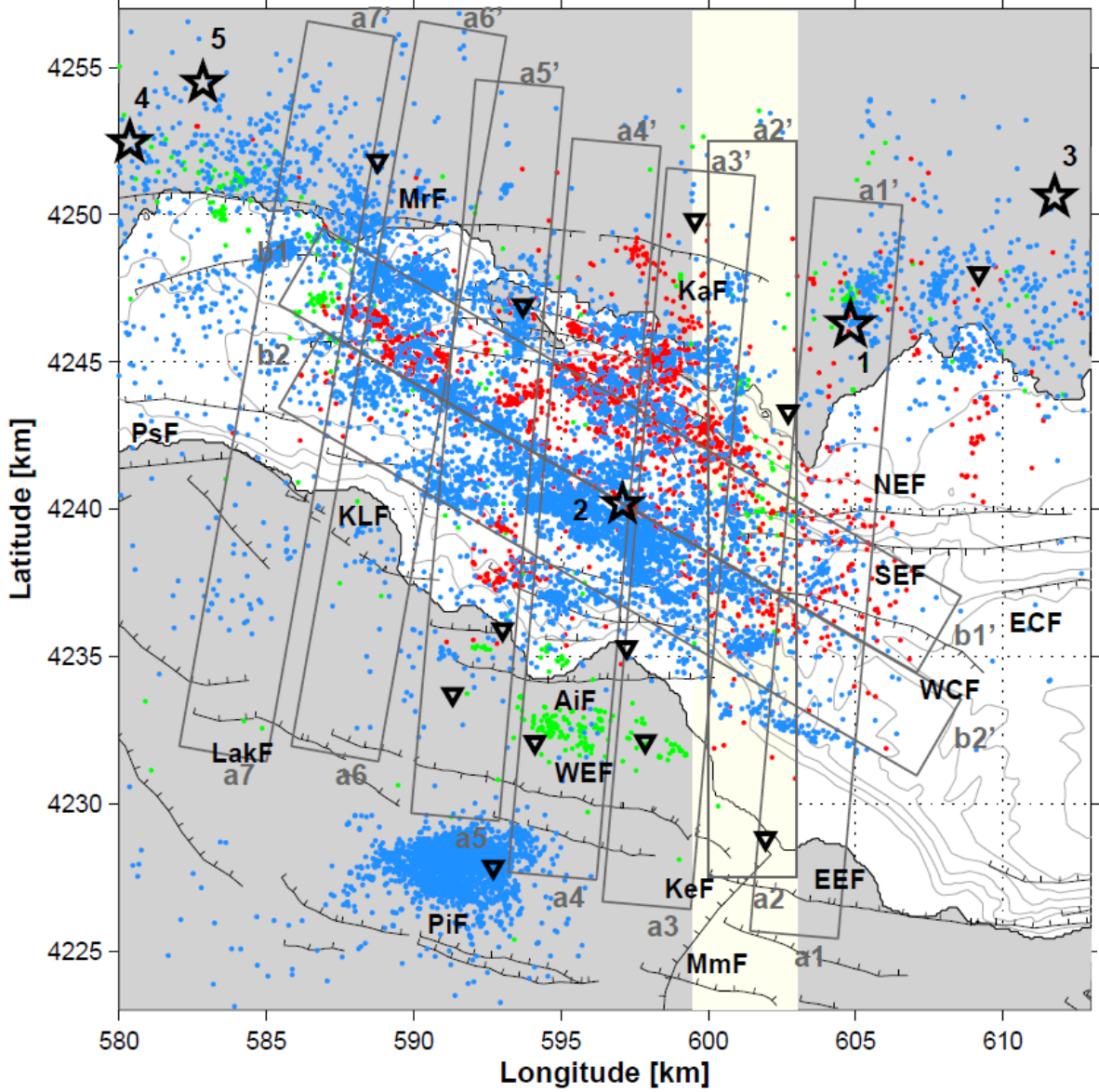
**Pure creep model:**  
 velocity strengthening



**Creep transients on the Psathopyrgos fault at various time scales:  
 hours (strain), days-weeks (seismicity), decades (GPS), centuries (paleo-shorelines)**

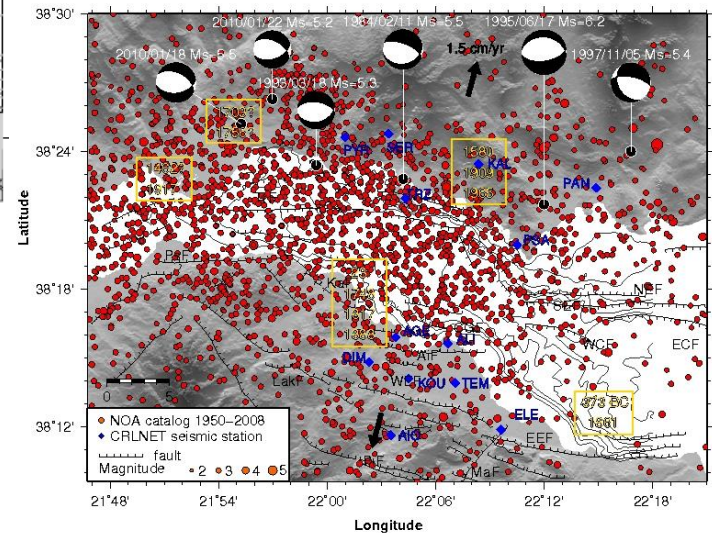
# Double-difference relocation, Lambotte et al. (2013)

- NW trend of micro-seismicity: at the root of the recent fault system
- strong segmentation
- isolated clusters



Blue: 2001-2007  
 Red: Aftershocks 1995  
 Green: July 1991

# NOA locations since 1975



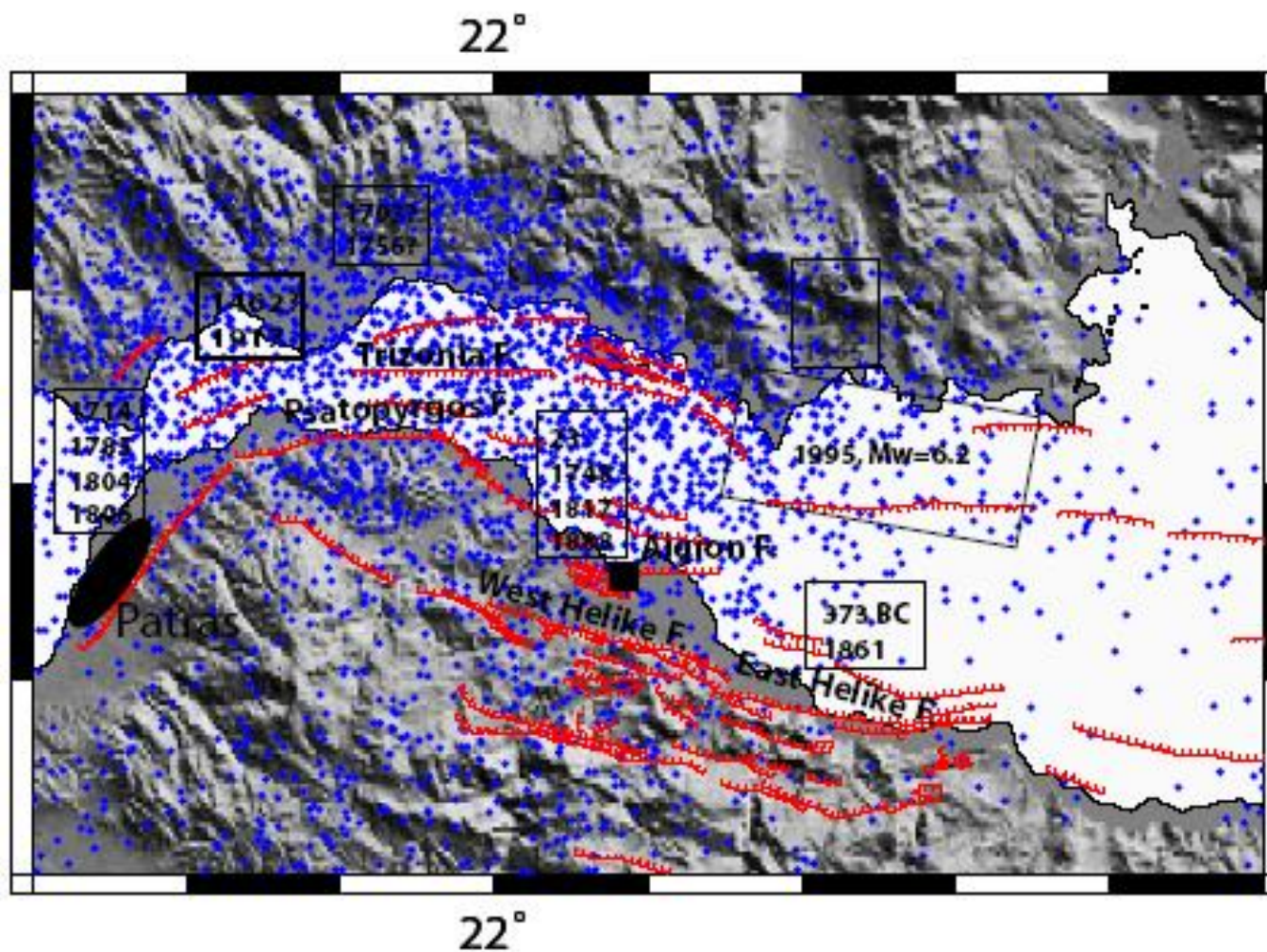
Legend:  
 ● NOA catalog 1950-2008  
 ◆ CRLNET seismic station  
 --- fault  
 Magnitude: ● 2 ● 3 ● 4 ● 5



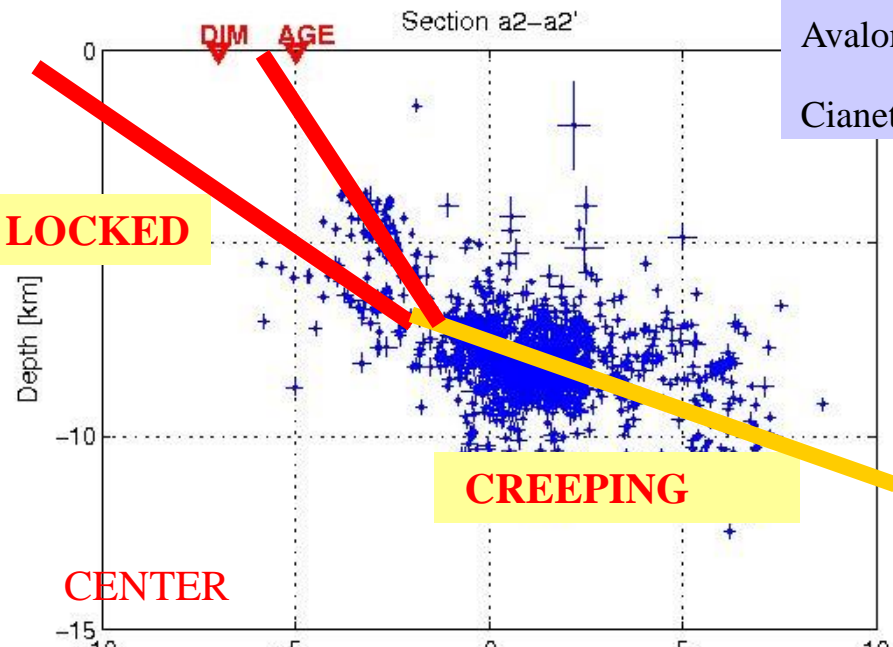
HELIKE FAULT



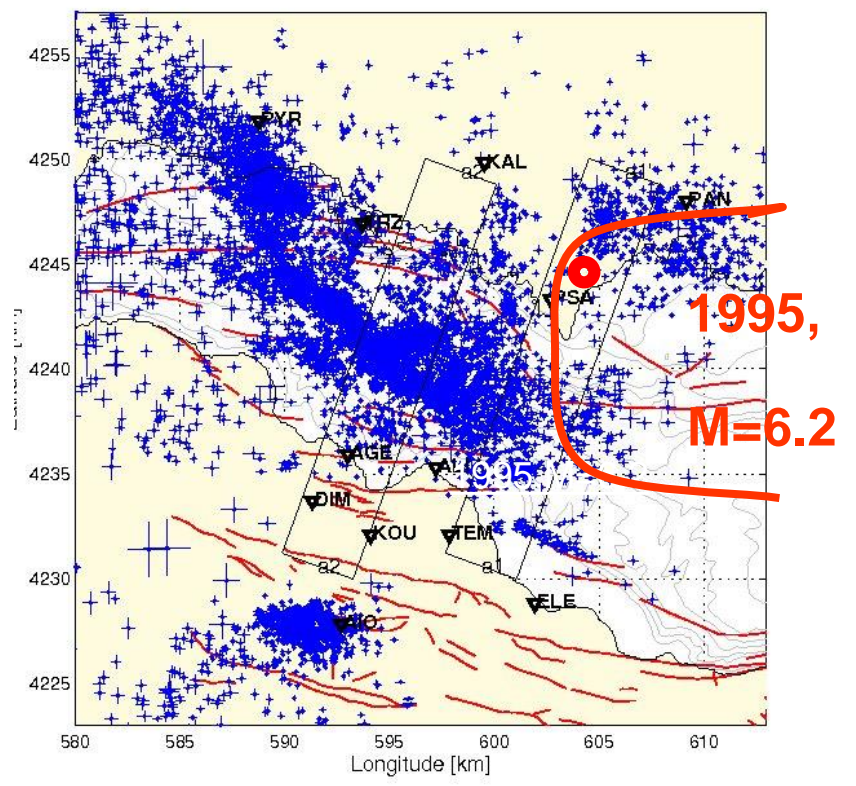
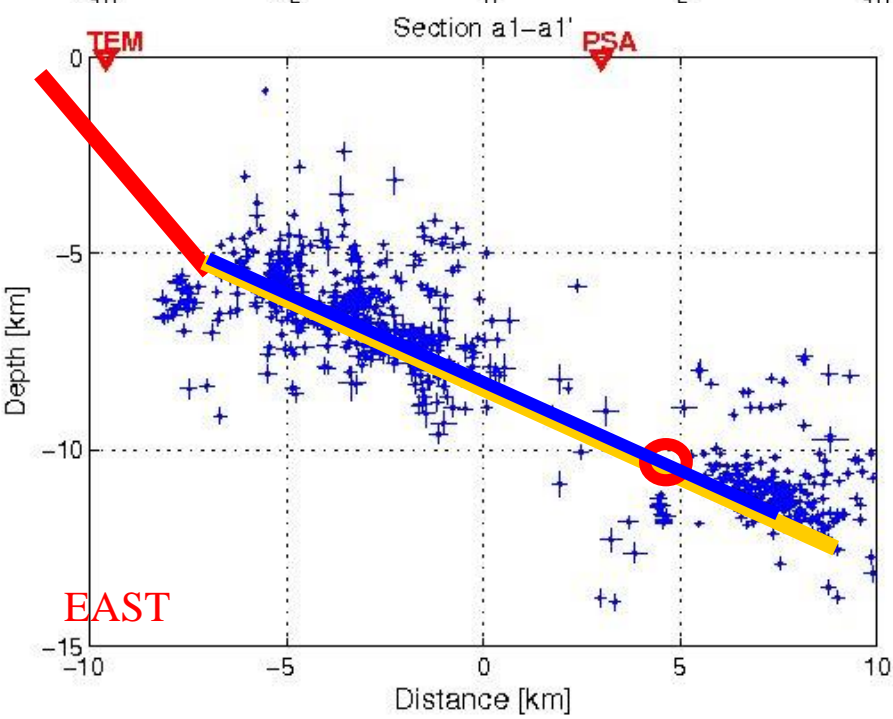
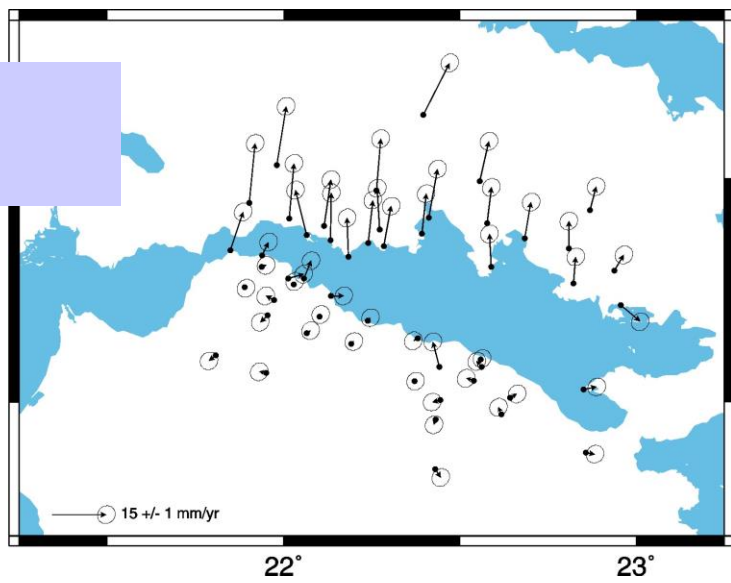








Avalone et al., 2004  
 Cianetti et al. 2008

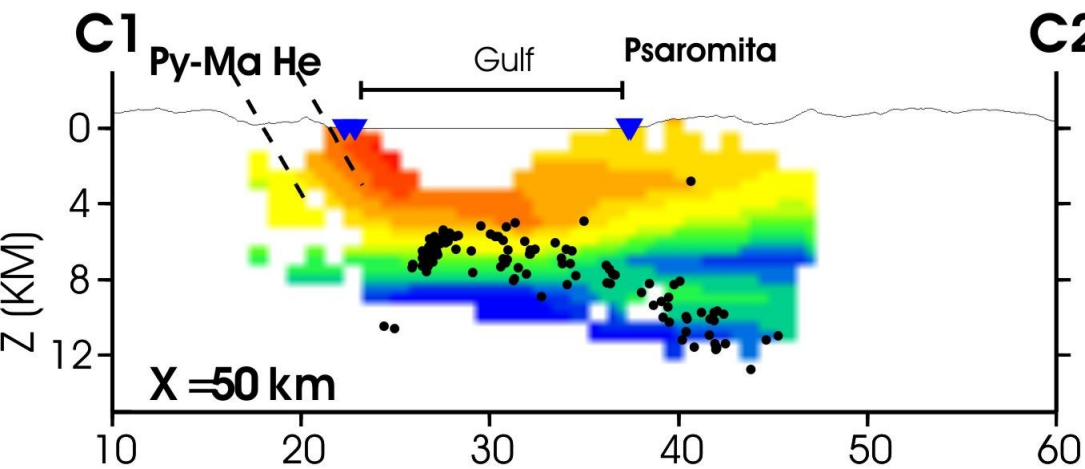
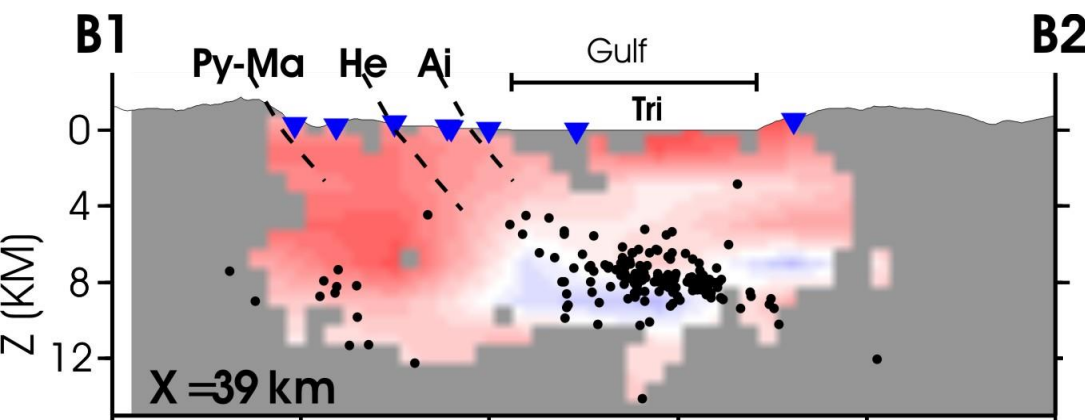
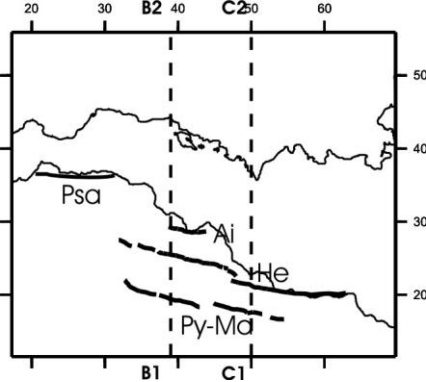


**SEISMICITY 2001-2007**

# Corinth Tomography: Rifting and fluids

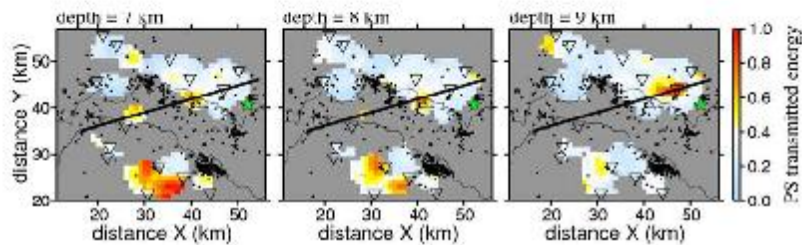
Latorre, 2004

Gautier et al. 2006

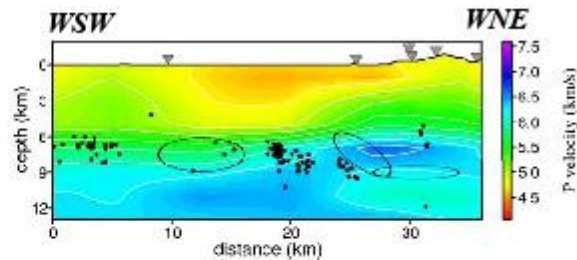
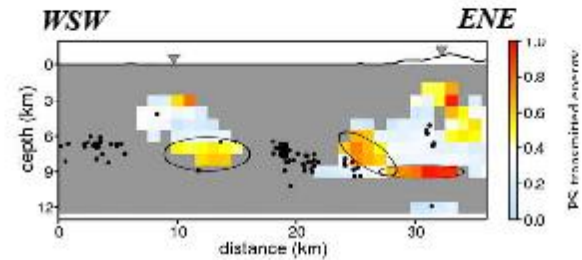


# RESULTS

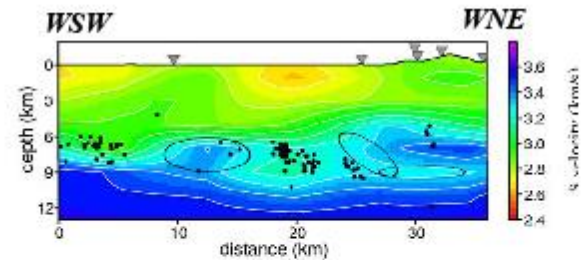
Map view of the migrated model a 7km, 8km and 9 km depth



Vertical cross-section oriented 70° N



P velocity model



S velocity model

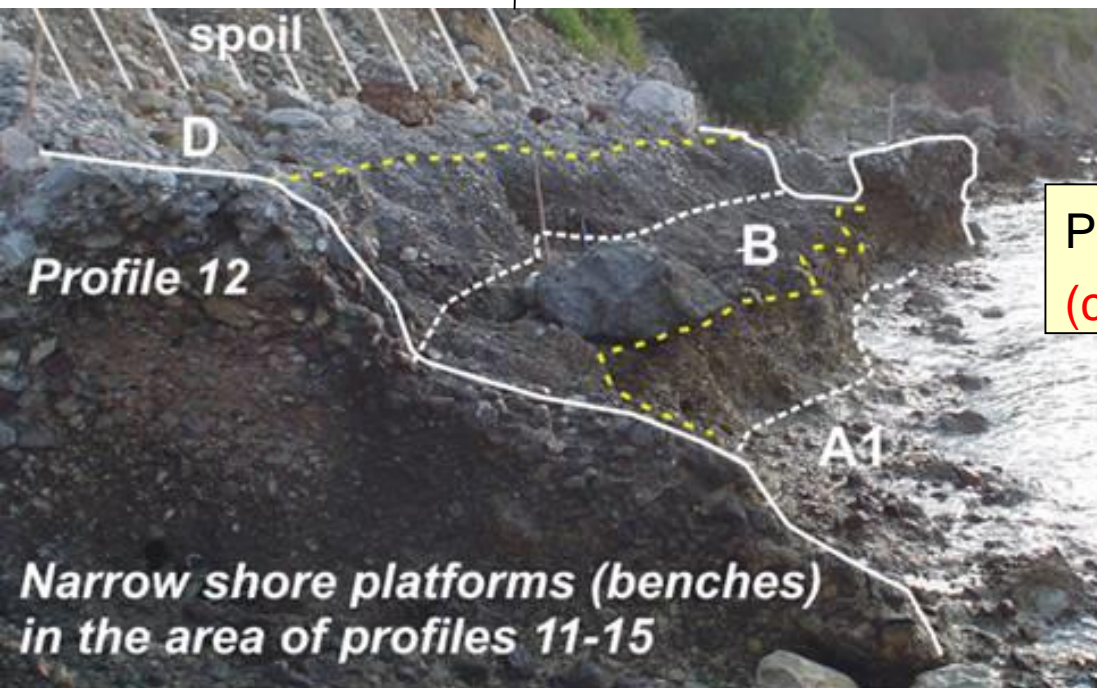
Latorre D., Virieux J., Monfret T., in preparation

Nancy 2-3 Juillet 2007

Le rift de Corinthe

9

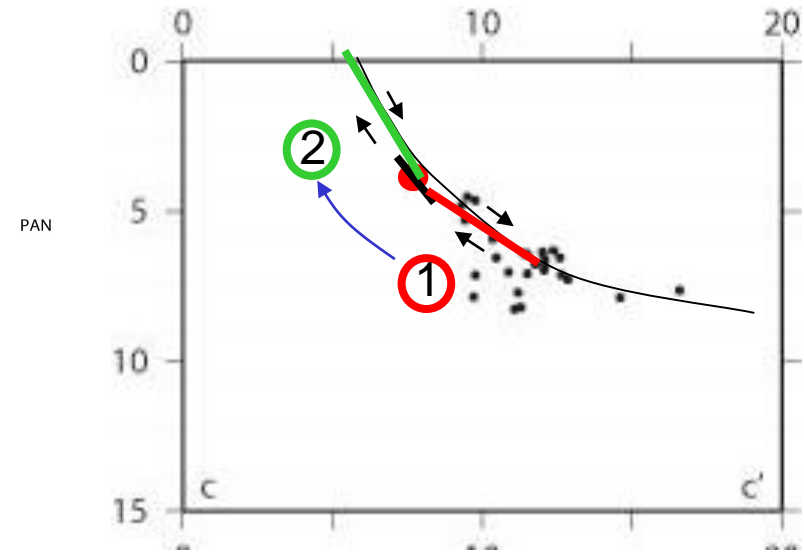
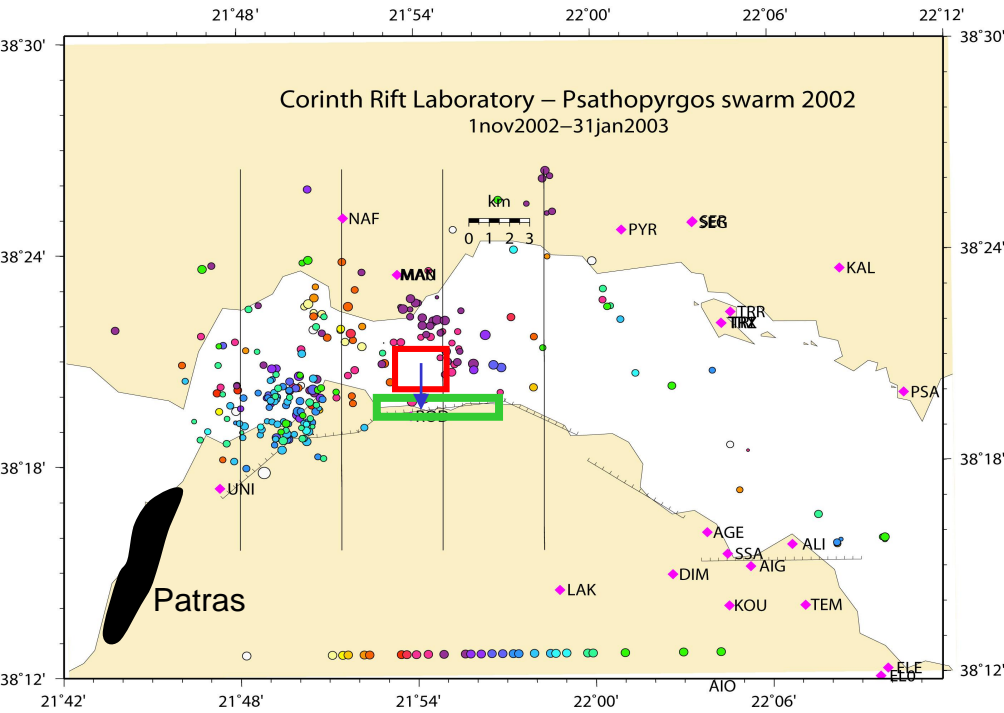
# Psathopyrgos fault



Paleoshoreline: 5 rapid uplift episodes (coseismic?) in 2000 years

Pantosti et al. 2007

# Transient creep on the Psathopyrgos fault



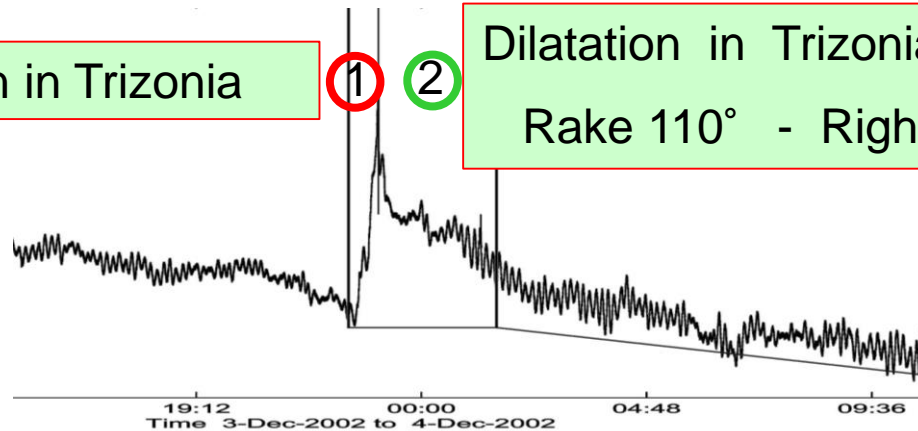
**Aseismic slip -  $\Delta u=10$  cm -  $M_w=5$**

Compression in Trizonia

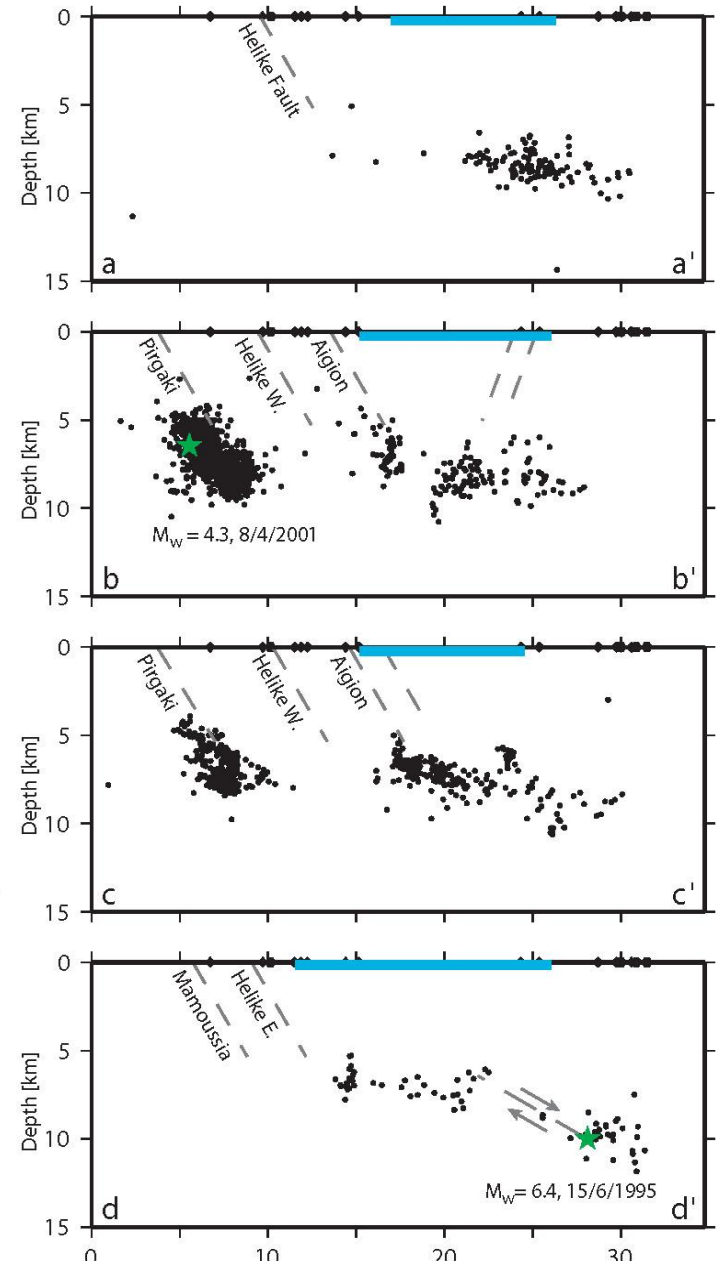
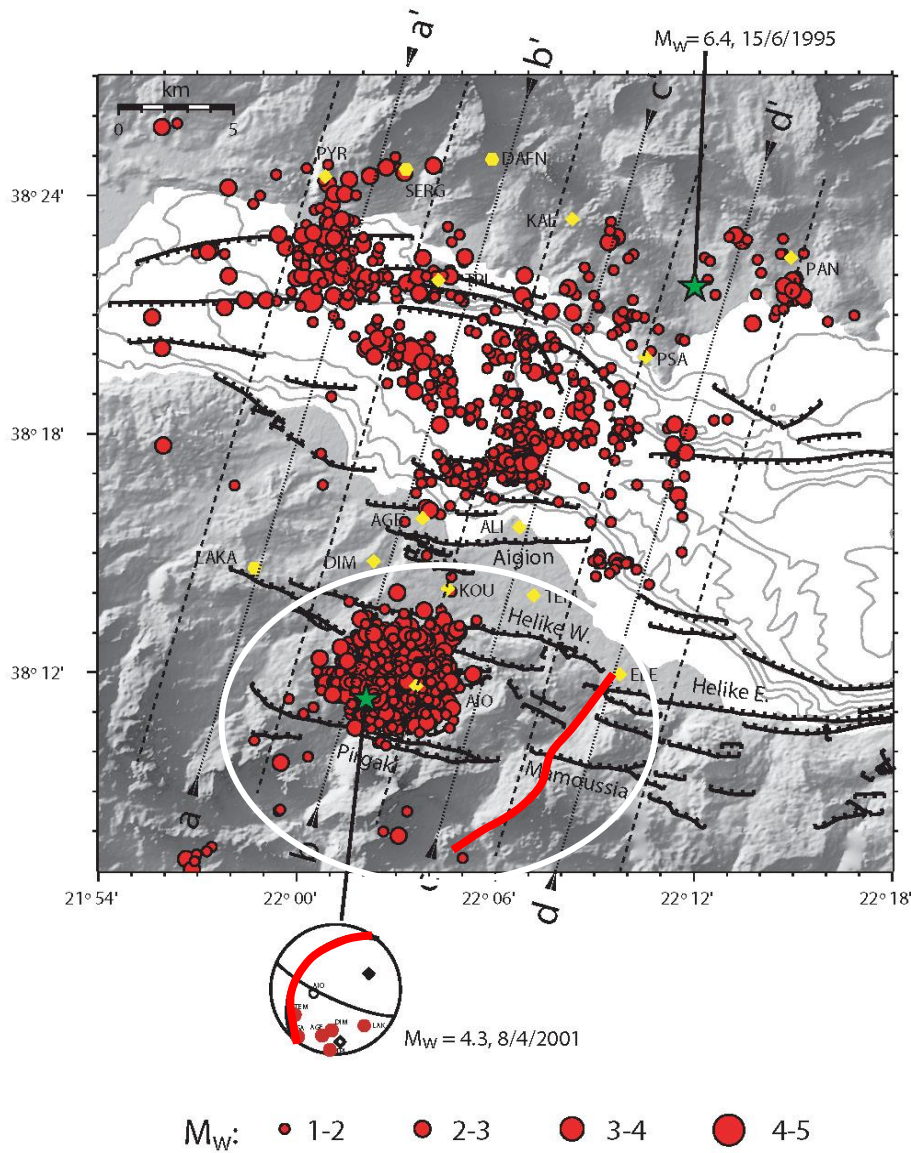
① ②

Dilatation in Trizonia

Rake  $110^\circ$  - Right-lateral

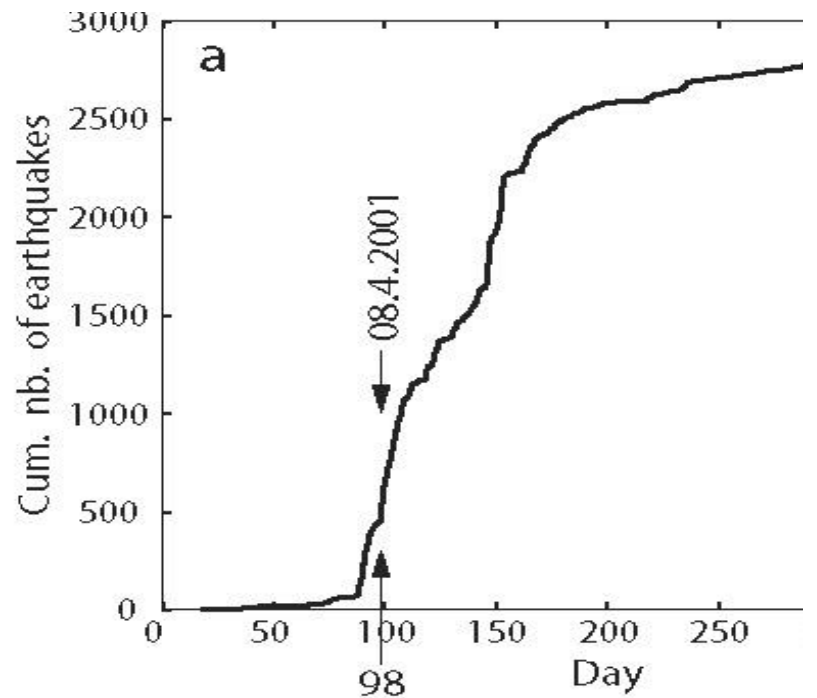
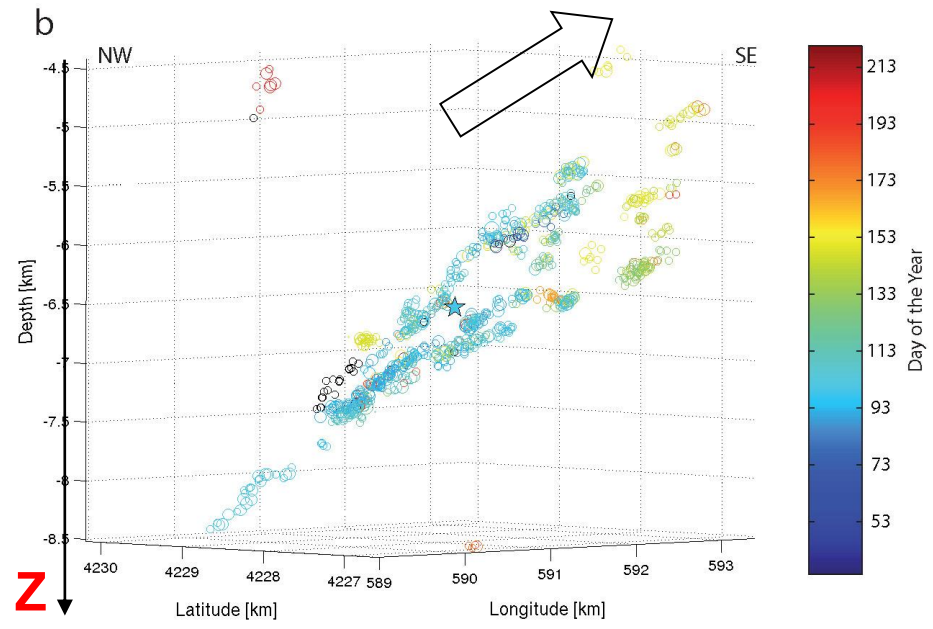
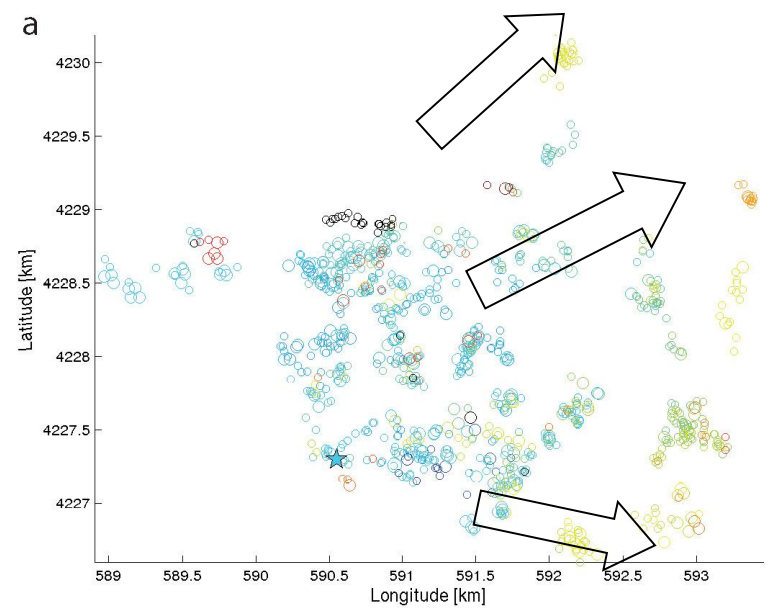


2001 seismic swarm - 3 months -  $M_{max}=4.2$   
 ancient fault - dip  $45^\circ$  NW - reactivated





# 2001 seismic swarm: relocated multiplets



**2 time scales of transient:**

**Week-month:**

**Upward migration :**

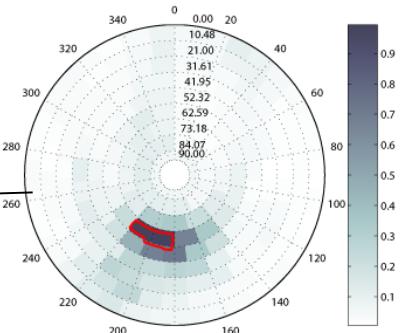
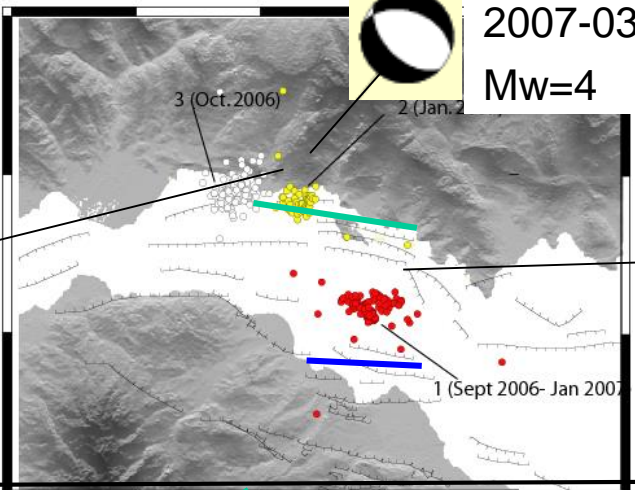
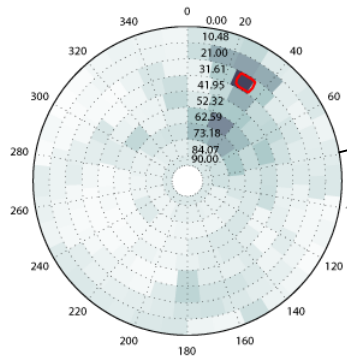
**pore pressure pulse?**

**Minute-day:**

**clustering Omori type**



2007-03-19  
Mw=4

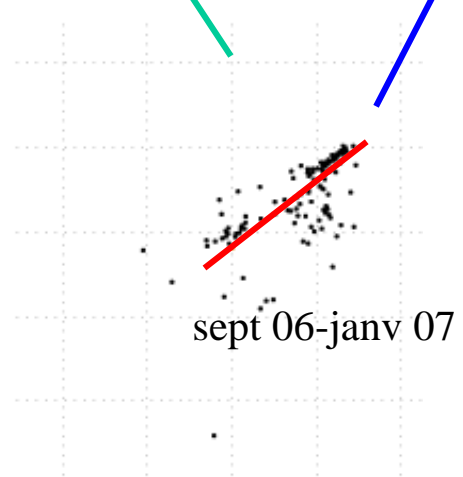
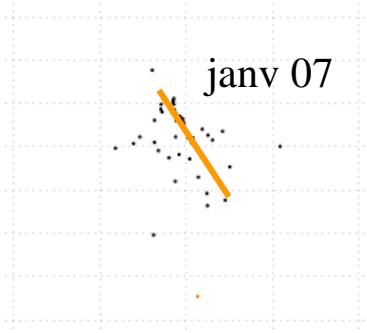


NNE ← 0 → SSW

Trizonia  
fault

Aigion  
fault

0  
5  
10  
km



### -Pendages failles principales vers le nord:

- Helike: inconnu
- Aigion: 60° jusqu'à 6 km
- Kamarai-Lambiri: 60° probable jusqu'à 6 km
- Psathopyrgos (2002 – Latorre – Rigo)

### Structure profonde:

#### - centre et est:

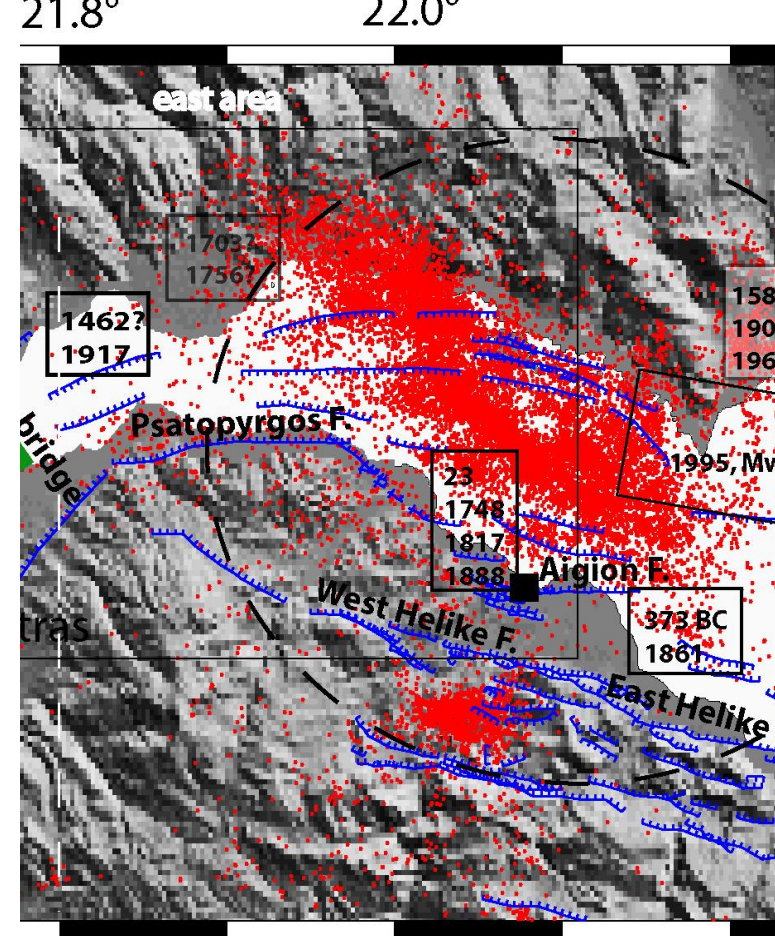
- épaisseur < 2km, liée à une zone de faible vitesse
- Pendage nord, entre 6 et 8 km coté sud, descendant jusqu'à 12-15 km coté nord
- approfondissement vers le NW, haut topo longitude Aigion
- petites failles pendage nord, 10°-30°, quelques petites structures décrochantes
- Petite faille oblique pendage NE pres de l'hypocentre du seisme de 1995

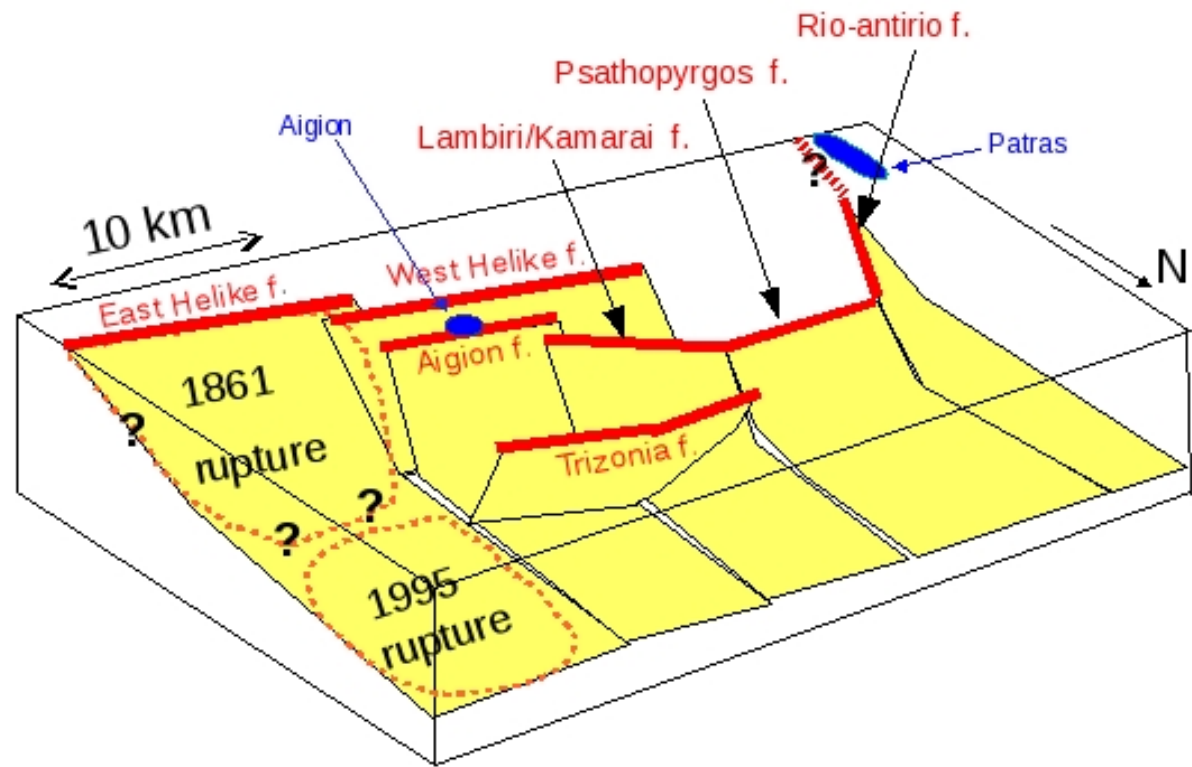
### -Failles obliques

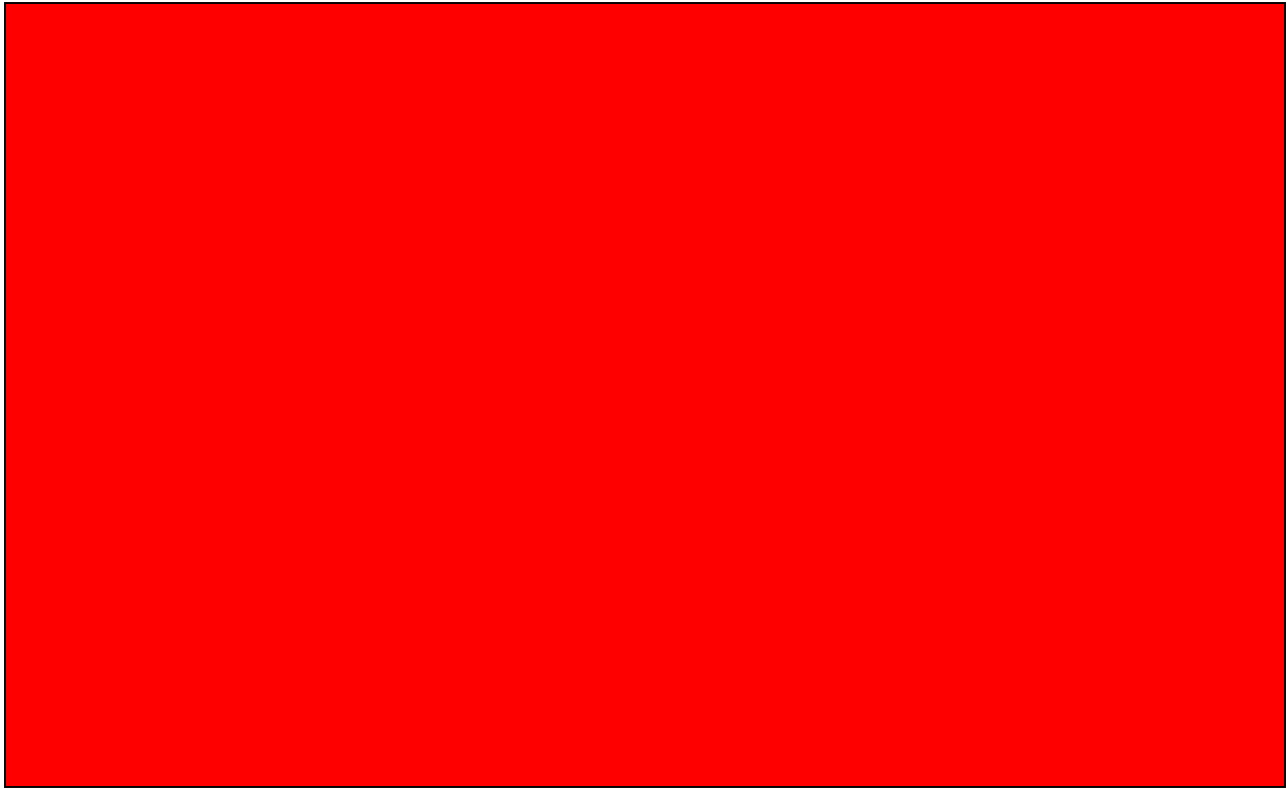
- faille de 2001: décrochante, 5-8 km, pendage NW, peut être liée à une structure majeure

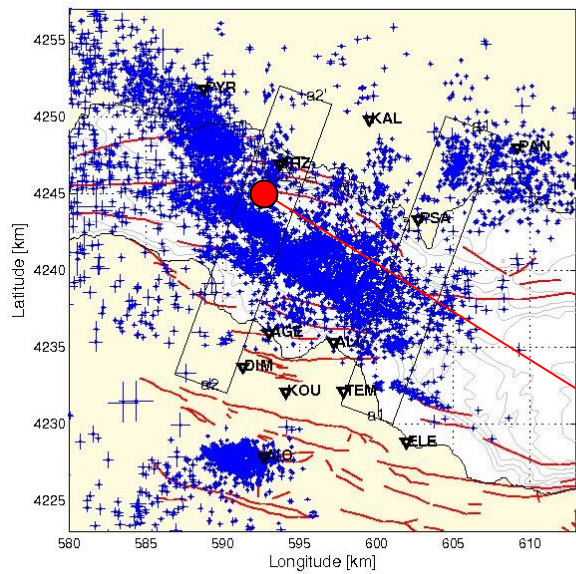
### -Failles antithétiques:

- faille de Trizonia: activité diffuse vers 5 km, pas d'activité de multiplet
- séisme de 2007: faille normale sous cote nord, plus au nord que Trizonia



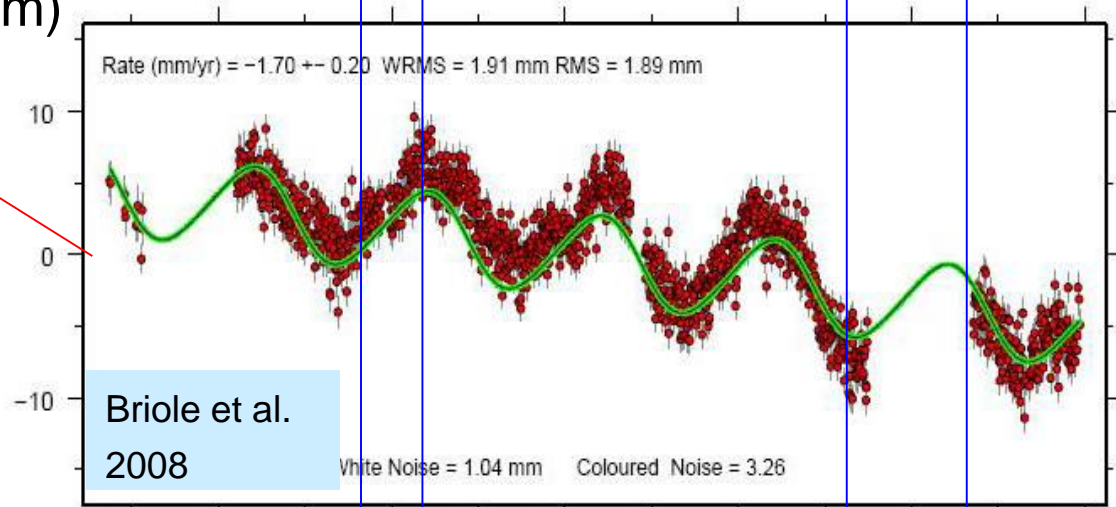






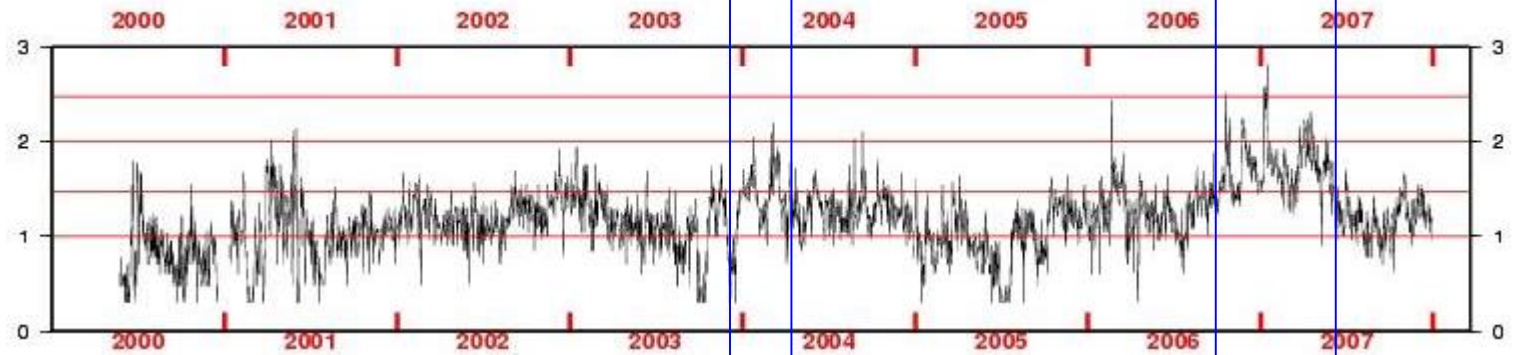
Displacement (N)  
(mm)

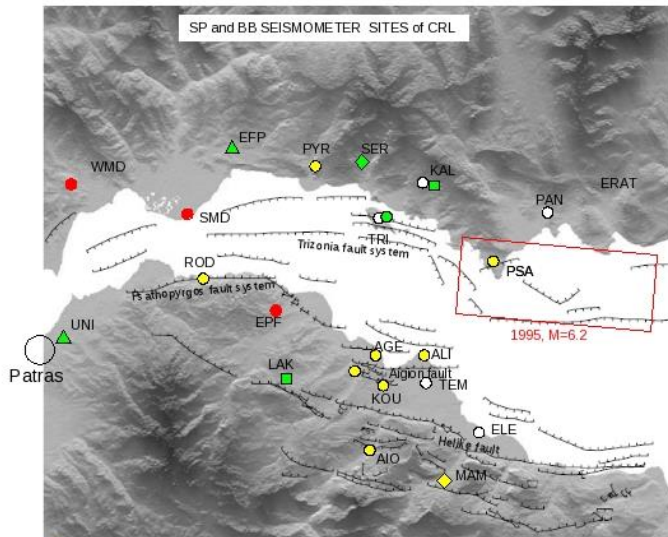
cGPS



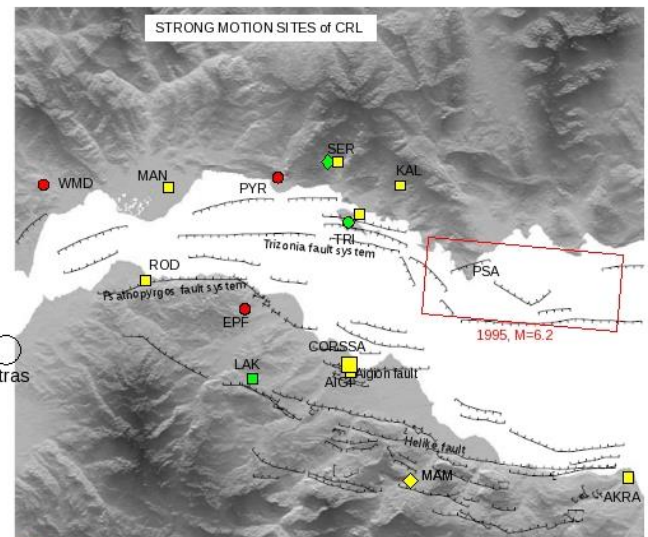
Log(seismicity rate)

/day

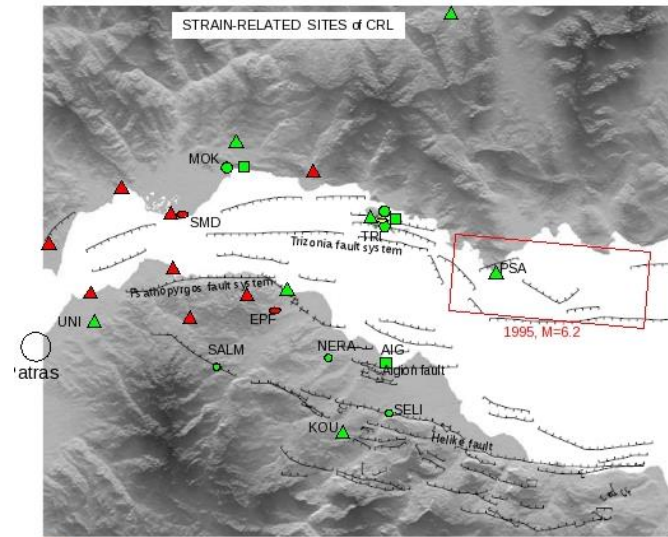




- NKUA
- ⬠ NOA
- ▲ PATRAS
- ◆ PRAGUE
- FRANCE CNRS
- ▲ in operation, to be maintained
- ▲ in operation, to be upgraded
- △ to be suppressed
- seismometer planned in SISCOR

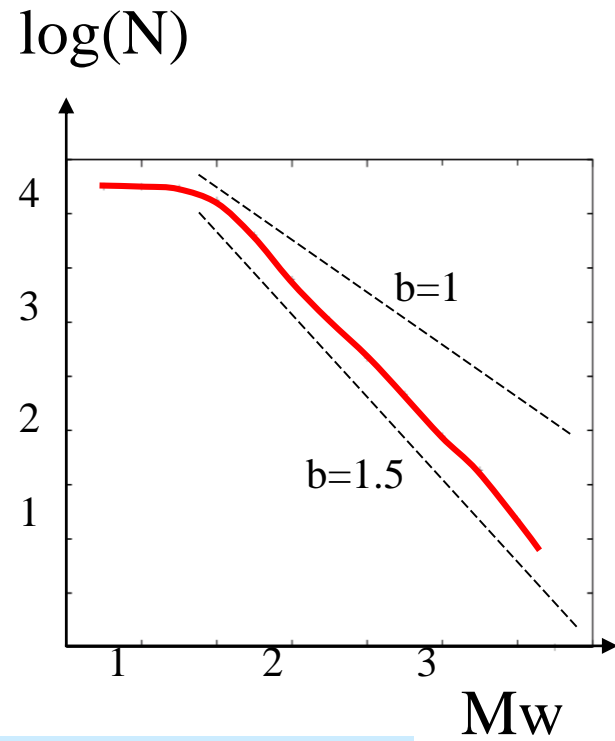


- NKUA
- ⬠ NOA
- ▲ PATRAS
- ◆ PRAGUE
- FRANCE
- ▲ in operation, to be maintained
- ▲ in operation, to be upgraded
- accelerometer planned in SISCOR

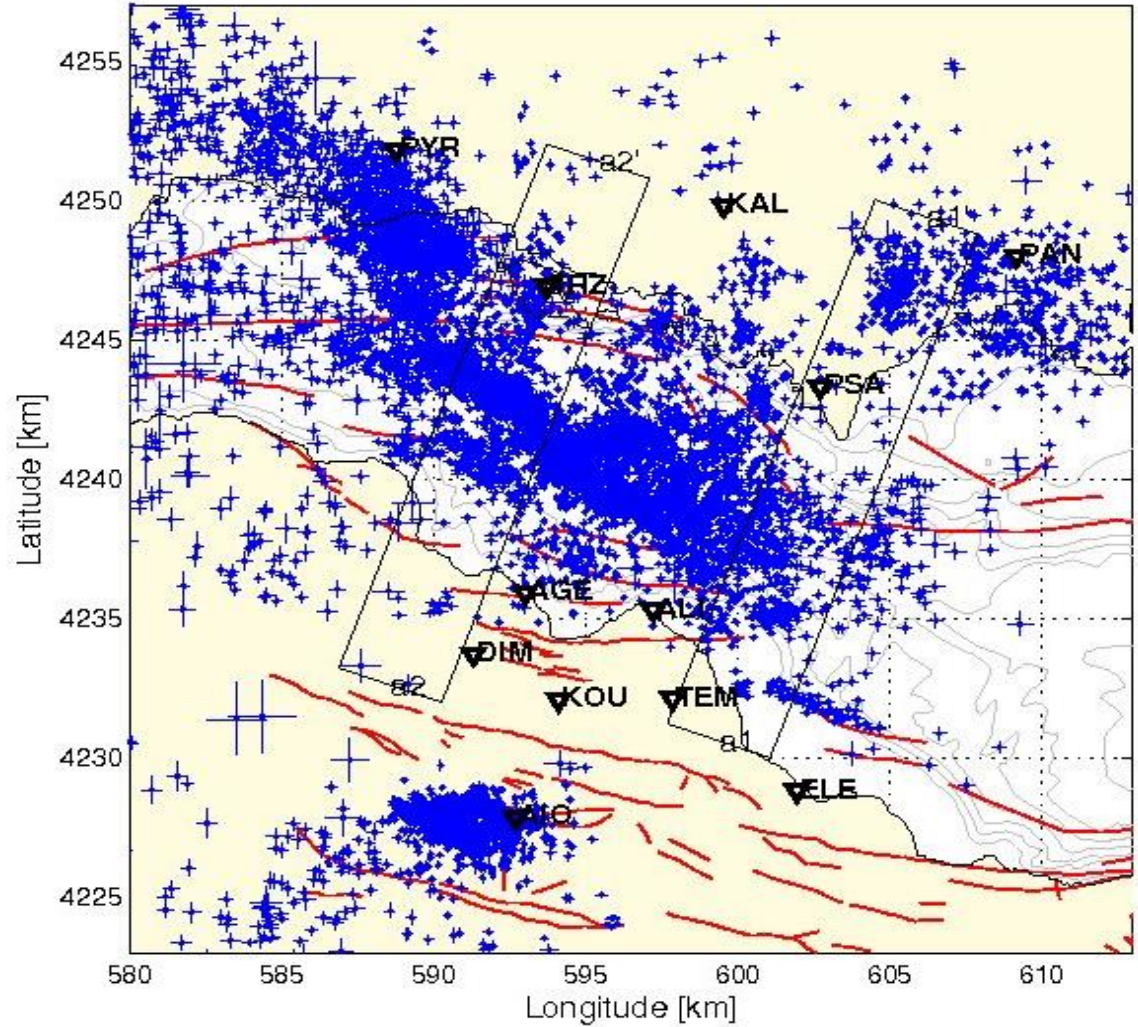


- ▲ GPS (ENS)
- Borehole Strainmeter (IPGP)
- Borehole tiltmeter (IPGP)
- Borehole pore pressure/water level (IPGP/UM)
- Long base tiltmeter (IPGP)
- Tide-gage (IPGP)
- ⬠ Meteo station (IPGP)
- ▲ in operation, to be maintained
- ▲ in operation, to be upgraded
- borehole tiltmeter planned in SISCOR
- ▲ continuous GPS planned in SISCOR

# Seismicity 2000-2007



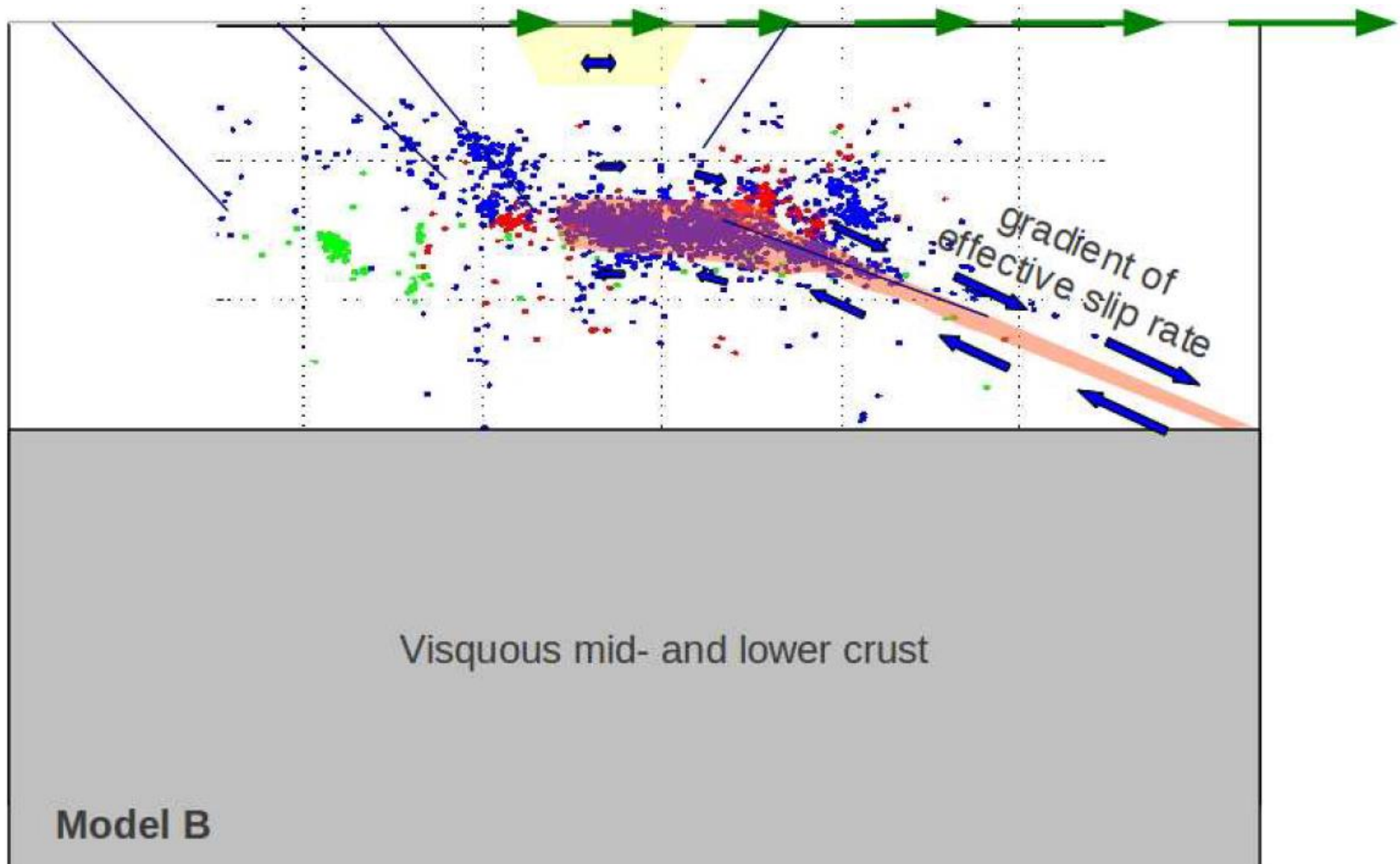
Pacchiani, 2006



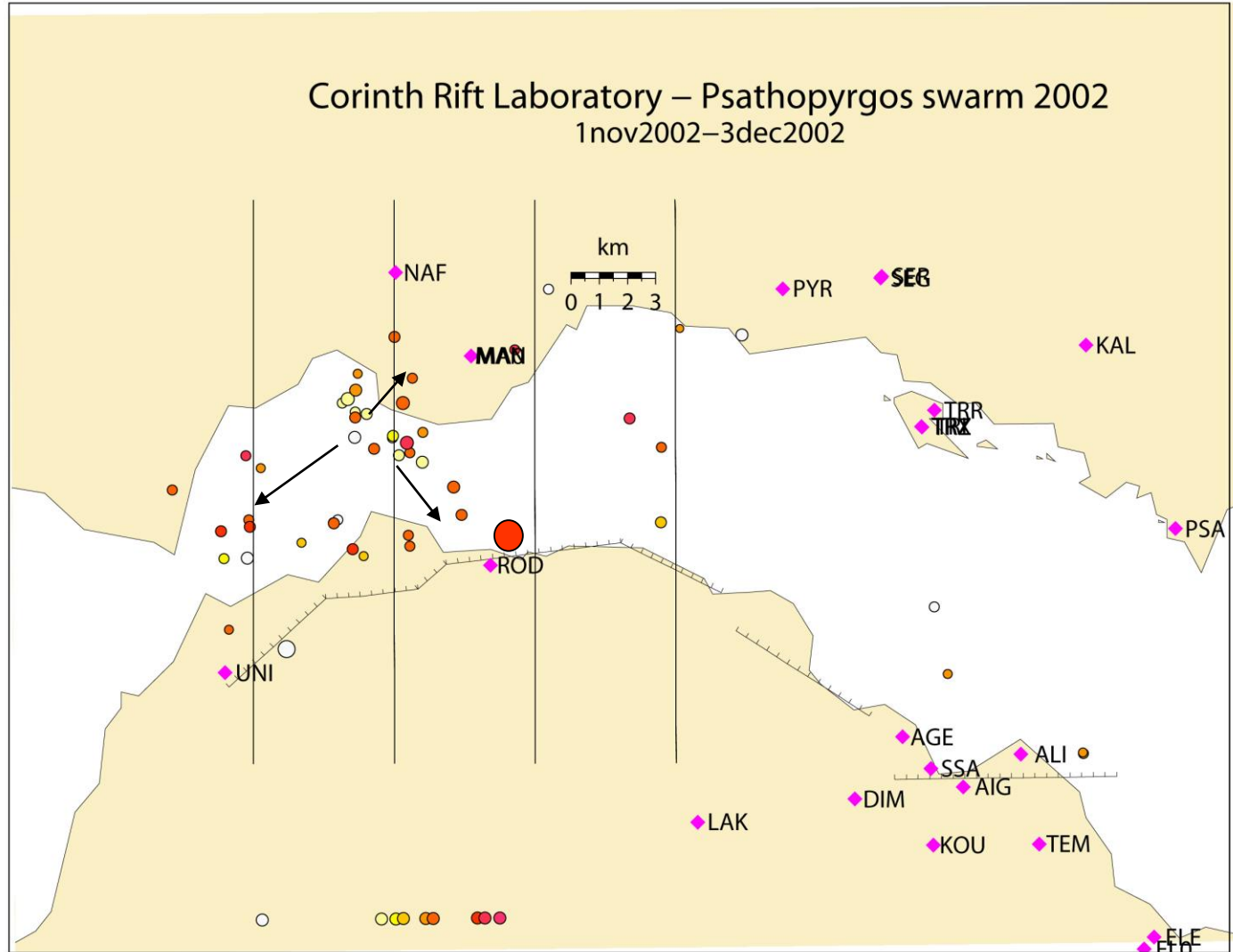
Lambotte, 2008



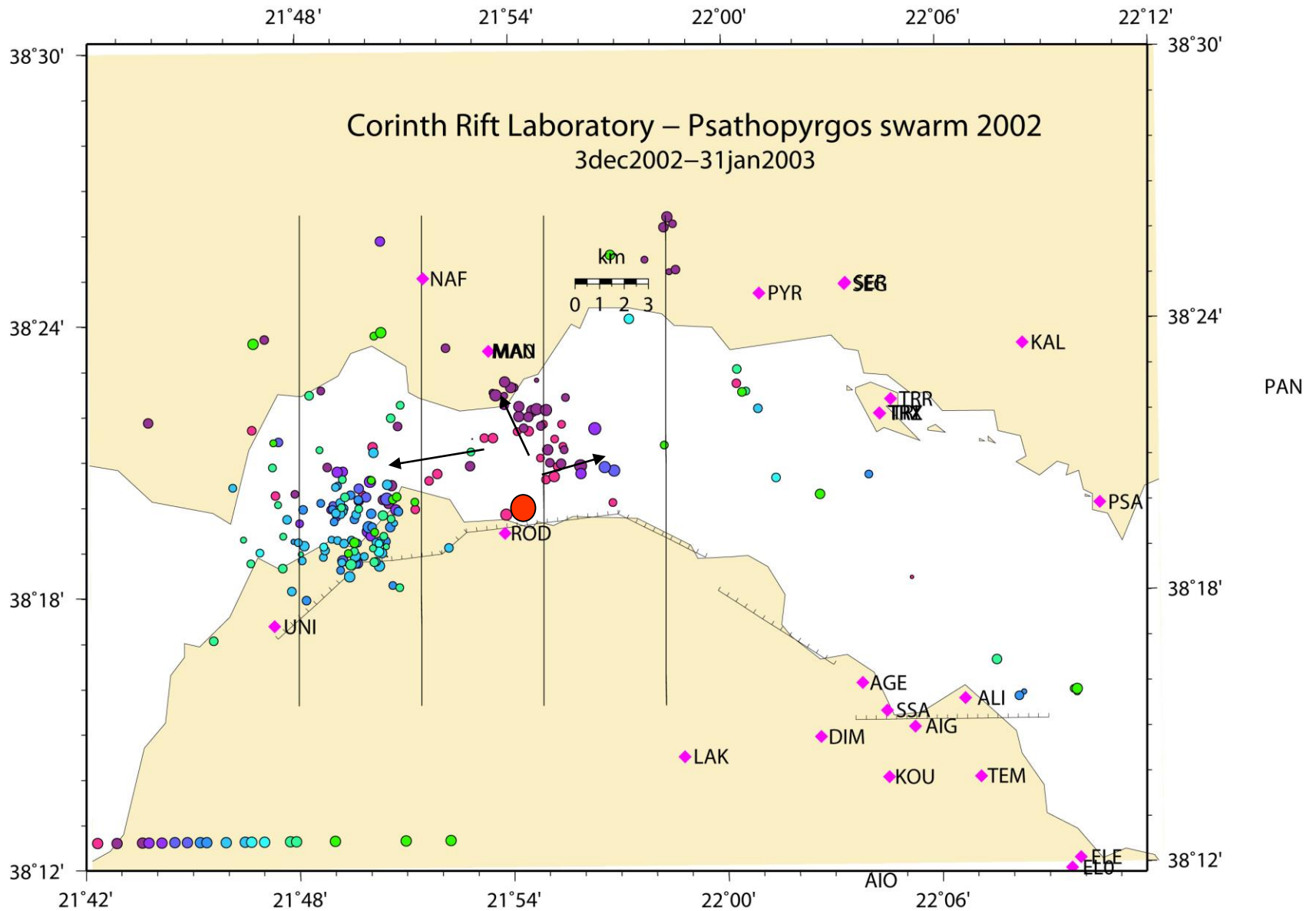
Detachment model with constant opening rate (remote) and constant friction

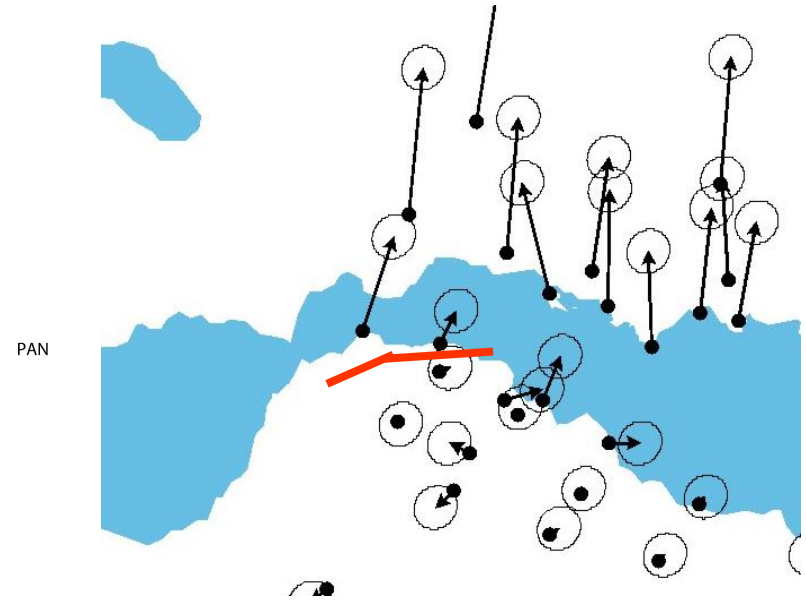
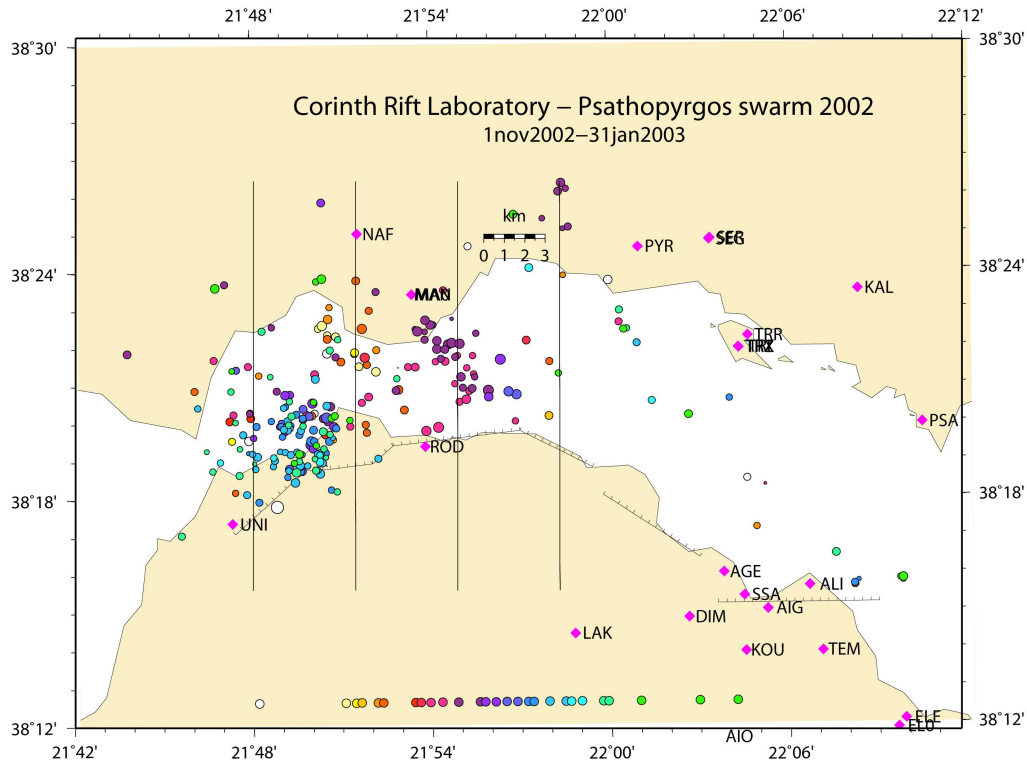


# Before 3 december earthquake

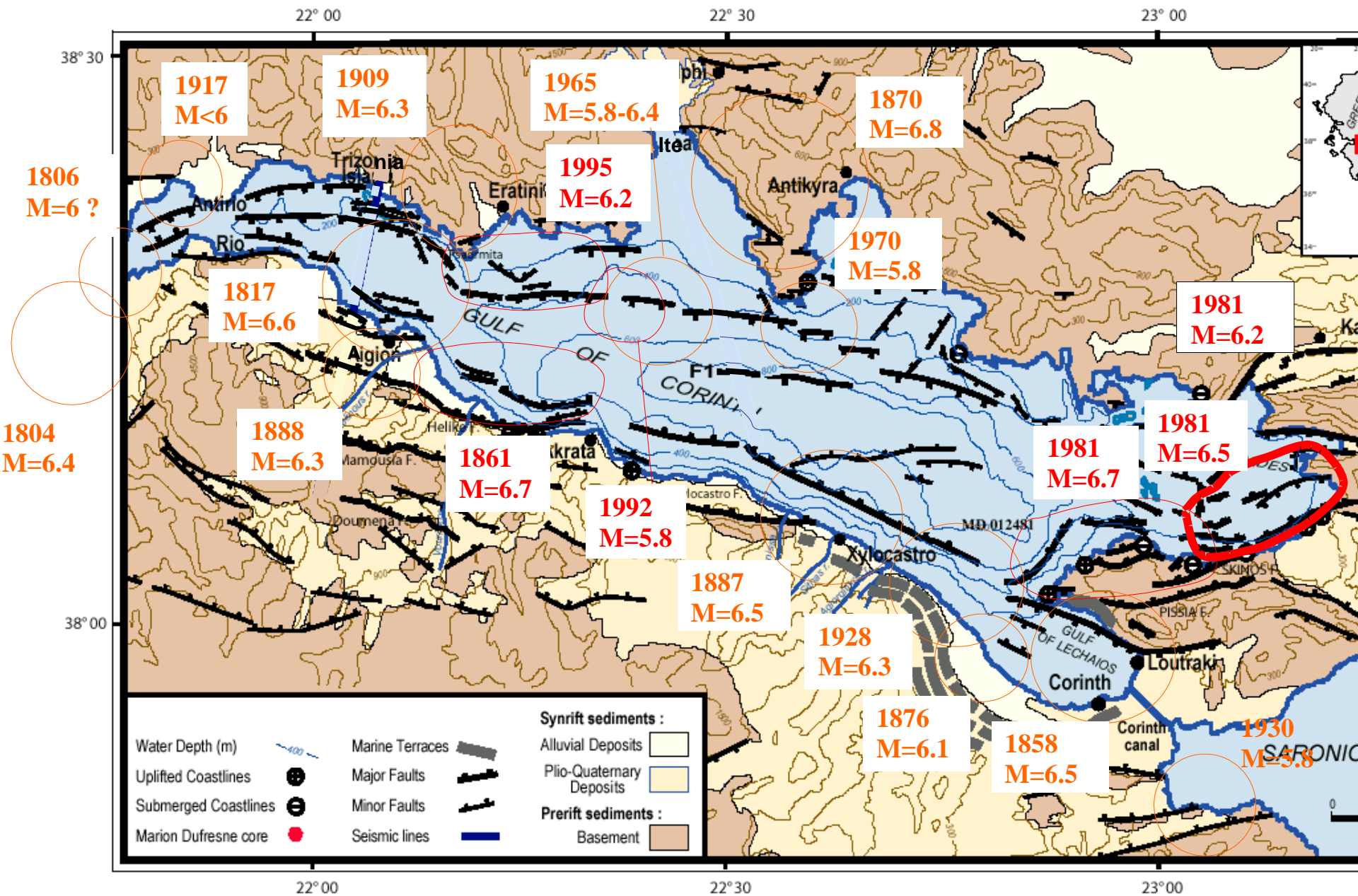


# After 3 december earthquake





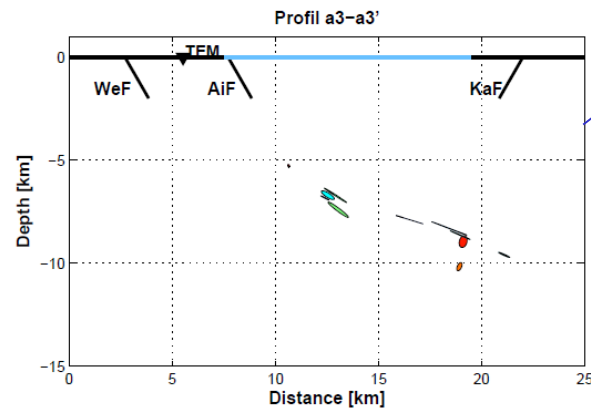
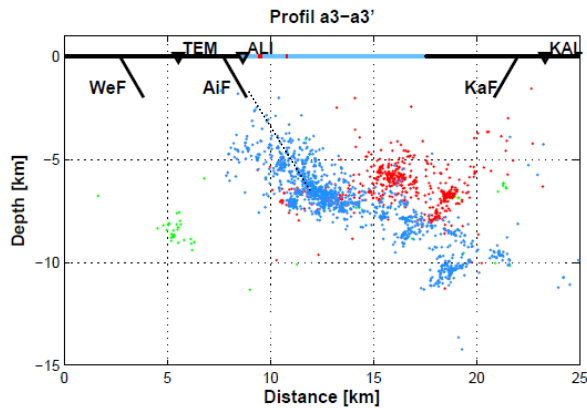
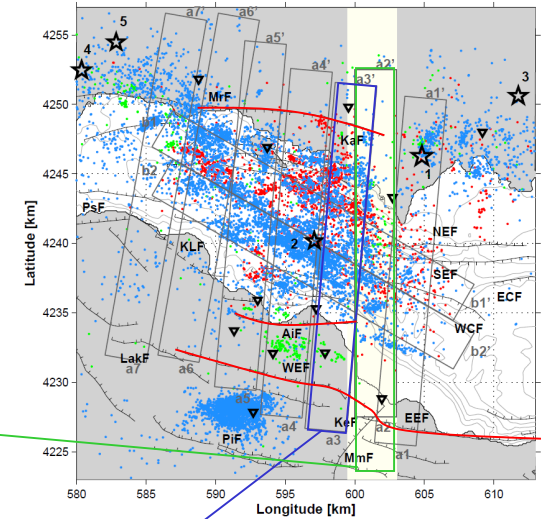
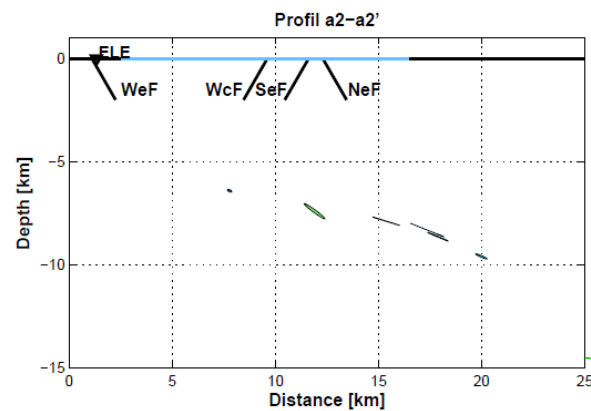
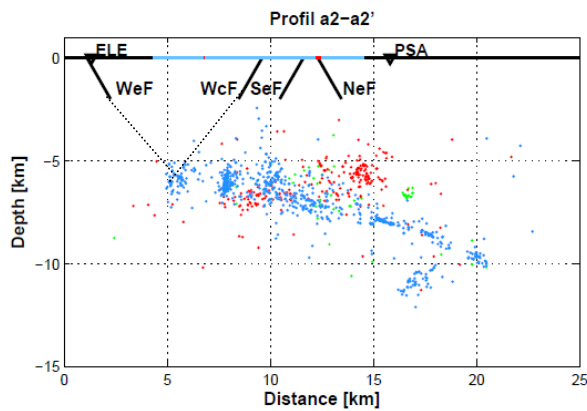
- No historical earthquakes (300 yrs)
- GPS  $>1.5$  cm/yr
- creeping and seismic

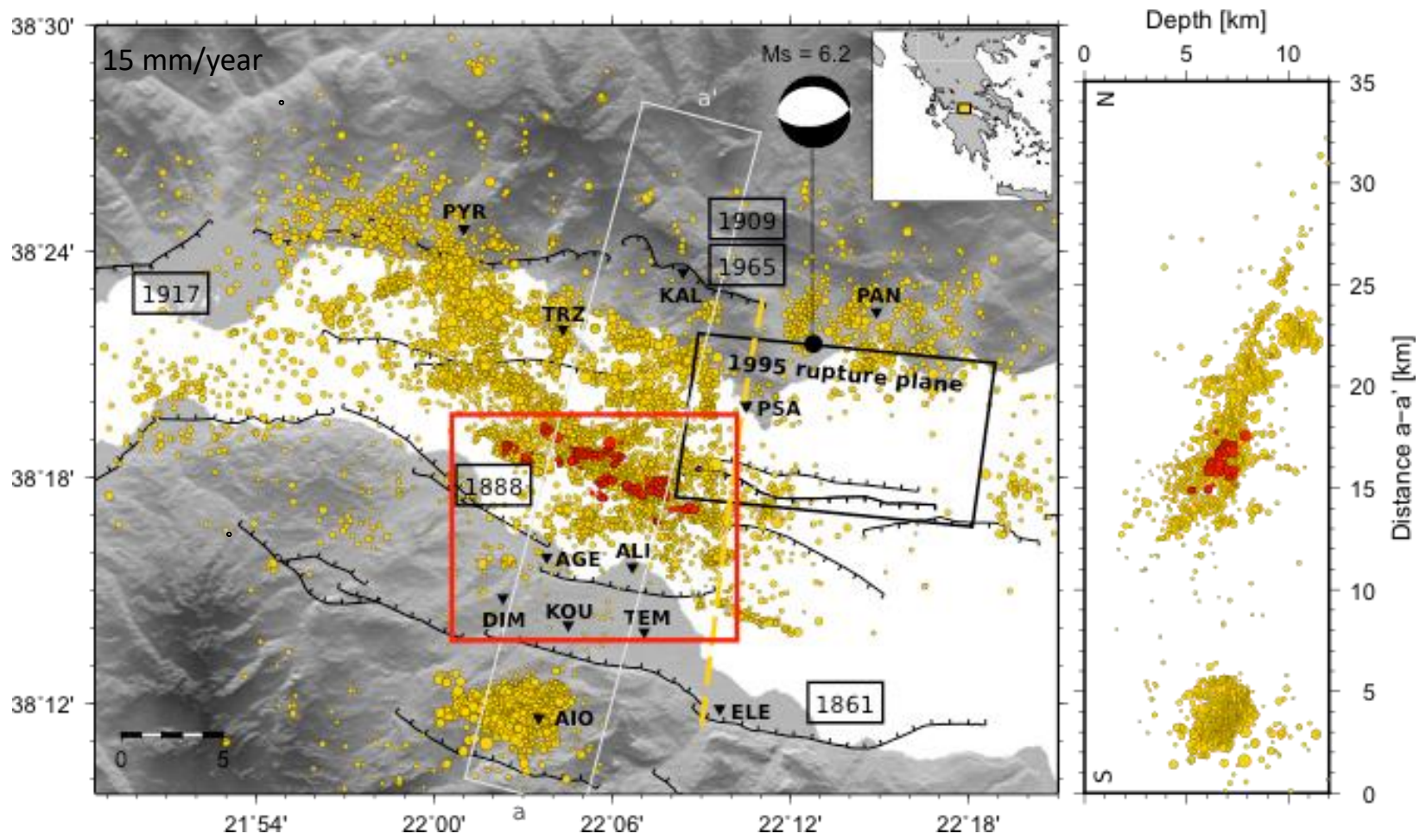


**1800-2008**

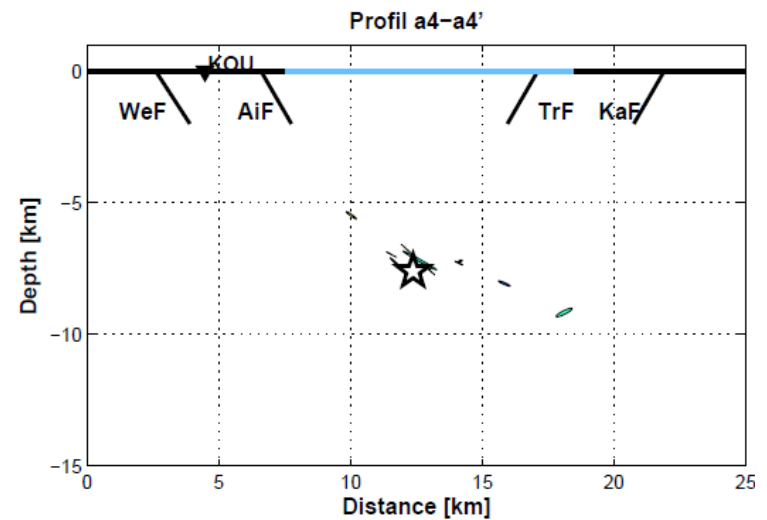
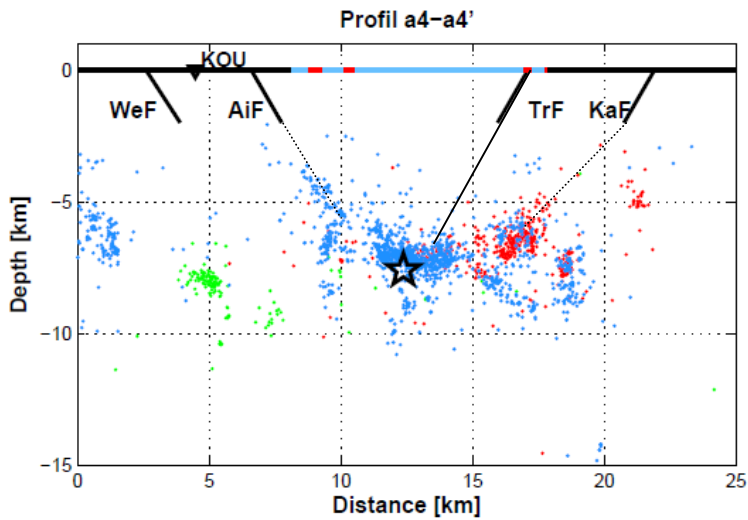
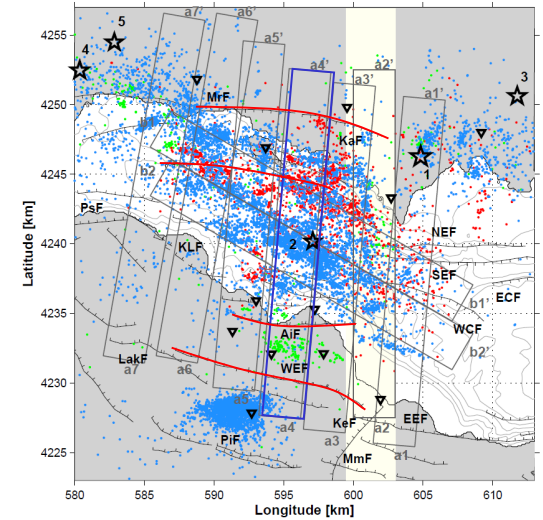
0 SSE

- Aigion et Helike faults
- antithetic faults ( Kalithea )
- low dip seismogenic layer , dipping north but:  
1995 aftershocks are shallower (red)  
Some multiplets have larger dip  
Active structures below the layer





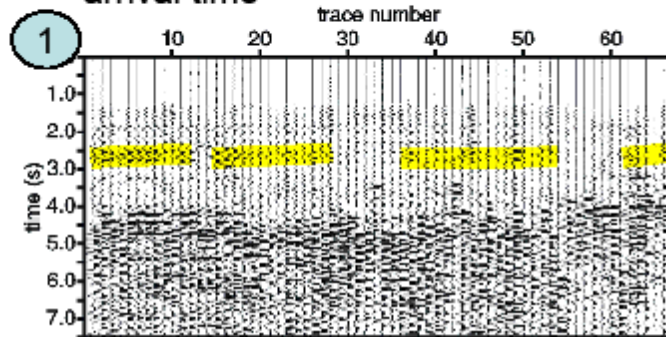
- Aigion fault: activity at 3 km, partial creep?
- West Helike Fault : silent
- Trizonia fault, antithetic: silent
- Kalithea fault: deep activity?
- + complex structure of the seismogenic « layer »



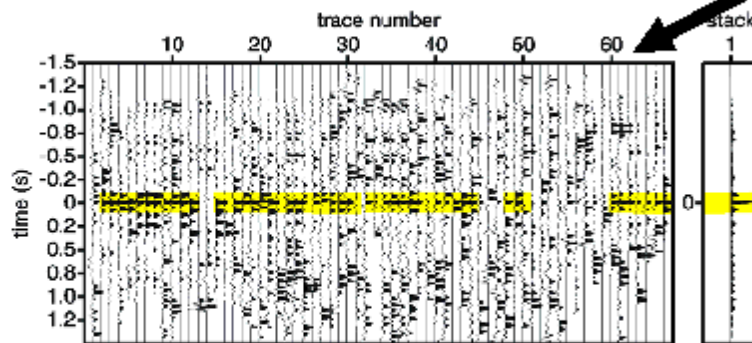
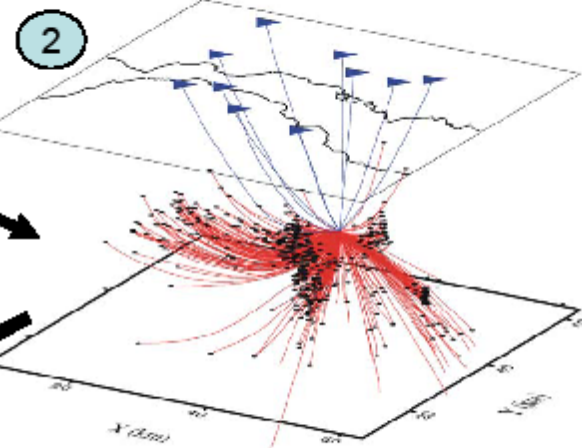


# Example on real data

Example of starting data: radial components aligned on the P-first arrival time

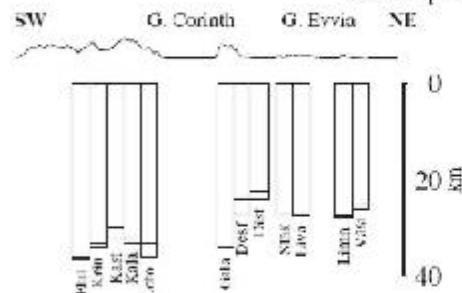
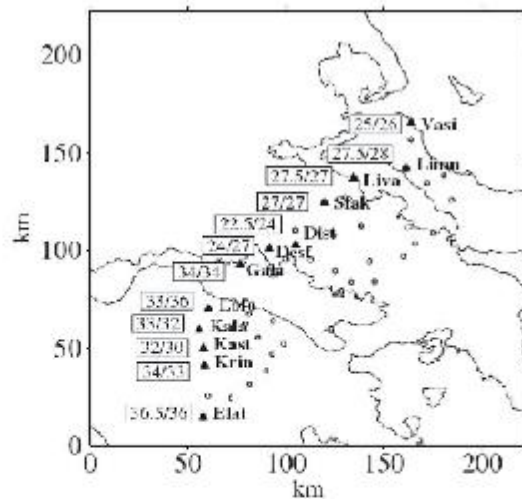
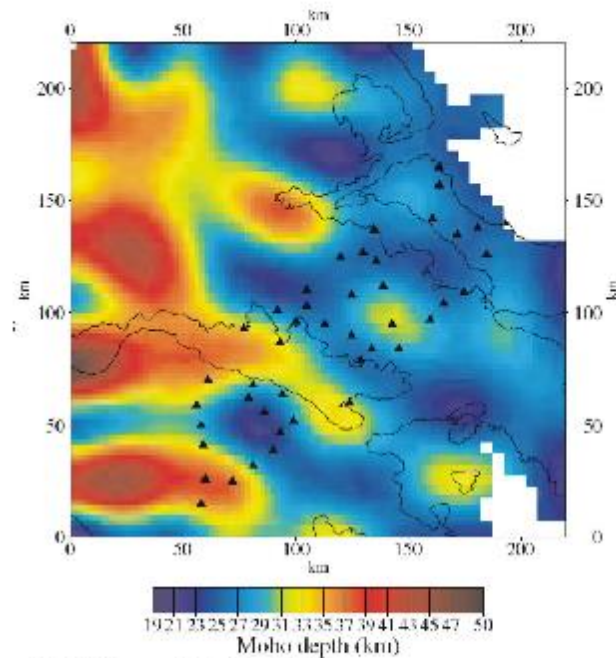


Ray tracing in the 3D velocity model for a PS converted wave at a given grid point

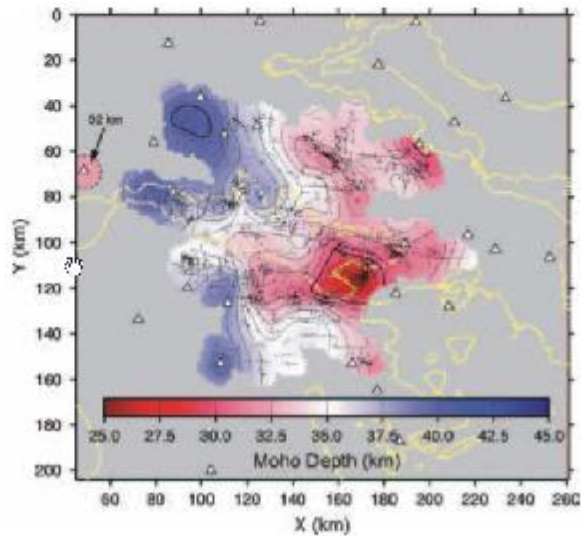


3 Stacking of shifted records: the migrated image at the grid point is computed by summing the weighted amplitudes at  $t=0$

# Inversion de données gravimétriques Comparaison à la tomographie aux fonctions récepteur

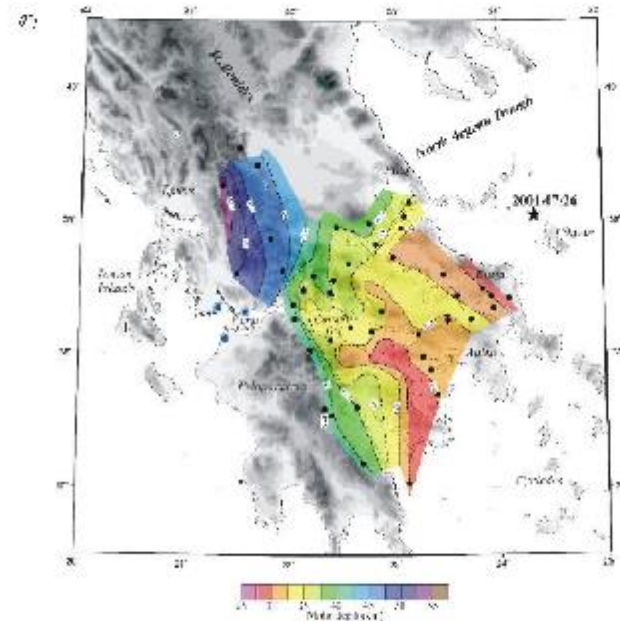


Tiberi et al.,  
Geophys..J.Int.,  
2001



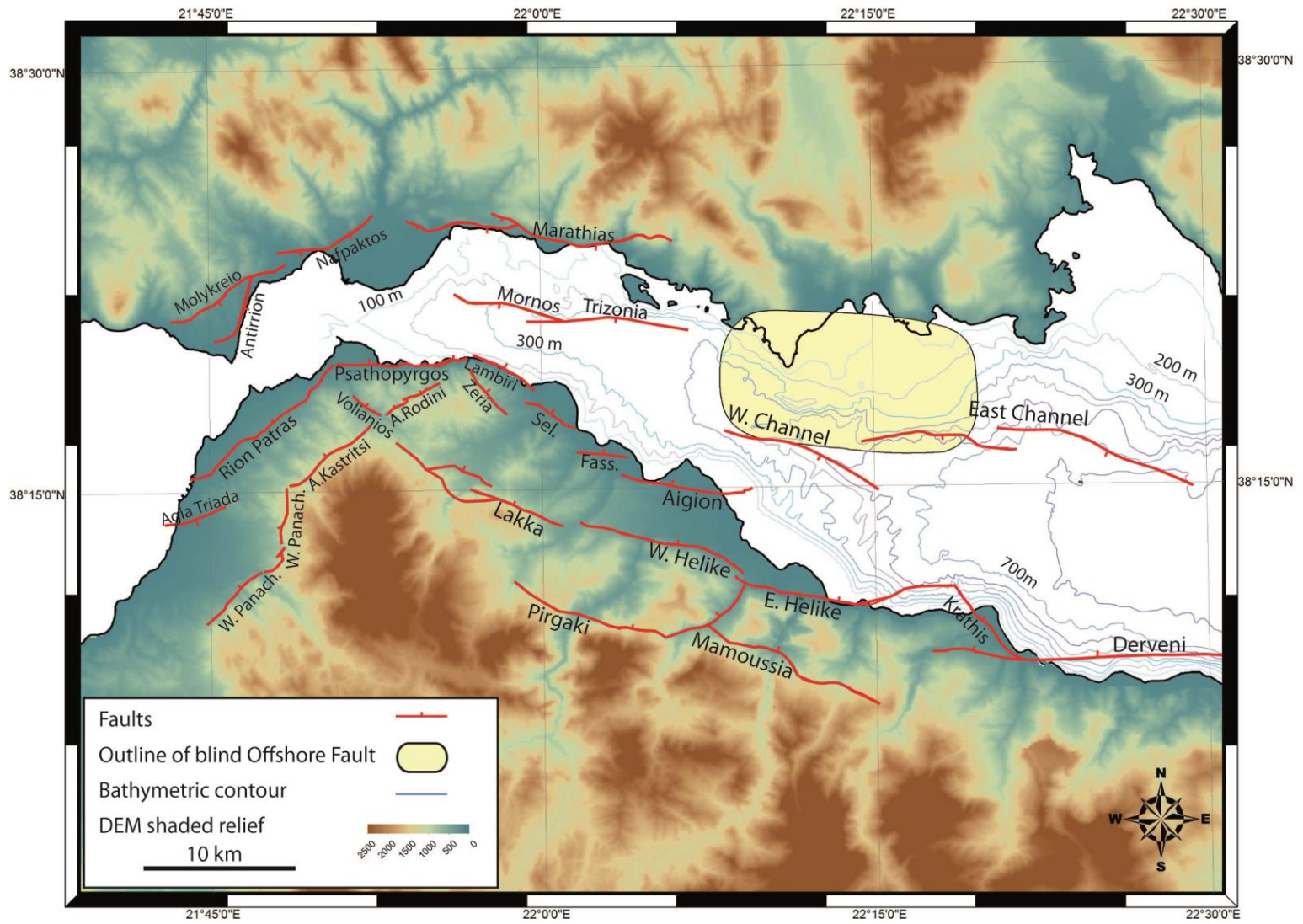
**Moho obtenu par tomographie  
2D des temps de réflexion PmP**

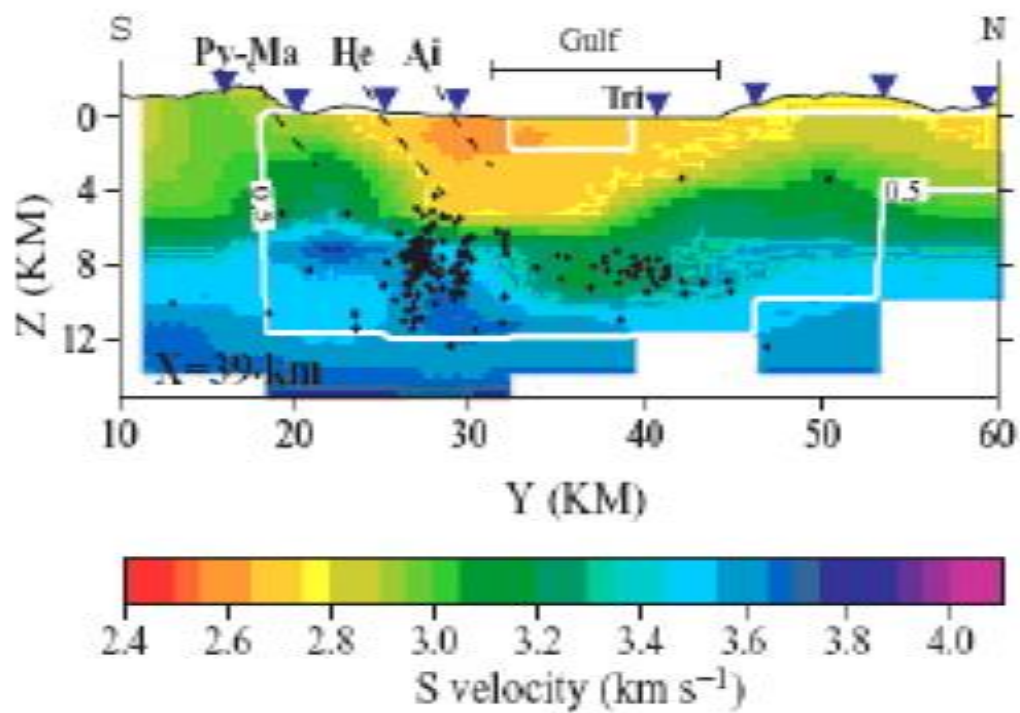
Zelt et al., Geophys.J. Int., 2005

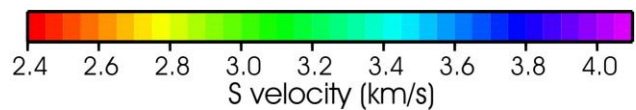
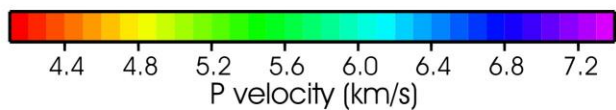
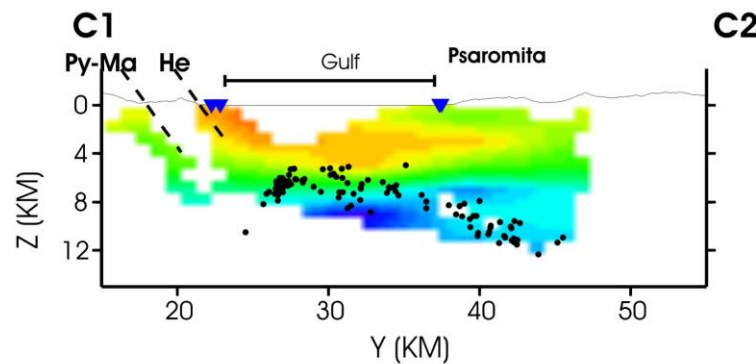
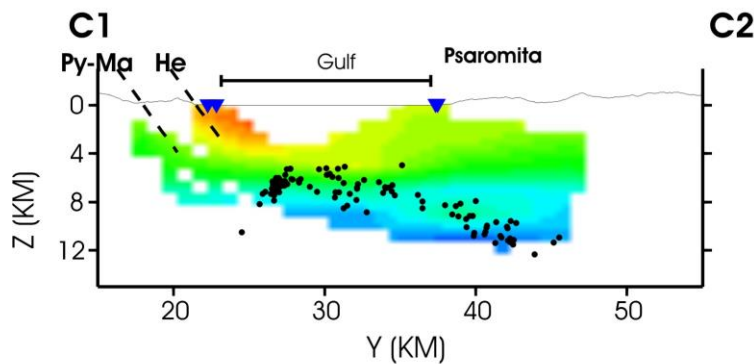
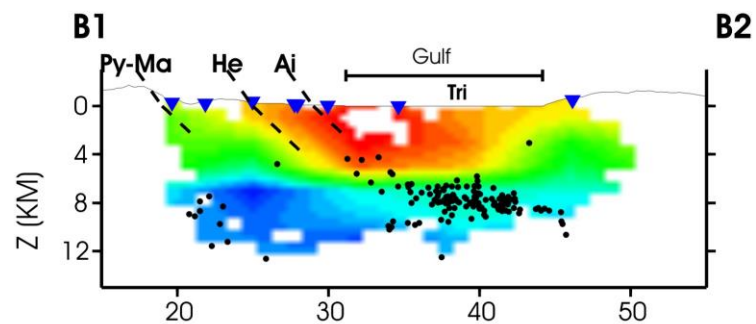
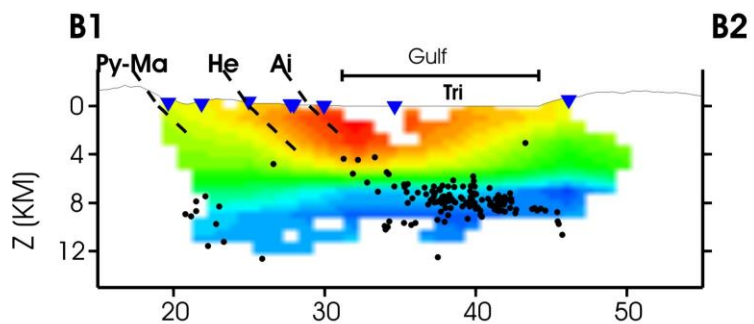
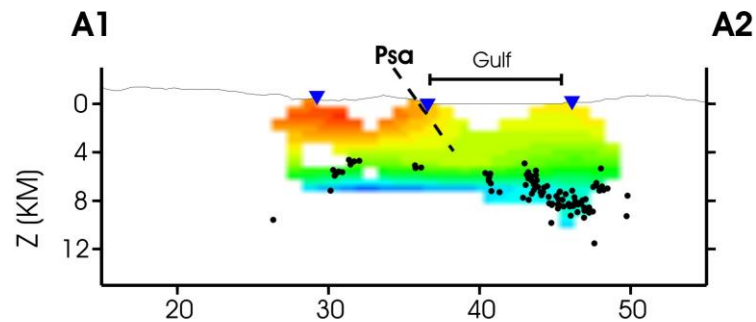
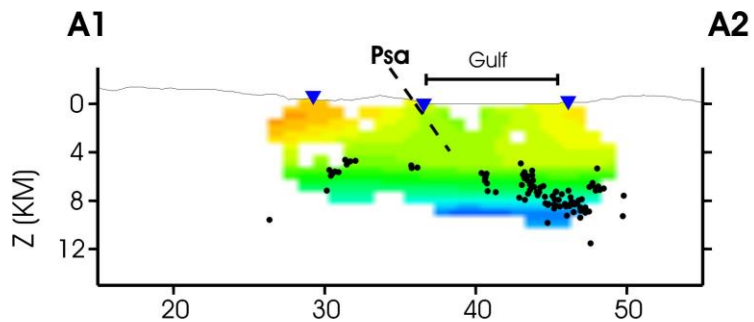


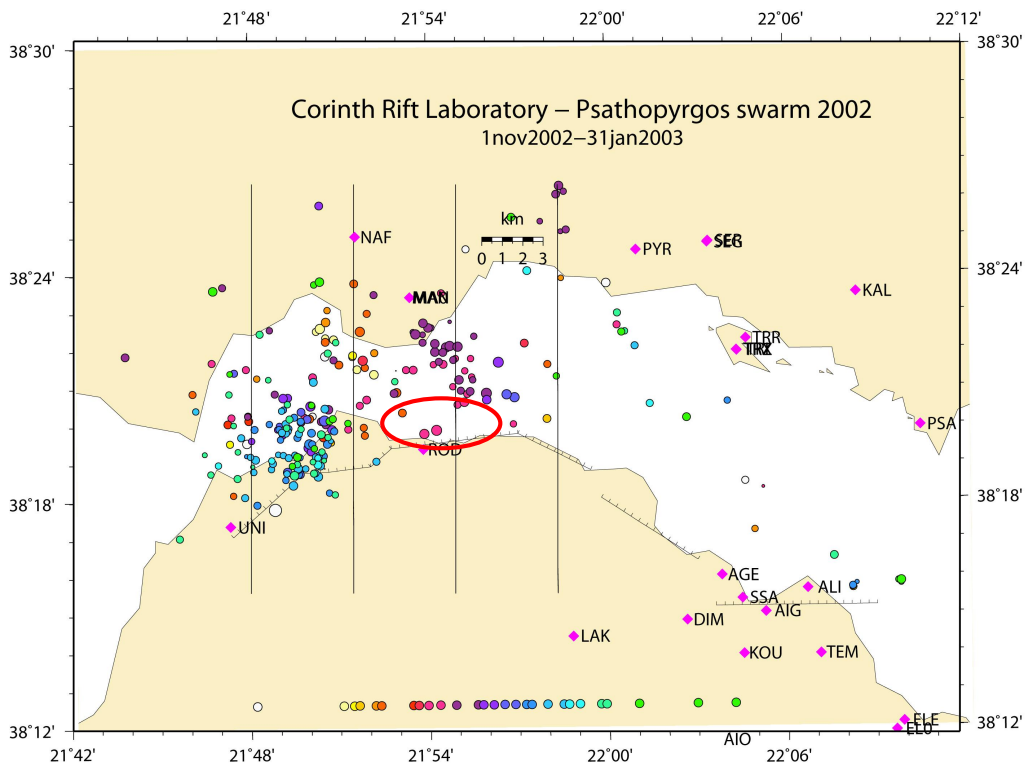
**Extension du modèle avec des Pn**

Sachpazi et al., Tectonophys. 2007

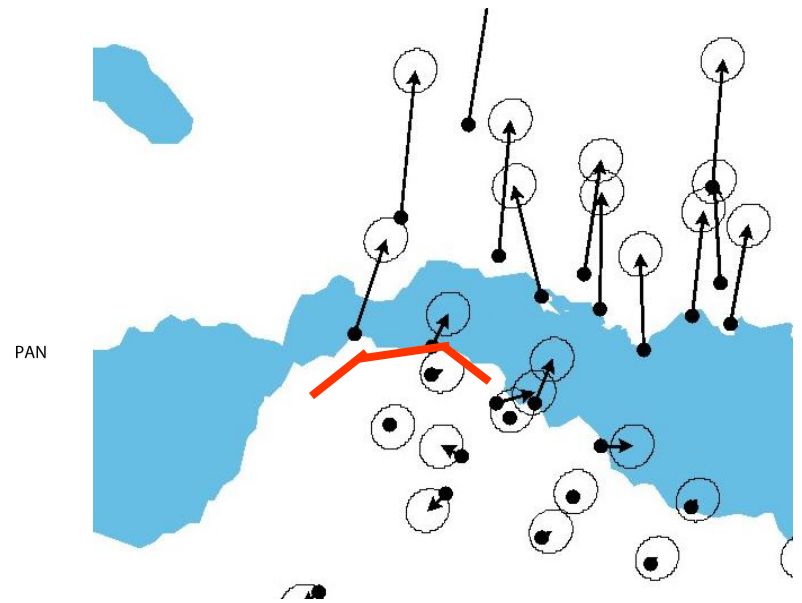








**Nov 02 – jan 03:  
Creep from seismic swarm  
and strain transient**



**Avalone et al. 2004**

**1990-2001:  
Creep from GPS  
1.5 cm/yr**